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4 Phosphorus speciation of clay fractions from long-term fertility

5 experiments in Sweden

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19 Abstract

20 Phosphorus (P) losses from agricultural soils constitute a main driver for eutrophication of the Baltic Sea. There is limited knowledge about sorption and release processes of P in these soils, 21 especially concerning the effects of fertilization. In this study, P speciation of the clay fractions 22 23 from six different soils in long-term fertility experiments in Sweden was investigated by P Kedge XANES spectroscopy. As expected, unfertilized soils had lower concentrations of acid-24 digestible P compared with fertilized soils. Based on best-fit standards that emerged from linear 25 combination fitting (LCF) of XANES spectra, phosphate sorbed on iron (Fe) (hydr)oxides was a 26 dominant P species in clay fractions from unfertilized soils containing more than 35 mmol kg⁻¹ of 27 oxalate-extractable Fe. In contrast, P sorbed on aluminum (Al) (hydr)oxides predominated in 28 29 soils with lower concentrations of oxalate-extractable Fe. A greater proportion of organically bound P was fit for soil samples containing >2 % organic carbon. The soils included one 30 31 calcareous soil for which a greater proportion of P was fit as apatite. After long-term fertilization, 32 P had accumulated mainly as Al-bound forms (adsorbed species and precipitates) according to 33 the XANES analysis. Our research shows that P speciation in fertilized agricultural soils 34 depended on the level of P buildup and on the soil properties.

1. Introduction

37 Phosphorus (P) is a main driver of eutrophication in waters such as the Baltic Sea. The largest source of P from Sweden into the Baltic Sea is from agricultural fields (Boesch et al., 2006). 38 Phosphorus in soils is found in mineral phases, as adsorbed species (surface complexes) on 39 40 (hydr)oxide solids, and in organic forms. Adsorption / desorption processes are important in 41 controlling P solubility in soils, but precipitation and dissolution of minerals may also affect the solubility, especially in soils enriched in P. Organically bound P species in soils are also 42 important, in which case immobilization and mineralization affect the solubility of phosphate. 43 Aluminum (Al), and iron (Fe) (hydr)oxides are important for phosphate adsorption (Hingston et 44 al., 1967). Amorphous Al (hydr)oxides may be of greater importance for phosphate sorption on 45 46 P-rich soil particles (e.g. Lookman et al., 1996; Pierzynski et al., 1990). Various sequential extraction methods are commonly used to estimate P speciation of soils. Examples are the 47 procedures developed by Chang and Jackson (1957) and Hedley et al. (1982). However, a 48 49 common criticism of these procedures is that they are not specific in dissolving particular chemical forms of P, and therefore they are not sufficient for determining P speciation (e.g. 50 Barbanti et al., 1994). 51

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A more direct method used for speciation of organic P is solution ³¹P-NMR following alkaline extraction of a soil sample. However, certain forms of P can be hydrolyzed during the extraction, leading to artifacts. In addition, solid-state ³¹P-NMR analysis of soil has been used without pretreatment for speciation (e.g. Cade-Menun, 2005; Lookman et al., 1996). There are, however, limitations with this method, e.g., high P and low Fe concentrations are needed to obtain an adequate spectrum with minimal paramagnetic effects (Cade-Menun, 2005).

Hesterberg et al. (1999), Beauchemin et al. (2003) and Toor et al. (2005) introduced the use of 60 61 XANES spectroscopy to characterize P species in environmental samples. Dominant species of P are commonly estimated using linear combination fitting (LCF), where a weighted sum of 62 XANES spectra from selected P standards are fit to the spectrum from a sample (Kelly et al., 63 64 2008). Different chemical species of P have unique spectral features, for example: (1) P associated with transition metals (e.g. Fe) gives a weak, but clear pre-edge shoulder (Franke and 65 Hormes, 1995); (2) P in calcium (Ca) phosphates gives clear continuum resonances and post-66 edge shoulders (Franke and Hormes, 1995); (3) Al phosphates give a weaker pre-edge resonance 67 that overlaps with the strong white line (Khare et al., 2007), but also weak post-edge features that 68 commonly occur at higher energies than those of Ca phosphates (Franke and Hormes, 1995); and 69 (4) organic P species commonly do not show any clear pre- or post-edge features, which makes it 70 hard to differentiate between different organic P species (Doolette and Smernik, 2011; 71 72 Hesterberg, 2010).

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Accurate characterization of P speciation in fertilized soils is important to create better models 74 for predicting P mobilization and movement to surface waters. Although the P speciation of 75 fertilized soils has previously been estimated using XANES spectroscopy (e.g. Beauchemin et al., 76 77 2003; Lombi et al., 2006), these studies did not address long term changes of P speciation over time as a result of fertilization. In one of the few studies devoted to analyzing P speciation 78 changes resulting from fertilization (Ajiboye et al., 2008), samples of a Vertisol and a Mollisol 79 80 were incubated in the laboratory for short time periods, then characterized by P K-edge XANES spectroscopy. The results suggested an important role of adsorbed P species. 81

82

The Swedish soil fertility experiments represent a unique set of field plots for assessing the long-83 84 term effects of fertilization on soil-chemical properties (e.g. Börling et al., 2001; Carlgren and Mattsson, 2001; Gustafsson et al., 2012). At several sites, soil plots have been amended with 85 different levels of nitrogen (N), P, and potassium (K) fertilizers for periods of 40 to 60 years. 86 Recently, the P speciation of selected sites was studied by means of ³¹P NMR spectroscopy 87 (Ahlgren et al., 2013). According to this study, neither the absolute amounts nor the speciation of 88 organic P changed as a result of fertilization, suggesting that added P was accumulating as 89 inorganic P species. This result agrees with a similar study conducted in Finland (Soinne et al., 90 2011). 91 92

93 The aim of our investigation was to evaluate changes in P speciation due to long-term fertilization of some fine-textured agricultural soils, and we used P K-edge XANES 94 95 spectroscopy. That is, our objective was to determine how added P was bound. Soil samples from 96 the long-term Swedish soil fertility experiments were used, which have been extensively characterized in earlier studies (Börling et al., 2001; Djodjic et al., 2004; Svanbäck et al., 2013). 97 Moreover, detailed speciation of organic P using ³¹P-NMR has already been analyzed for these 98 soils (Ahlgren et al., 2013), and our XANES analysis should be more sensitive to differences in 99 inorganic P species. 100

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- 102

2. Materials and methods

103 *2.1. Soil sampling and preparation*

104 Soil samples were collected from six different sites included in the Swedish soil fertility

- 105 experiments (Carlgren and Mattsson, 2001): Fors, Kungsängen, Vreta Kloster, Bjertorp, Ekebo
- and Fjärdingslöv. A full description of the sites can be found in Kirchmann (1991), Kirchmann et

al. (1999) and Kirchmann et al. (2005). The experimental plots were established between 1957 107 108 and 1969 and include two different crop rotations, with and without livestock (Carlgren and Mattsson, 2001). In this investigation we used samples from the plots with crop rotation without 109 livestock. Nitrogen, phosphorus and potassium were applied as inorganic fertilizers. In total, 110 111 there are 8 combinations of fertilizer treatment for each crop rotation (Carlgren and Mattsson, 2001). We used soil samples from plots receiving 125 kg N ha⁻¹ yr⁻¹ for the Fors. Kungsängen, 112 Vreta Kloster and Bjertorp soils, and 150 kg N ha⁻¹ yr⁻¹ for the Ekebo and Fjärdingslöv soils. 113 Samples from plots receiving three different phosphorus and potassium fertilization treatments 114 were used; (1) control plots with no added P or K fertilizer, referred to as A3; (2) plots amended 115 with P and K that replaced the P and K removed with harvest, plus 15 and 40 kg ha⁻¹ yr⁻¹, 116 respectively on Ekebo and Fjärdingslöv soils, or replacement of harvest + 20 P and 50 K kg ha⁻¹ 117 vr^{-1} on other soils, referred to as C3; (3) and plots receiving replacement + 30 P and 80 K kg ha⁻¹ 118 yr⁻¹, referred to as D3. Because of time limitations in collecting P K-edge XANES spectra, we 119 120 used samples from only one replicate of each treatment.

121

Soil cores were collected from 0 to 20 cm around a 1-m diameter circle at a random location but
at a minimum distance of 50 cm from the edge of each plot. All samples were collected during
spring and autumn 2011, then immediately air-dried and sieved to <2 mm. One sample for each
fertilization treatment (no replicates within or between plots) was taken at each site. The samples
were analysed for bulk and clay mineralogy using X-ray diffraction (Hillier, 1999; Hillier, 2003;
Omotoso *et al.*, 2006). No substantial mineralogical differences were found between samples
taken across treatments at each site (data not shown).

Clay fractions of $<2 \mu m$ (equivalent spherical diameter) were separated from the soils by sedimentation according to Stokes' law. A suspension of approximately 20 g soil and 200 cm³ water was treated twice with ultrasonic dispersion for 5 minutes and stirred thoroughly in between. The suspension was left in a cylinder to settle under gravity for 16 h, and the top 20 cm suspension was siphoned off. The upper 20 cm was refilled with water and the sedimentation repeated once. The clay suspension was freeze-dried, and the clay fraction was stored dry until analysis.

137

138 2.2. Soil characterization

The particle size distribution was analyzed according to ISO 11277 (1998). The soil pH was 139 measured in a suspension of 10 g air-dried soil to 30 cm^3 of deionized H₂O. The organic carbon 140 content (OrgC) was measured by combustion using a LECO CNS-2000 analyzer (LECO, St. 141 Joseph, MI). The calcite content was measured by full pattern fitting of XRD data (Omotoso et 142 al., 2006) of a bulk soil sample after spray drying (Hillier, 1999). Soil test P was determined 143 144 according to the Swedish standard ammonium lactate (AL) method of Egner et al. (1960), for which 5 g dry soil was equilibrated for 1.5 h with 100 cm³ solution containing 0.1 M AL and 0.4 145 M acetic acid. Hydrochloric acid-digestible P was measured by boiling 2 g dry soil in 50 cm³ of 2 146 M HCl for 2 h. Oxalate extractions were made according to van Reeuwijk (1995) using 1 g dry 147 soil to 100 cm³ solution and an equilibrium time of 4 h in darkness. The extract was then filtered 148 through a 0.2 µm single-use filter and diluted 1:5 in H₂O. Phosphate in the oxalate extract (PO₄-149 150 ox) was analysed colorimetrically according to Wolf and Baker (1990) using a Tecator Aquatec 5400 spectrophotometer with flow injection analysis. Oxalate-extractable aluminum (Al-ox) and 151 152 iron (Fe-ox) were determined by ICP-OES using a PerkinElmer 5300 DV instrument.

Pyrophosphate-extractable iron (Fe-pyro) and aluminum (Al-pyro) were determined in an extract 153 of 1 g dry soil in 100 cm³ of 0.1 M Na₂P₂O₇· 10H₂O equilibrated for 16 h. The extract was filtered 154 through a 0.2 µm single-use filter and diluted 1:5 in H₂O. Pseudo-total P (PsTotP) was measured 155 by acid digestion in aqua regia (modified from ISO 11466, 1995), for which 3 g dry soil was 156 equilibrated for 16 h with 30 cm³ agua regia solution, then the mixture was boiled for 2 h with a 157 water condenser to minimize evaporation. The condenser was flushed with 20 cm^3 of 0.5 M 158 HNO₃ and the rinsate was collected into the digestate, which was filtered and diluted to 100 cm³ 159 with 0.5 M HNO₃. The final solution was analyzed for P using ICP-OES. Digestion in aqua regia 160 161 is commonly used for measuring pseudo-total concentrations of elements, e.g. phosphorus, in soils (e.g. Stroia et al., 2013). Previous research shows that between 55 and 102 % of the total P 162 163 in soils is recovered by aqua regia digestion (Hornburg and Luer, 1999; Ivanov et al., 2012). 164 These investigations found strong correlations between PsTotP and total soil P determined by 165 HClO₄ extraction.

166

167 2.3. Phosphorus K-edge XANES of clay fractions

The quality of synchrotron P K-edge XANES data from soil samples depends on the total soil P 168 concentration, the intensity of incident synchrotron X-rays, the sensitivity of the fluorescence 169 detector used, and concentrations of other elements such as Si that contribute to the total 170 171 fluorescence signal to the detector. To improve the quality and reliability of our data, we collected XANES data on clay fractions separated from our fine-textured soil samples taken from 172 173 the long-term fertility experiments. Williams and Saunders (1956) found that the majority of soil 174 P in fine-textured soils is associated with the clay fraction. Our clay-fraction samples were packed into wells of acrylic holders (sample volume of 15 x 6 x 1.5 mm; w x h x d), and the 175 176 surface was smoothed with a spatula and covered with 5 µm-thick polypropylene X-ray film

177	(Spex Industries, Columbia, IL). The P K-edge XANES data were collected at Beamline X-15B
178	of the National Synchrotron Light Source at Brookhaven National Laboratory.

An Si(111) monochromator crystal was used, and the flux was approximately $1 \cdot 10^{12}$ photons s⁻¹. 180 181 The beam was focused to approximately 1.1 mm diameter with focusing mirrors. Fluorescence signals were collected in a He atmosphere using a single-element solid-state Ge detector. The 182 energy was calibrated to 2151 ± 2 eV at the 1st-derivative maximum of a hydroxyapatite standard. 183 Scans were recorded within the energy range of 2100 to 2470 eV. The step size was varied as 184 follows: 2 eV between 2100 and 2130 eV, 0.2 eV between 2130 and 2156 eV, 0.5 eV between 185 2156 and 2153 eV, 1 eV between 2153 and 2180 eV, 2 eV between 2180 and 2250 eV, and 5 eV 186 between 2250 and 2470 eV. Dwell times ranged from 2 to 6 seconds, with longer times used 187 across the edge region. To decrease the fluorescence signal from Si in the samples, the detector 188 window was covered with a ~40-µm thick polypropylene film. A minimum of 13 scans for the 189 190 unfertilized samples and 9 scans for the fertilized samples were collected for each sample. A hydroxyapatite sample was run between each sample to ensure stability of the energy calibration. 191 192

193 2.4. Phosphorus K-edge XANES data analysis

All data analyses were performed using the Athena software in the Demeter suite of programs (Athena v0.8.056; Ravel and Newville, 2005 and Demeter v0.9.18). All sample and standard spectra were energy-calibrated to a common energy scale, where the derivative max of hydroxyapatite was set at 2151 eV and the shift applied to sample data collected after each calibration. The calibrated scans for each sample were aligned and merged. The spectra were baseline-corrected by subtracting a linear regression through the pre-edge region (approx. -20 to -5 eV relative to E_0 set at the 1st derivative maximum for a given sample or standard) and

background-corrected using a quadratic function through the post-edge region (+30 to approx.

+58 eV; except for the Fors sample, for which the background correction was extended to +100

eV). The pre-edge range for baseline subtraction was approximately parallel to the normalization

range. The same data treatment procedure was used for all standards.

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- 206

2.5. Linear combination fitting analysis

Linear combination fitting (LCF) was performed across an energy range from -10 eV to +30 eV 207 relative to E₀ to investigate the P speciation. In total, 86 different standard spectra were used for 208 the unfertilized samples. This group of standards contained (1) mineral samples from Ingall et al. 209 210 (2011); (2) mineral standards and standards with adsorbed phosphate and organically bound P 211 from Hesterberg et al. (1999); and (3) mineral standards and standards with adsorbed phosphate 212 from Eveborn et al. (2009). Principal component analysis (PCA) was performed in the Demeter 213 software to limit the number of independent standards needed to fit the spectra (Beauchemin et 214 al., 2002). To determine which standards most likely accounted for the variation in the sample spectra and therefore to include in the fit, target transformation was performed using the same 215 216 software. To quantitatively estimate P speciation, LCF analysis was performed. However, because of the relatively low signal-to-noise ratio in the data and the insensitivity of P K-edge 217 218 XANES analysis to different species, a maximum of three standards were used in each fit (Beauchemin et al., 2003). To limit the number of standards in the LCF analysis, we excluded 219 220 standards with the lowest probability to account for the variation based on target transformation. 221

Two approaches to fitting were used: (1) sample spectra from unfertilized plots were fit with all standards selected from the target transform; and (2) the best-fit model spectrum generated for the unfertilized samples was included as one standard when fitting the sample spectra from

225	fertilized plots. The latter approach assumes that the speciation of the initial soil P did not change
226	due to fertilization. Fits having a sum of the weight factors in the range of 80 to 120 % of total P
227	were considered acceptable. The sums of fitted weighting factors were adjusted to a total sum of
228	100%. Additional standards needed to fit the fertilized sample, which were those used in the data
229	set of Hesterberg et al. (1999), but re-collected on NSLS Beamline X15B, were assumed to
230	represent the P species formed as a result of fertilization. Uncertainties in XANES fitting
231	parameters were calculated by the Athena software (v0.8.056; Ravel and Newville, 2005). Other
232	statistical analyses were done in MINITAB 16® 16.2.0.
233	
234	3. Results
235	3.1.Extractable P in comparison to P speciation
236	The concentrations of ammonium lactate- and oxalate extractable P increased after fertilization
237	for all soils (Table 1). This trend suggests an increase in inorganic P in adsorbed or mineral
238	species, which are expected to be dissolved by acid extractants (e.g. Hartikainen et al., 2010).
239	
240	The greatest relative increase of extractable P following fertilization was observed in the soil
241	from Ekebo and Fjärdingslöv. The smallest change was observed in the soil from Fors, which
242	also had the highest content of PsTotP (Table 1). By contrast, the soil from Fjärdingslöv had the
243	lowest concentration of PsTotP, but also a low concentration of oxalate- and pyrophosphate
244	extractable Al and Fe. The greatest change in PsTotP between the unfertilized sample and the one
245	receiving the highest level of P fertilizer was observed for the samples from Bjertorp, whereas the
246	smallest change was observed for the Fors soil.
247	

3.2. Phosphorus speciation in unfertilized samples

Linear combination fitting analysis on clay fractions from the samples with no added P showed 249 250 that P speciation differed between soil samples from different sites (Figure 1 and Table 2). Calcium phosphate (apatite) was observed in the best fit for all samples, except for the Ekebo and 251 Fjärdingslöv samples. The Ekebo sample had the lowest pH of all samples (5.9-6.1). However, 252 253 despite the low pH. Ca phosphates were fitted to the data in two of the five best fits; hence the 254 presence of trace amounts of Ca phosphates could not be excluded. In the case of the Fjärdingslöv sample, Ca phosphates occurred in all but the first of the five best fits. It should be 255 noted, however, that there might be signal overlap between Ca phosphates and organic P when 256 only small amounts occur in the sample (Doolette and Smernik, 2011). There were statistically 257 significant relationships between PsTotP in the soils from the unfertilized treatments and P bound 258 as Ca phosphates (r=0.89*). A relationship between the ratio between oxalate extractable Fe and 259 Al, and the amount of P adsorbed to Al surfaces was also found (r=0.92**). This suggests that for 260 261 unfertilized soils containing more oxalate-extractable Al than Fe, P is bound predominantly to Al 262 (hydr)oxides (see also Figure 2).

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For the Fors soil, the best fits included 65 % of the P present as carbonate hydroxyapatite fluorian 264 (a Ca phosphate standard from Ingall et al., 2011; Figure 1 and Table 2). Phosphate adsorbed to 265 Al and/or Fe (hydr)oxides was also included in the best fits for all samples. In soils with >35 266 mmol kg⁻¹ of oxalate extractable Fe (Table 1), XANES fitting indicated that a major part of P was 267 bound to Fe (hydr)oxides. The Kungsängen and Ekebo soils, which had concentrations of 268 oxalate-extractable Fe in the upper range (160 and 40 mmol kg⁻¹, respectively), also had the 269 highest concentrations of organic C (2.2 and 2.3 %, respectively). For these soils organic P (as 270 evidenced by P bound to phytate or lecithin) occurred in the best fit for the unfertilized samples. 271 272 For the Ekebo sample, the best fit also included phosphate adsorbed to Al (hydr)oxides,

273 consistent with its relatively high concentration of oxalate-extractable Al compared to Fe (Figure274 2).

275

For the clay fractions from the unfertilized Bjertorp and Fjärdingslöv samples the majority of the
P was bound to Fe (hydr)oxides. The molar ratio of oxalate-extractable P to Fe in these soils was
rather low (<0.08) indicating additional capacity for adsorption of phosphate to Fe (hydr)oxides.
The soils from Kungsängen and Ekebo had a P to Fe ratio higher than 0.08 in the oxalate extract;
however in these soils less P was bound to Fe (hydr)oxides according to the XANES analysis
(Figure 2).

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For the soils from Fors and Vreta Kloster, the XANES results showed no significant contribution of Fe-bound P, which is consistent with the lower concentration of oxalate extractable Fe in these soils (<35 mmol kg⁻¹; Figure 2). Instead the XANES results of the clay fractions showed that the speciation of P was dominated by Ca phosphate (apatite) and by phosphate adsorbed to Al (hydr)oxides.

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3.3. Phosphorus speciation in fertilized samples

For most samples the standard spectra giving the best fit for the phosphate added by fertilization was phosphate bound to Al (hydr)oxide (boehmite) or to Al-treated peat (Figure 1 and Table 3). The Vreta Kloster was an expection; here the P added by fertilization seems to have been bound as apatite. Although 20 % vivianite was fit to the XANES spectra for the Kungsängen sample (Figure 1 and Table 3), it seems more likely that other Fe-P mineral phases or sorbed P with similar spectral features may be present as this is a well-drained (well aerated) soil. In the best fit to the Kungsängen sample, the standard from phosphate adsorbed on Al-treated peat was also

included. However, the pyrophosphate extraction indicated that 7 mmol Al kg⁻¹ was bound to 297 298 organic matter, which is comparably low in this soil set. Phosphate adsorbed to Al-treated peat was also observed for the sample from Ekebo, which had the highest concentration of organic C 299 (2.3 %) and the highest concentration of pyrophosphate-extractable Al of all soils studied (70 300 mmol Al kg⁻¹). By contrast, the sample from Fjärdingslöv had a comparably low concentration of 301 extractable Al (9 mmol kg⁻¹) and Fe (12 mmol kg⁻¹), and also a low concentration of organic 302 carbon (1.37 %). In this sample, XANES fitting indicated that approximately 30 % of the added P 303 had instead been precipitated as Ca phosphates. 304

305

306 **4. Discussion**

307 *4.1.Phosphorus in clay fraction compared to soils*

In this study, the clay fraction was used for the XANES analyses instead of the bulk soil sample. 308 309 It has earlier been shown that a majority of the phosphorus is commonly found in the clay 310 fraction (Williams and Saunders, 1956). The P speciation in the clay fraction may, however, not be identical to the one in bulk soils for the following reasons: (1) there is a risk of overestimation 311 312 of organic phosphorus caused by an enrichment of organic P in the clay fraction (Williams and Saunders, 1956), and (2) there might be a risk of underestimation of crystalline calcium 313 phosphates. Liu et al. (2014) showed that a majority of the P in the colloidal fraction ($< 1 \mu m$) 314 from water extractions was associated with aluminum and iron. This indicated a smaller 315 quantitative importance of calcium phosphates in these finer soil fractions. However, no organic 316 317 P species was included in this investigation. Hence, because of this, the importance of iron and aluminum oxides may possibly be overestimated (Williams and Saunders, 1956). Clearly, further 318 studies are required to investigate to what extent clay fractions, as used in this study, may be 319 320 representative for the P in bulk soils.

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seems to be smaller than that of, e.g. sulfur (Prietzel et al., 2007). In addition, no chemical pretreatment was used in our method as described by Hillier (2003). 324 325 4.2. Speciation of P in unfertilized soils 326 For this diverse group of Swedish agricultural soils, the XANES fitting results show that P was 327 mainly bound as Ca phosphates (predominantly apatite) and adsorbed to Al and Fe (hydr)oxides 328 in the clay fraction. This result is consistent with other XANES studies that commonly reported 329 Ca phosphates in soil (e.g. Ajiboye et al., 2008; Beauchemin et al., 2003; Lombi et al., 2006). 330 331 Further, our study confirms that a large part of the soil P is adsorbed to Al and Fe (hydr)oxides. 332 333 The results also agree with those of Khare et al. (2004), who showed that phosphate is commonly

The isolation of the clay fraction may also by itself affect the P speciation, but for P this effect

334 bound to both Al and Fe (hydr)oxides when they occur in a mixture.

335

336 The identity of the Fe and Al sorbent phases is not clear, however. Concerning oxalate extractable Fe, it is likely that ferrihydrite is a main contributor, although other mineral forms 337 such as biotite or magnetite may also be dissolved in this extraction (Parfitt and Childs, 1988). 338 339 Concerning the Al phases the question remains even more open: allophane, amorphous Al

(hydr)oxide or hydroxy-interlayered Al phases all seem to be possible candidates. 340

341

4.3. Speciation of P in fertilized soils 342

The XANES fitting results suggest that P added to the soils through fertilization was adsorbed 343

344 mainly to the Al phases. This is consistent with earlier studies showing a strong correlation

345	between extractable Al and P in Swedish agricultural soils (Börling et al., 2004; Ulén, 2006). In
346	addition, our results confirm earlier NMR spectroscopic results showing that P added with
347	fertilization is bound as inorganic forms in these soils (Ahlgren et al., 2013). Khatiwada et al.
348	(2012) showed that directly after fertilization of a silt loam, P was mainly found as Ca phosphates
349	(which was the fertilizer) but after six months it was recovered primarily as an adsorbed phase.
350	Other studies have shown that added P was precipitated as Ca phosphates in soils with high pH
351	(>7.6) and calcium carbonate (>3.4 %) content, but with low concentrations of oxalate
352	extractable Al and Fe (< 27 mmol and < 56 mmol kg ⁻¹ , respectively; Ajiboye et al., 2008; Lombi
353	et al., 2006; Kar et al., 2012). However, in the sample from Fors, which should be comparable to
354	these samples, the added P was primarily adsorbed to Al (hydr)oxide surfaces. The Fors soil had
355	a somewhat higher concentration of oxalate extractable Al, indicating the existence of Al
356	(hydr)oxide phases in this soil.

However, for the Kungsängen soil, the XANES fitting analysis suggested the formation of Fe
phosphate precipitates. Precipitation of phosphates into Al and Fe phosphates has also earlier
been observed after P amendment to soils (Ajiboye et al., 2008). The differences between Fe
phosphate and P bound to iron (hydr)oxides are, however, rather subtle when these phases occur
in a mixture.

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It is clear both from this study and from earlier modeling studies (Cui and Weng, 2013;

Gustafsson et al., 2012) that the oxalate extractable Al in agricultural soils may represent a highly

reactive P-sorbing phase, probably because of a high surface area and a high affinity for P. Thus

367 to derive better mechanistically based geochemical models that can predict P sorption/desorption

and leaching patterns, more emphasis should be placed on understanding the interactions betweenAl and P in agricultural soils.

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371 *4.4.Comparison to earlier studies in the Swedish fertility experiments*

372 Svanbäck et al. (2013) investigated P losses from columns of topsoil using samples from four of the sites of our investigation, Vreta Kloster, Bjertorp, Ekebo and Fjärdingslöv. They showed that 373 phosphate losses from the unfertilized sites were of similar magnitude, except for the soil from 374 Fjärdingslöv, where the phosphate loss was higher. This is consistent with the results of the 375 376 present study, which showed that the latter soil was the one with the lowest concentration of 377 extractable Al and Fe, indicating low P adsorption capacity (Ulén, 2006). For the fertilized 378 samples, the losses observed by Svanbäck et al. (2013) were of similar magnitude as could be expected for the soil from Ekebo, where the P loss was lower. The present investigation showed 379 380 that the added P was adsorbed mainly to Al (hydr)oxides in this soil. Moreover Ekebo was the soil with the highest concentration of oxalate-extractable Al, and therefore it probably had the 381 greatest potential to adsorb the added phosphate. Svanbäck et al. (2013) also added manure and 382 383 measured the P losses one week after application. No increase in P losses was observed from the soil from Ekebo, which may be explained by high concentrations of non-crystalline Al-hydroxide 384 as indicated by oxalate extraction. Phosphorus losses by leaching were minor for the unfertilized 385 386 soil from Vreta Kloster, and somewhat greater for the soils from Bjertorp and Fjärdinglöv.

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5. Conclusions

Phosphorus K-edge XANES analysis indicated that in clay fractions of unfertilized soils from
long-term Swedish soil fertility experiments, the speciation of P was dominated by P adsorbed
onto Al and Fe (hydr)oxide phases, and by Ca phosphate (apatite). In soils with > 2 % organic C,

organic P was also indicated. In soils containing > 35 mmol kg⁻¹ oxalate extractable Fe, P was
mainly adsorbed to Fe (hydr)oxides, whereas P was adsorbed mostly to Al (hydr)oxides in soils
after fertilization. After long-term fertilization, acid-digestible P increased. Results from XANES
spectroscopy showed that P adsorbed to Al (hydr)oxide phases usually increased more than Fe
(hydr)oxide-adsorbed P, Ca-phosphate, or organic P fractions.

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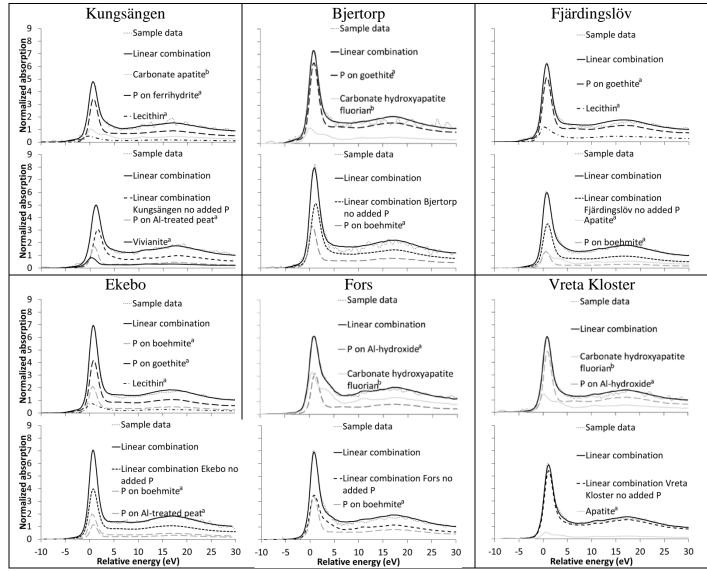
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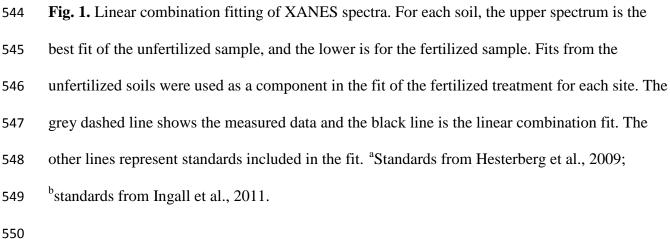
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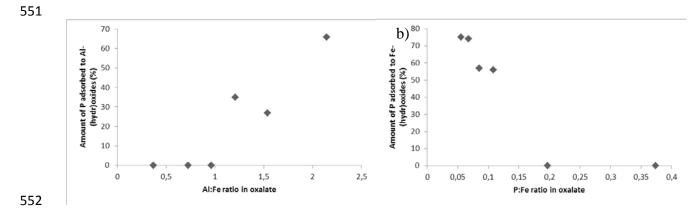


Fig. 2. Left: the relationship between the ratio of oxalate-extractable Al and Fe to the calculated
amount of phosphate adsorbed to Al (hydr)oxides for unfertilized samples. Right: the relationship
between the ratio of oxalate-extractable P and Fe to the calculated amount of phosphate adsorbed
to Fe (hydr)oxides for unfertilized samples.

Table 1

Site		Texture	Clay	pН	OrgC	Calcite	Al-ox	Al-py	Fe-ox	Fe-py	P-AL	P-ox	P-HCl	PsTotP
			(%)		(%	6)				(mmol i	kg ⁻¹)			
Fors	A3	silt loam	17	7.71	1.50	5.59	37.4	5.9	31.0	2.3	3.6	11.6	23.6	34.5
	C3	silt loam	15	7.72	1.47	7.39	33.0	5.6	31.0	3.0	4.8	13.2	22.9	36.5
	D3	silt loam	16	7.65	1.47	6.53	32.6	5.6	31.0	3.6	5.5	14.5	27.1	37.8
Kungsängen	A3	Clay	56	6.47	2.18	0.12	58.9	4.8	159.7	21.3	1.0	13.6	17.1	26.8
	C3	Clay	56	6.49	2.04	0.06	72.6	6.7	189.6	25.4	2.3	16.5	21.3	32.3
	D3	Clay	56	6.47	2.12	0.10	70.0	7.0	204.1	27.2	2.9	17.4	23.2	34.2
Vreta Kloster	A3	silty clay loam	39	6.52	2.03	0.12	63.0	13.3	29.4	6.4	1.0	5.8	14.2	20.7
	C3	silty clay	44	6.56	2.03	0.08	67.5	13.0	34.0	6.3	4.5	13.2	20.3	27.8
	D3	silty clay	45	6.56	2.04	0.13	61.2	13.3	32.2	6.4	5.2	14.5	22.9	30.7
Bjertorp	A3	silty clay loam ¹⁾	30	6.29	2.05	0.09	54.9	11.5	75.7	5.6	0.6	4.2	15.8	18.7
	C3	silty clay loam ¹⁾	31	6.39	1.98	0.11	51.9	12.6	66.3	6.1	2.6	11.6	23.6	25.8
	D3	silty clay loam1)	32	6.45	2.02	0.07	56.3	13.0	75.0	6.3	4.2	14.9	27.8	30.7
Ekebo	A3	Loam ¹⁾	16	5.91	2.32	0.05	63.7	36.0	41.5	25.6	1.0	4.5	19.0	19.4
	C3	Loam ¹⁾	14	6.07	2.31	0.05	81.9	52.6	44.2	26.5	3.2	11.3	22.9	23.9
	D3	Loam ¹⁾	12	6.03	2.55	0.14	94.1	70.0	39.9	24.5	6.5	16.1	30.7	30.7
Fjärdingslöv	A3	sandy loam ¹⁾	18	6.49	1.23	0.17	32.6	7.4	34.0	7.3	0.6	2.3	11.0	11.9
	C3	sandy loam ¹⁾	16	6.58	1.36	0.13	33.4	8.9	28.3	9.0	3.2	5.8	15.8	16.5
	D3	sandy loam1)	16	6.59	1.37	0.19	34.1	9.3	31.9	12.4	6.1	11.6	20.3	22.6

Selected properties of the soil samples studied.

Clay = clay content from pipette method, Al-ox = oxalate extractable Al, Al-py = pyrophosphate extractable Al, Fe-

ox = oxalate extractable Fe, Fe-py = pyrophosphate extractable Fe, P-AL= ammonium lactate extractable P, P-HC l= P after hot extraction with 2 M HCl, P-ox = Oxalate extractable PO₄, PsTotP = aqua regia-digestible P.

Table 2 566

Phosphorus speciation in clay fractions from unfertilized soils as evidenced from linear 567

combination fitting of XANES spectra. The five best fits are numbered from 1 to 5 in italics. 568

		CaP	FeP	AlP	P on Fe (hydr)oxides	P on Al (hydr)oxides	OrgP	R-factor ^a
Fors	weight (%)	65±1%				35±1%		0.003
	presence ^b	1,2,3,4,5				1,2,3,4,5		
Kungsängen	weight (%)	31±7%			57±4%		12±1%	0.004
0 0	presenceb	1,2,3,4,5			1,2,3,4,5		1,2,3,4,5	
Vreta K	weight (%)	34±2%				66±2%		0.005
	presenceb	1,2,3,4,5				1,2,3,4,5	4,5	
Bjertorp	weight (%)	25±1%			75±1%			0.008
5 1	presenceb	1,2,3,5			1,2,3,4,5		4	
Ekebo	weight (%)				56±11%	27±9%	17±1%	0.003
	presenceb	3,5			1,2,3,4,5	1,2,3	1,2,4	
Fjärdingslöv	weight (%)	*			74±1%		26±1%	0.005
5 0	presence	2,3,4,5	5		1,2,3,4,5	3	1	

569 570 CaP = calcium phosphates, FeP = crystalline iron phosphates, AlP = crystalline aluminum phosphates, P on Fe-

oxides = Phosphate adsorbed on iron (hydr)oxides, P on Al-ox = phosphate adsorbed on aluminum (hydr)oxides,

571 orgP = organic phosphorus. ^a R-factor calculated according to Ravel (2009).

573 **Table 3**

574 Phosphorus speciation in clay fractions from fertilized soils as evidenced from linear combination

		A3	CaP	FeP	AlP	P on Fe (hydr)oxid es	P on Al (hydr)oxides	Org P	R-factor
Fors	weight (%)	57±1%				03	43±1%		0.004
	presence	1,2,3,4,5				4	1,2,3,4,5		
Kungsängen	weight (%)	59±6%		20±1%			21±5%		0.003
	presence	1,2,3,4,5		1,3,4	4,5		1,2	2,3	
Vreta K	weight (%)	88±1%	12±1%						0.003
	presence	1,2,3,4,5				3	1,2,4,5		
Bjertorp	weight (%)	62±1%					38±1%		0.006
	presence	1,2,3,4,5				2,4	1,2,3,5		
Ekebo	weight (%)	56±5%					44±6%		0.001
	presence	1,2,3,4,5			3,5	4	1,2,3,4,5		
Fjärdingslöv	weight (%)	54±1%	29±2%				17±2%		0.002
	presence	1,2,3,4,5	1,2,3,4,5			3,4	1,2,5		

575 fitting of XANES spectra. The five best fits are numbered from 1 to 5 in italics.

576 A3 = fit for the unfertilized treatment from the same site, CaP = calcium phosphates, FeP = crystalline iron

577 phosphates, AIP = crystalline aluminum phosphates, P on Fe-oxides = phosphate adsorbed on iron (hydr)oxides, P on578 Al-ox = Phosphate adsorbed on aluminum (hydr)oxides, orgP = organic P. ^a R-factor calculated according to (Ravel,

579 2009).

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