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2 Spatio-temporal trends of nitrogen deposition and climate effects on
3 *Sphagnum* productivity in European peatlands
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38 Running title: Trends of N deposition effects on *Sphagnum*

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49 Abstract

50 To quantify potential nitrogen (N) deposition impacts on peatland carbon (C) uptake, we
51 explored temporal and spatial trends in N deposition and climate impacts on the production of
52 the key peat forming functional group (*Sphagnum* mosses) across European peatlands for the
53 period 1900 - 2050. Using a modelling approach we estimated that between 1900-1950 N
54 deposition impacts remained limited irrespective of geographical position. Between 1950-
55 2000 N deposition depressed production between 0-25% relative to 1900, particularly in
56 temperate regions. Future scenarios indicate this trend will continue and become more
57 pronounced with climate warming. At the European scale, the consequences for *Sphagnum*
58 net C-uptake remained small relative to 1900 due to the low peatland cover in high-N areas.
59 The predicted impacts of likely changes in N deposition on *Sphagnum* productivity appeared
60 to be less than those of climate. Nevertheless, current critical loads for peatlands are likely to
61 hold under a future climate.

62

63 Keywords: air pollution, carbon balance, critical load, ecosystem change, peat mosses

64

65 Capsule: Temperate *Sphagnum* peatlands are vulnerable to current and future N deposition
66 and current critical loads for peatlands are likely to hold under a future climate.

67

68 Introduction

69 Deposition of reactive nitrogen (N) has increased steadily in many parts of the world since the
70 industrial revolution (Galloway *et al.* 2003). At present N deposition levels are either
71 stabilizing at a high level (Europe) or continue to increase in other parts of the world (de Vries
72 & Posch 2011). N effects are numerous, ranging from biodiversity loss to increased leaching
73 of nitrate to aquifers and lakes (e.g. Bergström & Jansson 2006). To improve risk assessments
74 and support policy decisions, critical N deposition loads have been defined for many
75 ecosystems (Bobbink & Hettelingh 2011), including *Sphagnum*-dominated peatlands. These
76 extremely nutrient-poor ecosystems are among those most sensitive to N enrichment, and
77 show changes in species composition above N deposition levels of 5-10 kg ha⁻¹ yr⁻¹ (Bobbink
78 & Hettelingh 2010). Recent work suggests that N-impacts on *Sphagnum* depends on climatic
79 factors (Heijmans *et al.* 2008; Limpens *et al.* 2011), implying critical loads may vary across
80 geographical gradients and may change with climate warming. Our study is the first attempt

81 to explore the spatial distribution of N effects on *Sphagnum* production in peatlands for past,
82 present and future.

83

84 *Sphagnum*-dominated peatlands cover large areas of the boreal-temperate zone and store a
85 substantial amount of the global soil carbon (C) pool (Rydin & Jeglum 2013). Peat mosses
86 (genus *Sphagnum*) play a key role in these ecosystems as they efficiently engineer an
87 environment that facilitates their own growth while being hostile to vascular plants (van
88 Breemen 1995). Their unique properties are both directly and indirectly responsible for the
89 long-term C sequestration of northern peatlands. *Sphagnum* directly affects C-sequestration
90 by producing litter that decays at a slower rate than vascular plants (Freeman *et al.* 2001), thus
91 forming a much greater proportion of the peat than expected from their primary production
92 (Wallén, 1992). *Sphagnum* litter/peat is also responsible for the wet and acidic environment
93 that suppresses decay, thus indirectly facilitating long-term C-sequestration. Thus, *Sphagnum*
94 production and abundance are of vital importance for the functioning and C accumulation of
95 *Sphagnum*-peatlands (Johnson & Damman 1993). The proportion of annual *Sphagnum*
96 production that is eventually integrated into long-term C storage (peat accumulation) varies
97 between 0-10%, depending on environmental conditions (Rydin & Jeglum 2013). This can
98 translate into an average carbon accumulation of circa 20 g C m⁻² yr⁻¹ (Roulet *et al.* 2007).
99 Although peatland habitats are usually resilient to environmental change (Belyea & Baird
100 2006), nutrient alterations can lead to drastic changes in the composition of the vegetation. In
101 general vascular plants benefit from the surplus N and become more dominant while the
102 growth of peat mosses decreases due to litter burial and shading (Berendse *et al.* 2001). There
103 is also experimental evidence of more direct effects of increased N availability on *Sphagnum*
104 performance, such as increased photosynthesis, P limitation and even toxicity at very high N
105 levels (Granath *et al.* 2009, Granath *et al.* 2012, Limpens and Berendse 2003). Consequently,
106 N deposition has been suggested as a potential threat to C sequestration of boreal and
107 temperate peatlands (Gunnarsson *et al.* 2008; Dise 2009). Yet, attempts to quantify this threat
108 are lacking. This is in sharp contrast to forest ecosystems where N deposition effects on C
109 sequestration have been quantified at a European scale (de Vries *et al.* 2006; de Vries &
110 Posch 2011).

111

112 A modelling approach is necessary to quantify the effects of N deposition at a larger scale,
113 and until recently, no model was available for *Sphagnum*. A recent meta-regression analysis
114 (Limpens *et al.* 2011), was an important step forward in the process of understanding and

115 quantifying the effect of N on *Sphagnum* growth in peatlands. Using results of N application
116 experiments across the northern hemisphere, the authors identified some of the variables that
117 may affect *Sphagnum* growth response to N deposition. This analysis resulted in a statistical
118 model that describes the impact of N deposition on *Sphagnum* production in interaction with
119 climatic factors. By combining this statistical model with N deposition and climate data for
120 European peatlands for the past, present and future (1900-2050) we investigated, i) how N
121 deposition has affected *Sphagnum* production, temporally and spatially across Europe, over
122 the 20th century, ii) the effect of future N deposition on *Sphagnum* production, given future
123 climatic and N deposition projections, and iii) the overall consequences for *Sphagnum*
124 production and potential C-uptake at European scale.

125
126

127 Methods and Materials

128

129 *Modelling the effect of nitrogen deposition on Sphagnum*

130 We adapted the meta-regression model of Limpens et al (2011), further referred to as
131 Limpens-Granath model, to model the effect of N deposition on *Sphagnum* productivity under
132 different climatic conditions. Our adapted model is expressed as:

133

$$134 \log_e (P_{treat}/P_{control}) = \mu + a \times N_c + b \times N_{add} + c \times T_j + d \times P_a + C + Error \quad (1)$$

135

136 where μ is the intercept, N_c is the current mean annual N deposition at the site at the time of
137 treatment, N_{add} is the annual added N at treatment plots, T_j is the mean July temperature
138 (which is strongly correlated with mean summer temperature, June-August), P_a is mean
139 annual precipitation, $P_{control}$ is the *Sphagnum* productivity at the control plots receiving only
140 N_c , and P_{treat} is the *Sphagnum* productivity at the treatment plots, receiving $N_c + N_{add}$. C and
141 a, b, c, d are regression coefficients, where C also includes the effects of phosphorus
142 application and presence of vascular plants. The *Error* term captures the unexplained
143 between-experiment variation.

144

145 Adaptations to the Limpens-Granath model included the use of more recent climate and
146 background N deposition data, and a further simplification of parameters and adaptation of N_c

147 and N_{add} to make the model suitable for comparing N effects over different time periods
148 instead of between treatments. These adaptations are discussed below.

149

150 *Parameterization*

151 The model was parameterized using the same experimental data as in Limpens et al. (2011)
152 The most important details are given in Table S1 (sources, location, species, N application
153 rate). In contrast to Limpens et al. (2011), we used more recent climate and background N
154 deposition data. For the adapted model (1) we used gridded average climate data for a 10-year
155 period encompassing the timing of the experiments based on the Climate Research Unit
156 (CRU) data base (Mitchell *et al.* 2004). Furthermore, annual total background N deposition
157 (i.e. wet plus dry) was used instead of only wet deposition as in Limpens *et al.* (2011). Most
158 N application experiments were conducted in Europe and wet and dry N depositions for the
159 years of the experimental duration were extracted from the European Monitoring and
160 Evaluation Programme (EMEP MSC-W model results,
161 http://www.emep.int/mscw/index_mscw.html). For non-European sites, N deposition was
162 retrieved from original publications. More details on N deposition and climate data can be
163 found in the methods section below. The model was fitted using a Bayesian approach in the R
164 package MCMCglmm (Hadfield 2010) to more easily include parameter uncertainties in the
165 spatial models. For parameterization, we ran the model for 140 000 iterations after a burn-in
166 of 30 000 iterations, using non-informative priors. Changing the priors did not alter the results
167 and ensures little influence of the priors on the posterior (see Limpens *et al.* (2011) for details
168 on fitting hierarchical meta-regression models accounting for within-study dependence).

169

170 The Limpens-Granath model included the effect of microhabitat (moist lawn vs. dry
171 hummocks) and an interaction between microhabitat and temperature. As the proportions of
172 microhabitats in peatlands are not well-documented, we removed the terms involving
173 microhabitat in the model, generalizing our predictions. In addition, the Limpens-Granath
174 model included variables that were manipulated in the experiments: phosphorus application
175 and presence of vascular plants (some experiments removed vascular plants). In our
176 predictions we assumed that phosphorus was not added and vascular plants were not removed.

177

178 To test if the adapted Limpens-Granath model would lead to a loss of predictive value, we
179 compared both the original and adapted models using the experiment data set of Limpens *et*
180 *al.* (2011). Our simpler model performed as well as the original model, with predictors
181 explaining 51% (present model) and 53% (Limpens-Granath model) of the variation among
182 experiment outcomes (log response ratio), respectively. Model parameters estimated for our
183 revised model are given in Table 1.

184

185 Our model was applied on a 0.5×0.5 degree longitude-latitude grid (8539 grid cells with
186 peatlands) covering Europe except for the most eastern parts where the peatland distribution is
187 not available (Ukraine, Belarus, Moldova and Greece, see Figure 2). To make spatial
188 predictions, we used July temperature, annual precipitation and annual N deposition data for
189 each grid cell. Model uncertainty (parameter uncertainty and random variation) was
190 incorporated by performing random draws of parameter estimates from the joint distribution
191 of the meta-regression model. We made 5000 random draws for predicting the outcome at
192 each grid cell and time period. From these 5000 predictions we calculated the mean and the
193 90% and 50% uncertainty intervals.

194

195 Including all model uncertainties gave wide uncertainty intervals. A large part of that
196 uncertainty comes from the unexplained between-study variation (Table 1), which is mainly a
197 result of large within-peatland variation in *Sphagnum* species and vascular plant cover. Since
198 we were not interested in the within-peatland variation but rather in the aggregated response
199 of the whole peatland to N deposition, we removed the uncertainty related to *Sphagnum*
200 species for the uncertainty graphs. To this end we first analyzed a subset of the data used in
201 Limpens *et al.* (2011) that only contains studies with more than one species (see Table S1).
202 From this data we estimated that the species component explained about 25% of between-
203 study variation. This variation was removed when prediction intervals were created for the
204 graphs presented in the supporting information (see figures in Appendix S1 in Supporting
205 Information).

206

207 To summarize the effect over Europe we calculated weighted means of predictions using
208 proportions of peatlands in each grid cell as weights. The estimation of the peatland cover is
209 described later in the methods section. Effects on carbon uptake were calculated by assuming
210 a mean annual *Sphagnum* production of 200 g dry matter m⁻² yr⁻¹. This value is derived from
211 a meta-analysis of *Sphagnum* production (Gunnarsson 2005) enhanced by the production rates

212 observed in the control plots of the N addition experiments that formed the basis of our
213 model. The $200 \text{ g m}^{-2} \text{ yr}^{-1}$ is a representative, albeit conservative (e.g. Wieder *et al.* 2010),
214 estimate of long-term production values and thus highly suitable for the 50-year time step
215 adopted in our study. To convert the dry matter production to C-uptake we assumed a dry
216 tissue carbon content of 50% (Rydin & Jeglum 2013).

217

218 *Modelled time periods and scenarios*

219 We analysed N deposition effects between 1900-1950, 1950-2000 and four different scenarios
220 for the period 2000-2050; combining (1) two N deposition scenarios (current N deposition,
221 using means of 2000-2010, and future N deposition, according to the Global Energy
222 Assessment current legislation, GEA-CLE), with (2) two climate scenarios (current climate,
223 using means of 1950-2000, and future climate, according to the A1 climate scenario). For
224 more information on these scenarios see sub-section *Nitrogen deposition 1900-2050* and
225 *Climate data 1900-2050*.

226

227 *Modelling N effects over time*

228 The response variable in the Limpens-Granath model was the ratio of *Sphagnum* productivity
229 in treatment over control, i.e. the model predicts the *change* in *Sphagnum* productivity
230 compared to a reference point. Consequently, to apply the Limpens-Granath model based on
231 N addition experiments to field conditions, we needed to set a reference point with which to
232 compare the effects of increasing N deposition. We used the year 1900 as reference point,
233 assuming that N deposition at that time reflects pre-industrial times (de Vries & Posch 2011).

234

235 N load in the Limpens-Granath model was divided over N deposition at the experimental site
236 (N current = N_c) and experimentally applied N (N_{add}). For our adapted model we also divided
237 N deposition over the terms N_c and N_{add} , where N_c presented the impact of long-term historic
238 N deposition, and N_{add} the impact of additional N deposition. We used one or two allocation
239 steps, depending on the change in N deposition between 1900 and 2050. If N deposition
240 changed less than 0.4 g m^{-2} between 1900 and 2050, no reliable predictions could be made
241 between the periods as the Limpens-Granath model assumes a minimum N_{add} of 0.4 g m^{-2} .
242 One allocation step was used when the change in mean N deposition between 1900-2050
243 exceeded $0.4 \text{ g m}^{-2} \text{ yr}^{-1}$. Then the N deposition in 1900 was allocated to N_c , and all other N to
244 N_{add} . If the change in N deposition between 1950-2000 to 2000-2050 was higher than 0.4 g m^{-2}

245 $^2 \text{ yr}^{-1}$ we used two allocation steps: N deposition in 1900 was allocated to N_c for the period
246 1950 to 2000, and the average N deposition between 1950 and 2000 was allocated to N_c for
247 the period 2000-2050. In the latter case, the outcome was multiplied with the first outcome
248 (effect 1900 to 1950-2000) to keep comparisons to 1900.

249

250 *Model evaluation*

251 The sensitivity of our model was evaluated in two ways; 1) to investigate the effect of
252 potential uncertainty of our modelled predictors we ran the model with an uncertainty
253 (standard deviation set to 5% of the value) around modelled predictor values (N deposition,
254 temperature and precipitation), and, 2) to test the effect of different ways of allocating N
255 deposition over N_c and N_{add} . In addition, we explored an alternative model where the
256 parameters N_c and N_{add} were merged into a single parameter (total N addition, N_{c+add}). The
257 test of this alternative model was also restricted to the $0.4 \text{ g m}^{-2} \text{ yr}^{-1}$ minimum increase since
258 1900.

259

260 *Nitrogen deposition 1900-2050*

261 Oxidised and reduced N deposition were calculated with the atmospheric transport model of
262 EMEP/MSC-W (Simpson et al. 2012). Historic NO_x and NH_3 emissions were taken from
263 Lamarque *et al.* (2010). Predictions for the period 2000-2050 were based on a (1) constant N
264 deposition scenario, i.e. fixing N deposition to the current levels, here defined as the mean for
265 the period 2000-2010, and, (2) Global Energy Assessment (GEA) current legislation scenario
266 (GEA-CLE). The GEA-CLE deposition scenario assumes full implementation of all current
267 and planned air pollution legislation world-wide until 2030.

268

269 *Climate data 1900-2050*

270 Historical meteorological data were taken from a high resolution European data base
271 (Mitchell *et al.* 2004) that contains monthly values of temperature, precipitation and
272 cloudiness for the years 1901–2000 for land-based grid-cells west of 32°E of size $10^\circ \times 10'$
273 (approx. $15 \text{ km} \times 18 \text{ km}$ in central Europe). In the simulations, 10-year averages were taken
274 centred around 1900, 1910, ..., 1990 (1900 is the average of 1901–1905; 1910 of 1906–1915
275 etc.) to smooth the climate pattern. For future values we used two climate scenarios: (1)
276 constant climate after 1990 (using the average of the 1961-1990 period); (2) from 2010
277 onwards the A1 storyline from the Special Report on Emission Scenarios (SRES) with its
278 related climate scenario (Nakićenović & Swart 2000). This scenario, which is from the IPCC

279 scenario family with the highest emissions, reflects a future world with globalization and
280 rapid economic growth, low population growth, and the rapid introduction of new and more
281 efficient technologies everywhere. The A1 scenario was chosen in this study as the most
282 likely future scenario, considering that current CO₂ emission trends exceed this most
283 pessimistic IPCC scenario (Raupach *et al.* 2007). Mitchell *et al.* (2004) used results from the
284 HADCM3 General Circulation Model to derive monthly meteorological data for the years
285 2001–2100 from this (and several other) scenario, which we again averaged over 10-year
286 periods for use in the simulations. For the transition period, 1990–2010, data were
287 interpolated linearly. Under the A1 climate scenario, temperatures over Europe increase by
288 about 3±1°C between 1990 and 2050, whereas median precipitation hardly changes, with
289 precipitation at southern sites decreasing and at northern sites increasing compared to 1990
290 (see also de Vries & Posch 2011).

291

292 *Spatial distribution of peatlands in Europe*

293 We improved the estimation of the coverage of open *Sphagnum* dominated peatlands by
294 integrating the CORINE (CO-ordination of INformation on the Environment) land cover
295 2006 of the European Environmental Agency (EEA) with the Forest Map of the Joint
296 Research Centre (JRC) of the European Commission. The CORINE land cover is derived
297 from high resolution satellite data (e.g. Landsat-TM) by computer assisted visual
298 interpretation in combination with ancillary data. The minimum mapping unit is 25 ha and for
299 line elements the minimum width is 100 m. Büttner and Maucha (2006) validated the
300 CLC2000 and found a total reliability of CLC2000 was 87.0 ± 0.8%. The JRC Forest Cover
301 Map is a 25 m spatial resolution raster Pan-European Forest / Non Forest Map with target
302 years 2000 and 2006 derived from high resolution satellite data such as LISS III and SPOT
303 4/5 imagery. In the JRC forest map, forests are defined as vegetation dominated by trees, with
304 canopy closure more than 30% and trees higher than 5 m. For its integration with CORINE,
305 the JRC Forest Map was resampled to a 100 m grid (same resolution as CORINE) using a
306 majority filter.

307

308 From CORINE land cover we selected the land cover class 4.1.2 Peat bogs, which has been
309 defined as: peatland consisting mainly of decomposed moss and vegetable matter which
310 might be exploited (Bossard *et al.* 2000). It includes minerogenic peatlands (fed by ground
311 water) with mosses (mostly *Sphagnum*), ombrogenic peatlands (fed only by precipitation and
312 dominated by *Sphagnum*) such as boreal bogs, boreal bogs with reticulated structure (aapa),

313 blanket bogs, and fossil arctic peat bogs (palsa). The integrated map of JRC forest and
314 CORINE land cover (2006), focusing on non-forested peatlands, was analysed over European
315 sites where N application experiments used in the meta-analysis were performed, and over
316 other sites where *Sphagnum* production had been recorded in the literature. Our combined
317 map showed an improvement over CORINE 2006 concerning open, *Sphagnum* dominated
318 peatlands.

319

320 The gridded peatland distribution map was first re-projected and then overlaid with the
321 gridded climate/N deposition map (0.5×0.5 longitude-latitude grid) to calculate the proportion
322 of peatland per cell. This was done using the R packages *raster* (Hijmans & van Etten 2012)
323 and *rgdal* (Keitt *et al.* 2012). All maps are shown with the EPSG:3035 projection.

324

325

326 Results

327 *N deposition trend and peatland cover*

328 Estimated N deposition varied over time and space. At the beginning of the 1900s, N
329 deposition was generally low, except for a region in central Europe and the southern British
330 Isles. Between 1900 and 1950 N deposition increased slightly, but not more than $0.4 \text{ g m}^{-2} \text{ yr}^{-1}$
331 on average over the whole period (Fig 1a). From 1950 onwards, N deposition rose quickly,
332 reaching peak values around 1980-1990, after which it decreased slightly and levelled off to
333 the present levels in many areas. The total change in N deposition exceeded $0.4 \text{ g m}^{-2} \text{ yr}^{-1}$ in
334 most areas (Fig 1b). Predictions for 2000-2050 under the current legislation N emission
335 scenario (GEA-CLE) vary per region, with N deposition increasing or decreasing, depending
336 on location (Fig 1c). In the constant (mean 2000-2010) scenario, most areas showed a minor
337 decrease compared to the 1950-2000 period (not shown).

338

339 Combining CORINE land cover data and the JRC forest map, we estimated open/semi-open
340 *Sphagnum*-dominated peatland cover in Europe to be 10.5 million ha. About 20% of these
341 peatlands are defined as forested. The highest cover of peatlands is found in Fennoscandia and
342 the British Isles (Fig 2).

343

344 *Spatial predictions for Sphagnum productivity*

Trends of N deposition effects on Sphagnum

345 For the period 1900-1950 we could not make reliable predictions for changes in *Sphagnum*
346 productivity as the change in N deposition remained below 0.4 g m^{-2} (the minimum amount
347 needed for modelling predictions, see Methods) for all regions. For 1950-2000, large areas of
348 Europe did show a sufficient increase in N deposition for making predictions. This area
349 includes approximately 1 million hectares of *Sphagnum* dominated peatlands. Over this
350 period, our model estimated that N deposition depressed *Sphagnum* production between
351 between 0 and 25% in western, central and eastern Europe (Fig 3a), but increased production
352 in south-eastern Sweden. N deposition in parts of Ireland and northern Fennoscandia did not
353 change enough to allow reliable predictions. When summarized over all predictions and
354 taking peatland cover into account, the overall N effect on *Sphagnum* production, and
355 consequently on the *Sphagnum* C balance of the living parts, was close to zero (Table 2).

356

357 Over 2000-2050 N deposition is predicted to depress *Sphagnum* production further under both
358 the constant N emission scenario (using average N deposition of 2000-2010; Fig 3c) and the
359 current legislation N emission scenario (GEA-CLE deposition scenario; Fig 3b). Without a
360 change in climate however, predictions are similar to 1950-2000 (Fig. 3d). The strongest
361 overall reduction of *Sphagnum* production occurs under the constant N deposition and A1
362 climate scenario (Table 2, Fig 3c) and the weakest is under the reduced N deposition and
363 constant climate scenario (Table 2, Fig 3d). The predicted impacts of N deposition appeared
364 to be less (on average 3 percentage points difference between constant and reduced N
365 deposition) than those of climate change (on average 7-8 percentage points difference
366 between constant climate and climate change according to A1; see table 2). The consequences
367 of the different scenarios for C uptake show large spatial variation, but little-no difference
368 when upscaled to Europe (Table 2).

369

Model uncertainty and sensitivity

371 Including all model uncertainties gave wide uncertainty intervals even when correcting for
372 contribution of within peatland uncertainty. Hence parameter uncertainty is the largest
373 contributor to the overall uncertainty (Supporting information, Fig. S1, S2).

374

375 When looking at the sensitivity of the model, a standard deviation of 5% for a predictor value
376 (e.g. an N deposition value of $1.0 \text{ g m}^{-2} \text{ yr}^{-1}$ is assumed to be estimated with a standard
377 deviation of $0.05 \text{ g m}^{-2} \text{ yr}^{-1}$) did not change our predictions substantially. If we did not use

378 allocation steps for N_c , but only one fixed N_c (N deposition in 1900) for all time periods, our
379 future predictions (2000-2050) suggested less negative effects of N deposition. Also, the
380 model where N_c and N_{add} were merged into one parameter (N_{c+add}) did not diverge much from
381 our original model. In this case, however, the model produced more extreme predictions near
382 the boundaries of model space and stronger negative effects were predicted at lower N
383 deposition levels because a larger proportion of N deposition was allocated to N_c . The
384 instability of this model (N_{c+add}) makes it less useful despite its simplicity and thus the results
385 from the two-step allocation model are the only ones presented here.

386

387 Discussion

388

389 *Nitrogen deposition impacts*

390 For the second half of the 20th century, our model suggests that N deposition will depress
391 *Sphagnum* production in many areas by 0-25% relative to 1900. This reduction nearly equals
392 the year-to-year variation due to weather conditions observed in contemporary *Sphagnum*
393 production studies (e.g. Vitt 1990). These modelled trends correspond to independent field
394 observations of peat accumulation (Gunnarsson *et al.* 2008), vegetation changes (Kapfer *et al.*
395 2011) and *Sphagnum* N concentration (Malmer & Wallén 2005).

396

397 The N deposition impact varied spatially, with N deposition depressing *Sphagnum* production
398 mostly in peatlands approximately south of 60° latitude, with the exception of Eastern
399 Sweden. North of 60° latitude, changes in *Sphagnum* production could generally not be
400 modelled because N deposition changed less than 0.4 g m⁻² yr⁻¹ relative to 1900. However
401 based on our model, it is likely that N deposition had little effect on *Sphagnum* production and
402 probably stimulated growth. Unfortunately, there are no long-term data sets on *Sphagnum*
403 production that we can use for direct validation of our model results, restricting validation to
404 comparisons of patterns observed in vegetation composition and litter production which may
405 be viewed as indirect proxies for *Sphagnum* production. Below we compare our results with
406 patterns observed in vegetation composition and litter production in bogs. Our spatial
407 distribution of N deposition impacts corresponds with that observed in studies focussing on
408 peatland vegetation composition. For example, in Southern Sweden (60° latitude) where our
409 model estimates a net positive effect over the 20th century, Gunnarsson *et al.* (2000) found no
410 evidence of N-induced changes in vegetation composition, whereas in England where our

411 model estimates negative effects, Chapman & Rose (1991) reported negative effects on
412 vegetation, notably on *Sphagnum* cover. Another interesting example is the peatland Store
413 mosse in SW Sweden. Malmer & Wallén (2004) estimated that about 50% less litter
414 (*Sphagnum* and vascular plants) was produced in the 20th century compared to the 19th
415 century. A direct comparison with our estimate (20-25% reduction) is not straight forward
416 since we focus on *Sphagnum* and single out the effect of N deposition. In their study, Malmer
417 & Wallén (2004) acknowledge that N deposition is a likely important factor causing the
418 decrease of litter from *Sphagnum*, together with general drier conditions. Although being a
419 single peatland, this gives some quantitative support for our predictions that needs to be
420 scrutinized further. The scarcity of suitable validation data highlights the need of long-term
421 monitoring of processes, such as productivity, instead of only state variables, such as
422 vegetation composition.

423

424 Our future scenarios suggest that N deposition impacts predicted for the period between 1950-
425 2000 will continue in the 21st century, becoming more pronounced with climate warming. At
426 the European scale, however, the net N deposition impact on carbon uptake by *Sphagnum*
427 remains small due to the low peatland cover in high N deposition regions: ranging between -6
428 to -15 g C m⁻² yr⁻¹ relative to 1900 for the period 1950-2050, depending on N deposition
429 scenario and climate warming. However, at the peatland scale N deposition could depress
430 *Sphagnum* C uptake up to 50-60 g C m⁻² yr⁻¹. This is twice as large as recent estimates of
431 annual C sequestration in northern peatlands based on long-term eddy covariance monitoring
432 (Yu *et al.* 2011). Assuming 10% of the *Sphagnum* production eventually enters long-term
433 storage in the anoxic zone (Rydin & Jeglum 2013), then long-term C sequestration of
434 *Sphagnum* would be reduced by roughly 5 g C m⁻² yr⁻¹ in the most negatively affected
435 peatlands. If this long term decrease in C sequestration by *Sphagnum* can be offset by
436 increased production of vascular plant litter is highly uncertain as non-woody vascular plant
437 litter is readily decomposed and is less likely to enter into the anoxic zone as peat (Rydin &
438 Jeglum 2013). Moreover, reduced *Sphagnum* growth may decrease the self-regulation
439 capacity and resilience of these ecosystems to further disturbances such as changes in climate
440 or land-use.

441 Regions that would be affected most in the future are SW Sweden, Ireland, England, NW
442 Germany, Netherlands and the Alps. Here the N deposition rates remain too high to halt
443 negative N impacts, even in the best-case N deposition scenario without climate change. This
444 suggests that the oceanic-temperate peatlands in Europe will continue to suffer under the

445 predicted N deposition levels. The area estimated to be affected is likely underestimated in
446 southern Europe because of the many small peatlands which may fulfil the CORINE size
447 criteria. Many of these temperate peatlands are protected nature reserves and may continue to
448 face lower *Sphagnum* production, invasion of vascular plants (Berendse *et al.* 2001;
449 Tomassen *et al.* 2004) and continued losses of DON to groundwater and surface waters
450 (Bragazza & Limpens 2004).

451

452 *Prediction uncertainties*

453 Despite the numerous N application studies carried out in peatlands, making quantitative
454 estimates remains a challenge. In our study we did not include any uncertainties in climate
455 change (e.g. Ruete *et al.* 2012) or N deposition scenarios. Instead we explored one model for
456 climate change (A1) and N deposition (GEA-CLE) and compared them with a constant
457 scenario, including all combinations. Sensitivity analyses showed that uncertainties in these
458 variables did not much affect predictions. Likewise, uncertainty in peatland cover was not of
459 major influence and point predictions were mostly influenced by model parameter
460 uncertainties.

461

462 Predictions presented here are associated with considerable uncertainties, reflecting both
463 parameter uncertainty, but also unknown variables affecting the effect of N deposition. It is
464 reasonable to assume that a large proportion of the random variation among N application
465 experiments is mostly associated with unknown microhabitat and species-specific responses
466 and factors that could not be included here because of the limited amount of data. For
467 example, different effects of N form (dry versus wet) and vascular plant cover. We argue that
468 at the peatland scale our average effects are reasonable and realistic, although we should be
469 cautious in our interpretation of the large effects reported (e.g. future *Sphagnum* production
470 reduction by 50% in some regions). By using higher temperatures we are making predictions
471 near the limits of model space. Nevertheless, these results suggest that interactions between
472 the effect of N and climate need to be addressed in future research.

473

474 Another potential uncertainty resides in the assumption of our model that short term
475 experiments (here 1-6 years) can be extrapolated over a longer time span. This assumption
476 was supported by time series analyses in Limpens *et al.* (2011) which did not detect an effect
477 of experiment length on the effect of N application. Accumulated N deposition (i.e. total N
478 load) may, however, be as important as deposition rates since low deposition levels may over

479 time reach critical loads. We were unable to predict N deposition effects for a large part of
480 northern Europe, because N deposition change is too low to allow reliable modelling. N
481 experiments in peatlands have so far focused on rather large N loads, preventing accurate
482 assessment of the impact of a more realistic low, but chronic N deposition rates. Evidence
483 from other ecosystems, such as forest, suggests that low but chronic N loads induced by
484 experimental N fertilization of forests in Sweden, Finland and the USA for a period of about
485 10- 30 year causes enhanced carbon sequestration by increased tree growth without signs of N
486 saturation (Högberg et al., 2006; Pregitzer et al. 2007; Hyvönen et al., 2008). However, at
487 elevated N deposition, N saturation, occurrence of soil acidification and the elevated
488 incidence of pest and diseases may lead to adverse impacts on forest growth (Magill *et al.*
489 2004). These results imply that experiments with high N loads in forests tend to overestimate
490 adverse impacts as compared to low chronic loads. Our model suggest generally positive
491 effects of low N levels on *Sphagnum* production, which is supported by the field observations
492 in areas that received low N deposition ($\sim 0.5 \text{ g m}^{-2} \text{ yr}^{-1}$) during an extensive time period (e.g.
493 50 years - Gunnarsson *et al.* 2000, 34 years - Vitt *et al.* 2003). While intact *Sphagnum*-
494 dominated peatlands show a high resistance to disturbances such as moderate drainage or low-
495 intensity fire, the disruption of keystone species such as *Sphagnum* may reduce the resilience
496 of peatlands and cascade into disequilibrium or ecosystem change (Luo & Weng 2011).
497 Nevertheless, in the case of N deposition, it is also possible that thresholds do not exist, or are
498 below a detection limit, and the ecosystem is changing in a continuous manner (Payne *et al.*
499 2013).

500

501 *Evaluating critical loads*

502 Overall, our analyses support the critical load range of $0.5 \text{ to } 1.0 \text{ g m}^{-2} \text{ yr}^{-1}$ for peatlands
503 (Bobbink & Hettelingh 2011). The lower bound of that range suggests that negative N-
504 deposition impacts already occur at a very small increase ($0.2\text{-}0.3 \text{ g m}^{-2} \text{ yr}^{-1}$ assuming a pre-
505 industrial N deposition of $0.2\text{-}0.3 \text{ g m}^{-2} \text{ yr}^{-1}$) of N deposition. In some areas, *Sphagnum*
506 production showed reductions at the minimum N deposition increase that could be reliably
507 modelled in our study ($0.4 \text{ g m}^{-1} \text{ yr}^{-1}$), while in other areas a greater increase of N deposition
508 (up to $1 \text{ g m}^{-1} \text{ yr}^{-1}$) was needed to achieve a significant production decrease. Overall, our
509 results suggest the current critical load holds for the range of climatic conditions predicted for
510 2050.

511

512 *Conclusions*

513 Our modelling analysis suggests that the negative impact of N deposition was rather limited
514 throughout the 1900s, but the negative effects observed in the end of the 20th century (up to
515 25% production reduction) are likely to continue or even increase in the 21st century.
516 Peatland-rich areas in Fennoscandia are, however, less impacted and *Sphagnum* production
517 may even be temporarily stimulated in some areas. Our results stress the vulnerability of
518 temperate peatlands to N deposition impacts and highlight the need for more studies on the
519 effect of small but chronic elevations of N deposition. Furthermore, our results suggest the
520 current critical loads for peatlands are also valid under predicted climate warming.

521

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529

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684 Table 1. Results of the Hierarchical Bayes Linear Models (HBLM) with \log_e response ratio.
 685 Negative coefficients indicate that an increase in the predictor depresses the response of
 686 *Sphagnum* to adding N. Categorical levels are compared to the intercept which is set to
 687 without P addition and without vascular plants. Upper and lower 95% credible intervals are
 688 given. Residual heterogeneity represents between study variation not explained by predictors.
 689 Of the total between study variation, 51% was explained by the included predictors. N=107.

	coefficients	lower-95%	upper-95%
Intercept	1.491	0.847	2.124
Nitrogen (N) application rate	-0.032	-0.056	-0.009
Background N deposition	-0.200	-0.393	-0.009
Mean July temperature	-0.049	-0.083	-0.014
Mean annual precipitation	-0.000372	-0.000741	-0.000013
Presence of vascular plants	-0.377	-0.585	-0.181
Phosphorus application	0.225	0.023	0.432
Residual heterogeneity (τ^2)	0.062	0.032	0.097

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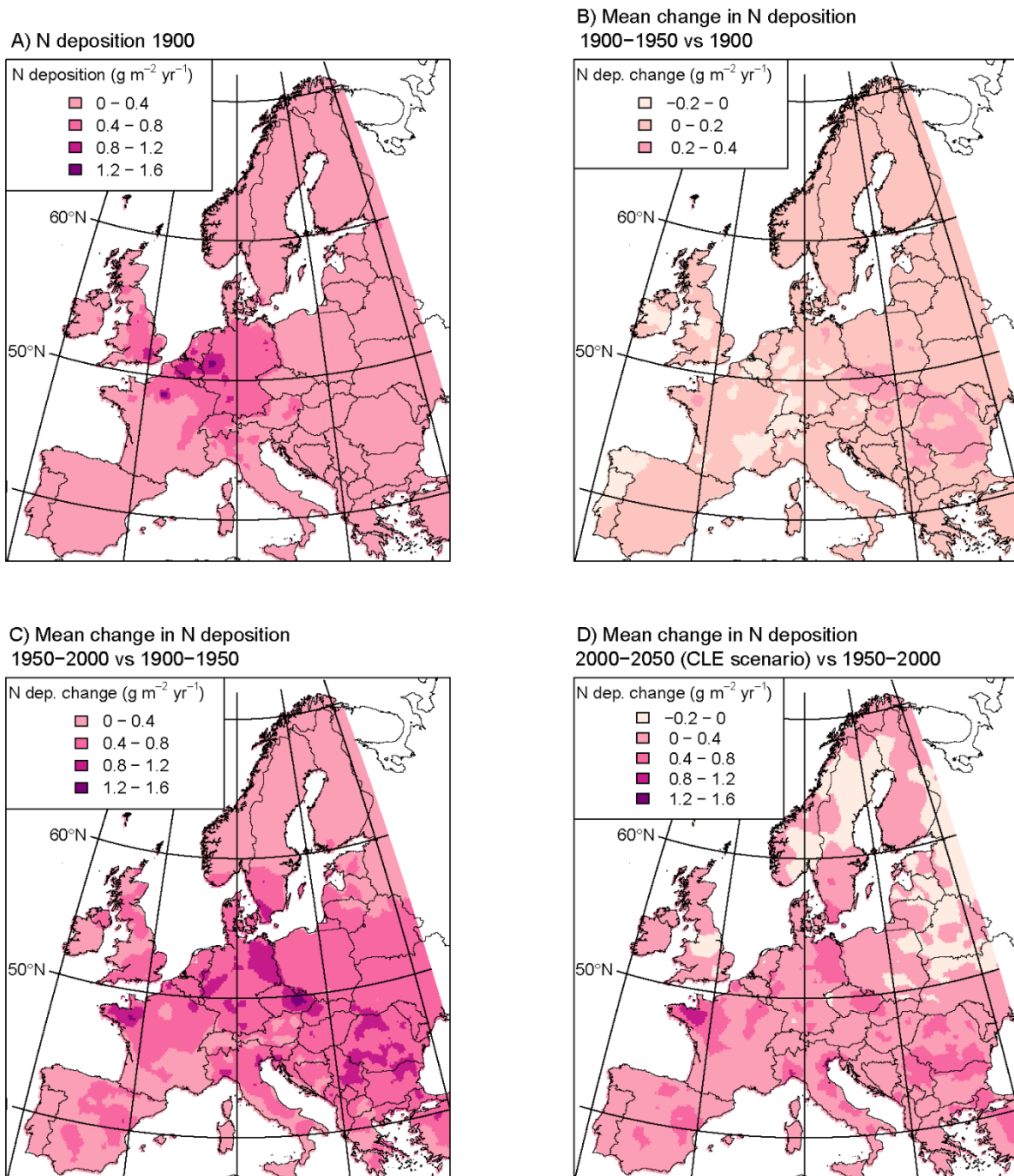
Trends of N deposition effects on Sphagnum

708 Table 2. Summary statistics of model predictions at the European scale (see Fig.1 for
 709 coverage). CLE N refers to the current legislation N emission scenario (GEA-CLE, here
 710 called CLE N) and A1 clim change to the A1 climate scenario. Constant N keeps N deposition
 711 at the mean of 2000-2010 and constant climate uses the mean of 1961-1990. Average effect of
 712 N deposition on *Sphagnum* production and *Sphagnum* C balance of *Sphagnum*-dominated
 713 peatlands are calculated as weighted means using peatland cover in grid cells as weights. Area
 714 indicates total peatland area where *Sphagnum* is estimated to have been, or to be, affected
 715 negatively or positively. The 1st and 99th percentiles are given to show the range of the grid
 716 cell predictions.

Period	Scenario	Average effect on production (1 st , 99 th %ile) (% change)	Average effect on C balance (1st, 99th %ile) (g m ⁻² yr ⁻¹)	Area (mil. ha)	
				neg	pos
1900 - 1950		-	-	-	-
1950 - 2000		-6 (-46, 9)	-6 (-46, 9)	0.84	0.36
2000 - 2050	CLE N, A1 Clim Change	-12 (-65, 16)	-12 (-65, 16)	1.11	0.44
	CLE N, Constant climate	-5 (-58, 25)	-5 (-58, 25)	0.92	0.63
	Constant N, A1 Clim Change	-15 (-52, 0)	-15 (-52, 1)	1.54	0.05
	Constant N and Constant climate	-8 (-49, 11)	-8 (-49, 11)	1.22	0.37

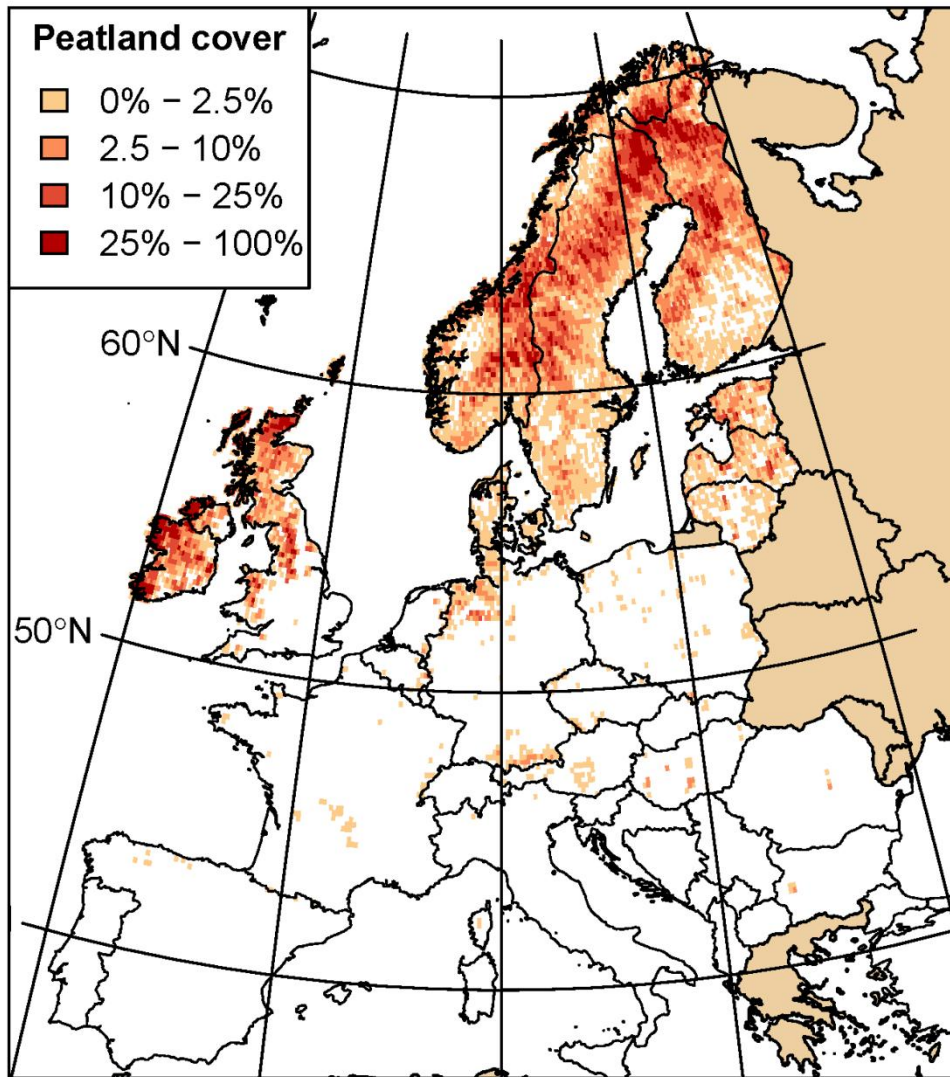
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Trends of N deposition effects on Sphagnum



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733 Fig 1. (A) Modelled nitrogen (N) deposition over Europe in the year 1900 and changes in N
734 deposition over three time periods; (B) mean 1900-1950 compared to 1900, (C) mean 1950-
735 2000 compared to the 1900-1950 mean, and (D), mean 2000-2050, compared to the 1950-
736 2000 mean. Predictions for 2000-2050 (C) are based on the GEA-CLE deposition scenario.

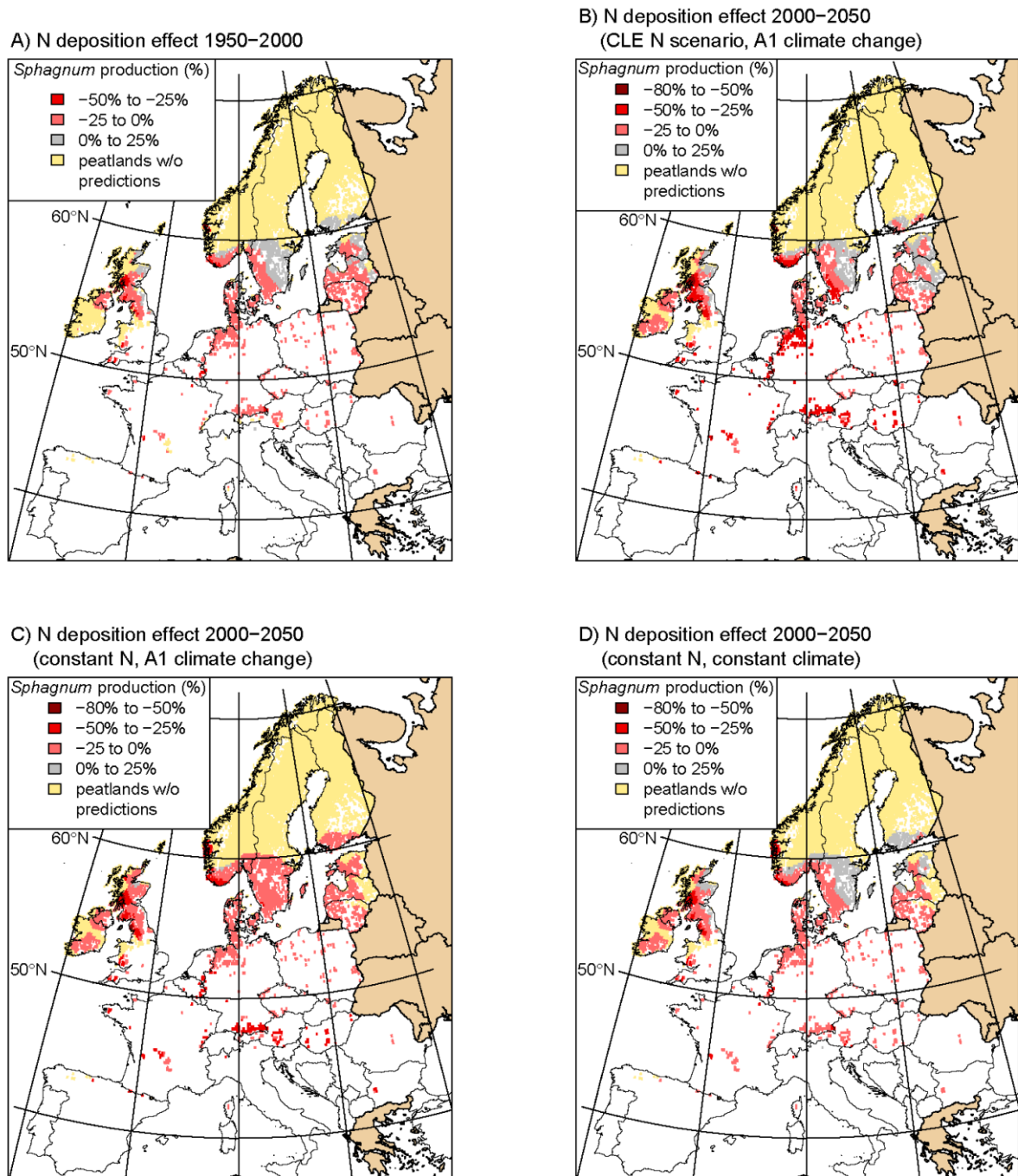


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738 Fig. 2. Map showing the abundance (proportion of total land area) and distribution of open to
739 semi-open *Sphagnum*-dominated peatlands in Europe. Data are based on the JRC forest map
740 and the CORINE 2006 land cover map. The light brown colour indicates countries without
741 data.

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Trends of N deposition effects on Sphagnum



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744 Fig 3. Effects of N deposition on *Sphagnum* production presented as percentage change
 745 relative to 1900. Annual mean effect between (a) 1950 to 2000, (b) 2000 to 2050 under a
 746 current legislation emission scenario (GEA-CLE, here called CLE N scenario) and the A1
 747 climate change scenario, (c) 2000 to 2050 assuming constant N deposition (mean of 2000-
 748 2010) and climate change scenario A1, (d) 2000 to 2050 assuming constant N deposition and
 749 constant climate (1961-1990 average). Only open to semi-open *Sphagnum* dominated
 750 peatlands were modelled. Yellow illustrates areas where N deposition were too low (< 0.4 g

Trends of N deposition effects on Sphagnum

751 $\text{m}^{-2} \text{yr}^{-1}$ increase since 1900) to make reliable predictions. Light brown colour indicates
752 countries without data on peatland distribution.

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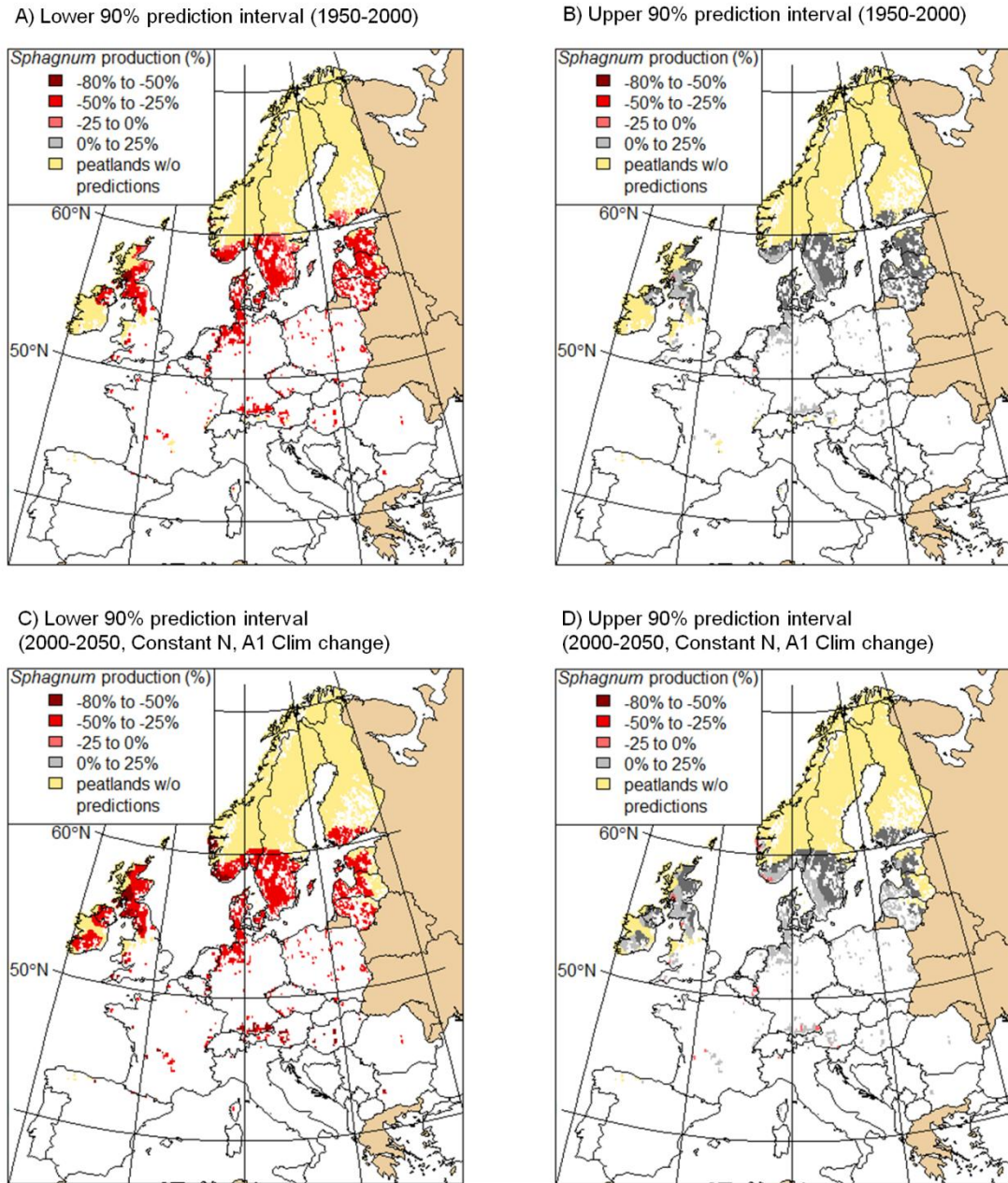
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778 **Supporting Information**

779 **Appendix S1.** Figures illustrating the uncertainties of some of the predictions

780 **Fig. S1.** Prediction uncertainties - 90% credible intervals



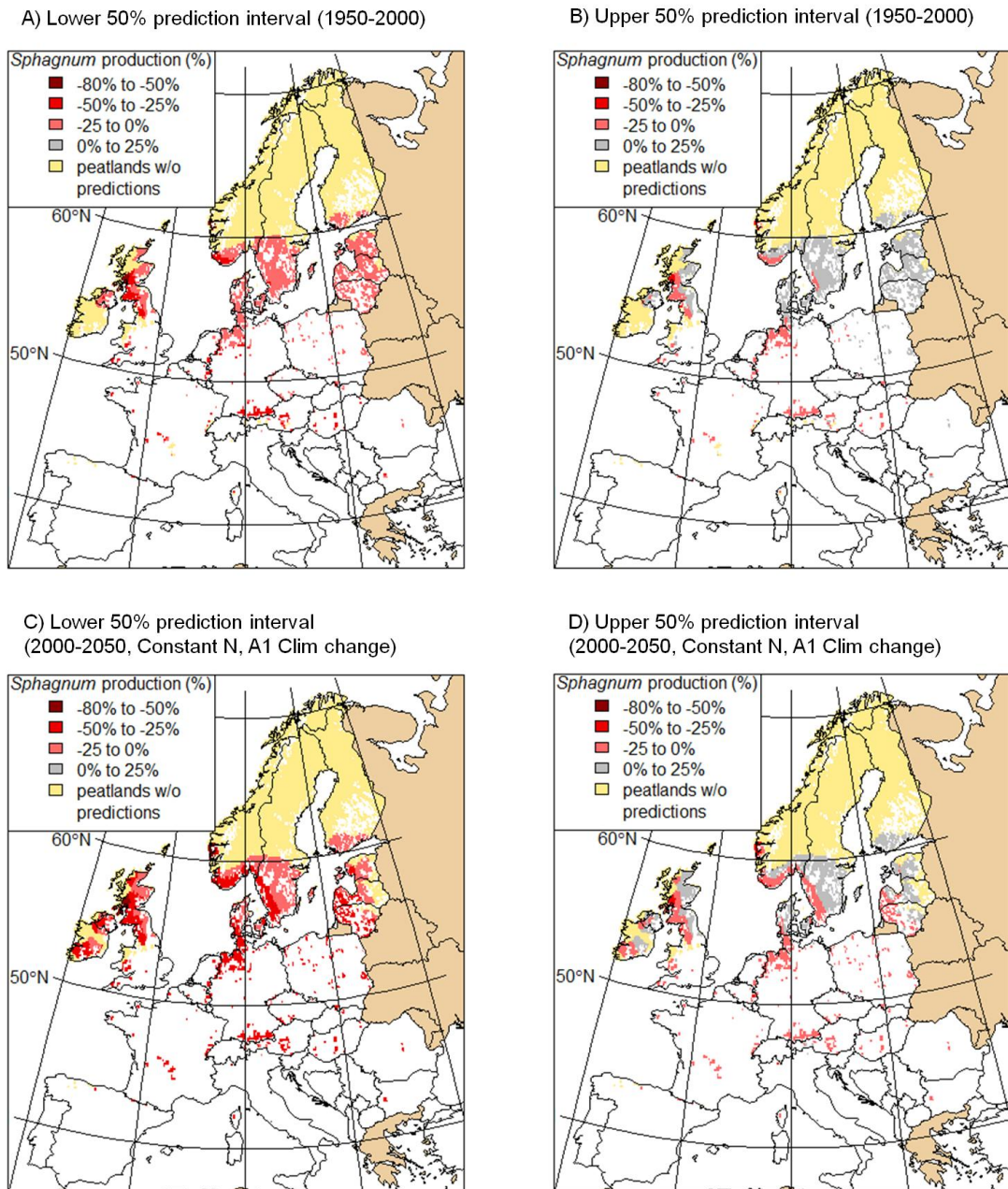
781

782 Fig. S1. Examples of prediction uncertainties (90% credible intervals) for two models, a-b)

783 1950-2000, and, c-d) 2000-2050 (Constant N, A1 Climate change scenario)

784

785 **Fig. S2.** Prediction uncertainties - 50% credible intervals



786

787 Fig. S2. Examples of prediction uncertainties (50% credible intervals) for two models, a-b)
 788 1950-2000, and, c-d) 2000-2050 (Constant N, A1 Climate change scenario)

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