

Emerging multiscale insights on microbial carbon use efficiency in the land carbon cycle

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Microbial carbon use efficiency (CUE) affects the fate and storage of carbon in terrestrial ecosystems, but its global importance remains uncertain. Accurately modeling and predicting CUE on a global scale is challenging due to inconsistencies in measurement techniques and the complex interactions of climatic, edaphic, and biological factors across scales. The link between microbial CUE and soil organic carbon relies on the stabilization of microbial necromass within soil aggregates or its association with minerals, necessitating an integration of microbial and stabilization processes in modeling approaches. In this perspective, we propose a comprehensive framework that integrates diverse data sources, ranging from genomic information to traditional soil carbon assessments, to refine carbon cycle models by incorporating variations in CUE, thereby enhancing our understanding of the microbial contribution to carbon cycling.

Earth System Models (ESMs) are indispensable tools for predicting the planetary response to climate change¹. The accuracy and reliability of ESMs are crucial for informing climate projections that guide policy decisions. Soils store more carbon (C) than plants, the surface ocean or

the atmosphere, and thus are critical for the functioning of the Earth system². While ESMs are becoming increasingly complex, their predictions of soil organic C (SOC) stocks have improved only marginally in recent decades^{3,4}.

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Microbial communities process most of the C entering the soil, thereby shaping its fate^{3,6}. Microbes metabolize multiple C sources, including detritus, root exudates, and microbial metabolites⁷. The energy needed to acquire C depends on whether the compounds can be taken up directly or require prior enzymatic degradation⁸. Additionally, microbial community composition and functioning are influenced by prevailing climatic conditions^{9–11}. The general omission of microbial community structure and related processes in C cycle models has been suggested as one of the causes for their poor performance in predicting SOC stocks and their responses to climate change^{12,13}.

Recognizing the impracticality of representing every conceivable microbial metabolic pathway, many models combine a spectrum of microbial processes into a single metric referred to as microbial C use efficiency (CUE)^{14,15}. CUE, as a model parameter or as a system property emerging from multiple co-occurring processes, represents the fraction of C uptake allocated to the production of new microbial

biomass¹⁶. Using this definition, CUE declines as more C is used for respiration to generate energy (for substrate uptake, cellular maintenance, enzyme production) or for exudation (extracellular enzymes, polysaccharides)^{17,18}. This pragmatic approach streamlines the modeling of soil C cycling by incorporating the diverse fates of microbial C, including biomass production, respiration, and exudation, thereby providing a more comprehensive understanding of microbially-mediated C-pathways.

However, accurately integrating the spatial or temporal dynamics of microbial CUE into soil C models remains a significant challenge. Most of the current C cycle models either lack explicit representation of CUE or treat it as a constant value⁴, despite our understanding that CUE varies under different environmental conditions. For example, observations indicate significant variability in CUE at the global scale⁸, which may be partially attributed to inconsistencies among measurement techniques (Fig. 1a). Moreover, comparisons across ecosystems reveal that CUE is generally higher in grasslands than in croplands, with

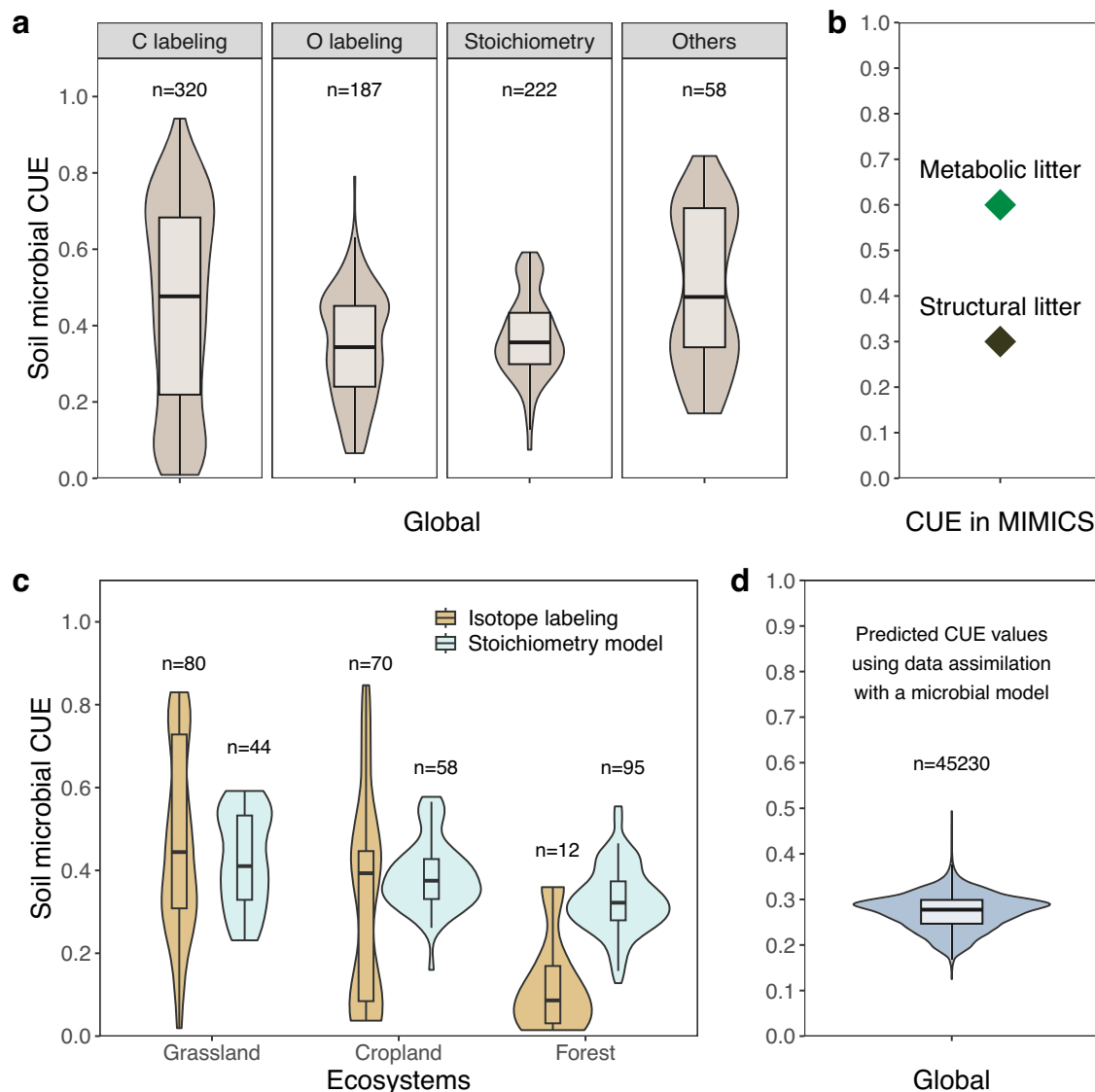


Fig. 1 | Variability of carbon use efficiency (CUE) at a global scale. **a** Observation-based CUE estimates at the global scale from C (¹³C and ¹⁴C) and ¹⁸O isotopic labeling, stoichiometric modeling and other methods. Data were collected from^{19,21,49,95,114}. **b** CUE constants used in the Microbial-Mineral Carbon Stabilization model (MIMICS) for two litter types (diamonds). Metabolic litter comprises plant litter that decomposes easily, whereas structural litter is more resistant to

decomposition¹³¹. **c** Observation-based estimates for different ecosystems using isotopic labeling¹¹⁴ or stoichiometric modeling¹⁹. **d** CUE values predicted using a microbial model assimilating information on SOC profiles²¹. Data assimilation integrates observed data into predictive models to refine model parameters and improve estimation accuracy.

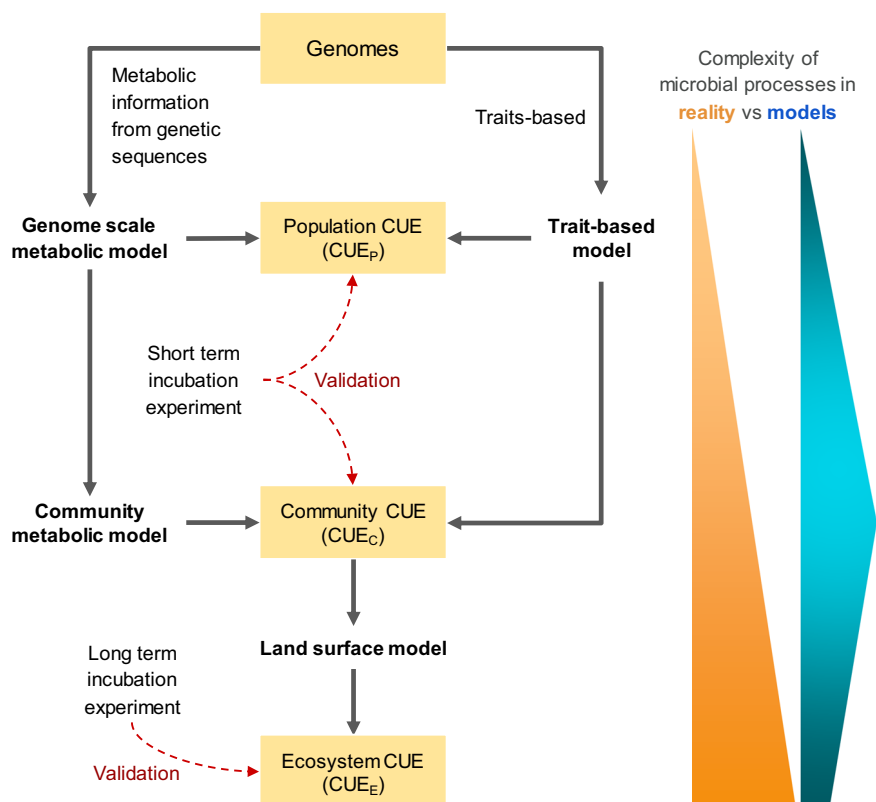


Fig. 2 | Schematic representation of a cluster of models integrating observational constraints on CUE at population (CUE_p), community (CUE_c) and ecosystem (CUE_e) scales. The genome-scale metabolic model predicts the movement of metabolites within a cell based on its genomic information. CUE_p and CUE_c can

be validated by short-term incubation measurements, while CUE_e requires long-term incubation measurements. Although the scales and processes governing CUE expand from individual cells to entire ecosystems, there is a practical limit to the extent they can be resolved in C cycle models.

forests consistently showing the lowest CUE values, regardless of the measurement approaches used^{19,20} (Fig. 1c). CUEs derived from data assimilation²¹ are also lower than those from more direct measurement approaches (Fig. 1d).

Several attempts have been made to reflect or incorporate CUE variations into models of litter²² or soil organic matter^{9,13} decomposition with the aim of assessing the implications for soil C cycling. For example, incorporating an empirically-derived negative relationship between microbial CUE and temperature into a microbial-explicit SOC model improved the simulation of contemporary soil C stocks²³. Zhang et al.²⁴ introduced the effects of substrate quality and soil fertility on microbial respiration, highlighting the joint control of litter quality and quantity on the steady-state SOC stocks. Wieder et al.²⁵ enhanced the understanding of CUE variation by including two types of decomposers with differing substrate preferences and CUE (Fig. 1b). These examples suggest that more realistic representations of microbial C transformations have the scope for improving model predictions of soil C^{23,26}. However, these predictions were poorly constrained by observational data, calling their reliability into question^{21,27,28}.

In this Perspective, we synthesize our understanding of CUE regulatory factors and databases for constraining numerical models, with the aim of clarifying complexities, addressing controversies, and providing a holistic perspective on pathways to adequately reflect CUE variations in C cycle models and their consequences for simulated soil C stocks.

Data availability and challenges

Terminology and definitions of microbial CUE

The concept of microbial CUE, the fraction of C uptake that is used to produce microbial biomass^{16–18}, is intuitively straightforward, but CUE definitions vary depending on the ecological processes involved,

measurement methods, and scales of biological organization (e.g., population, community and ecosystem)^{14,17}. Therefore, CUE can be regarded as an emergent parameter, encapsulating multiple processes within a single metric. It is useful in modeling as the number of processes that can be modeled is constrained by practical limitations (e.g., availability of data for calibration). Consequently, ecosystem models often simplify microbial process complexity, which in reality, escalates from the genomic to the ecosystem level (Fig. 2).

CUE is quantitatively expressed as the ratio of microbial growth (μ) to C uptake (U)^{16,29}, that is, $CUE = \mu/U$. This ratio encapsulates the efficiency with which microorganisms convert assimilated C into biomass. Microbial uptake involves C assimilation for growth (μ), respiration (R), and the secretion of extracellular enzymes and metabolites (EX). Geyer et al.¹⁴ introduced a nested conceptual framework for understanding CUE across different biological organization levels: population (CUE_p), community (CUE_c), and ecosystem (CUE_e). This framework is useful for integrating C fluxes mediated by soil microbes into models at various ecological scales (Fig. 2).

CUE_p reflects the species-specific functioning of microbial taxa (e.g., biosynthesis rate, exudate production) and thermodynamics of C substrate metabolism that limits the proportion of C uptake used for biosynthesis versus C lost from the cell (e.g., mineralized or exuded as metabolites). Typically measured in cultured populations, the CUE_p formula adjusts for respiration (R) and exudation (EX) losses from the uptake, expressed as $CUE_p = \frac{U-R-EX}{U}$. CUE_c incorporates additional environmental and community factors influencing microbial metabolism in natural communities consisting of multiple populations. It focuses on gross microbial production prior to the recursive substrate recycling of necromass and exudates, capturing the metabolic response of microbial communities to substrates over short durations (hours), and is similarly expressed as $CUE_c = \frac{U-R-EX}{U}$.

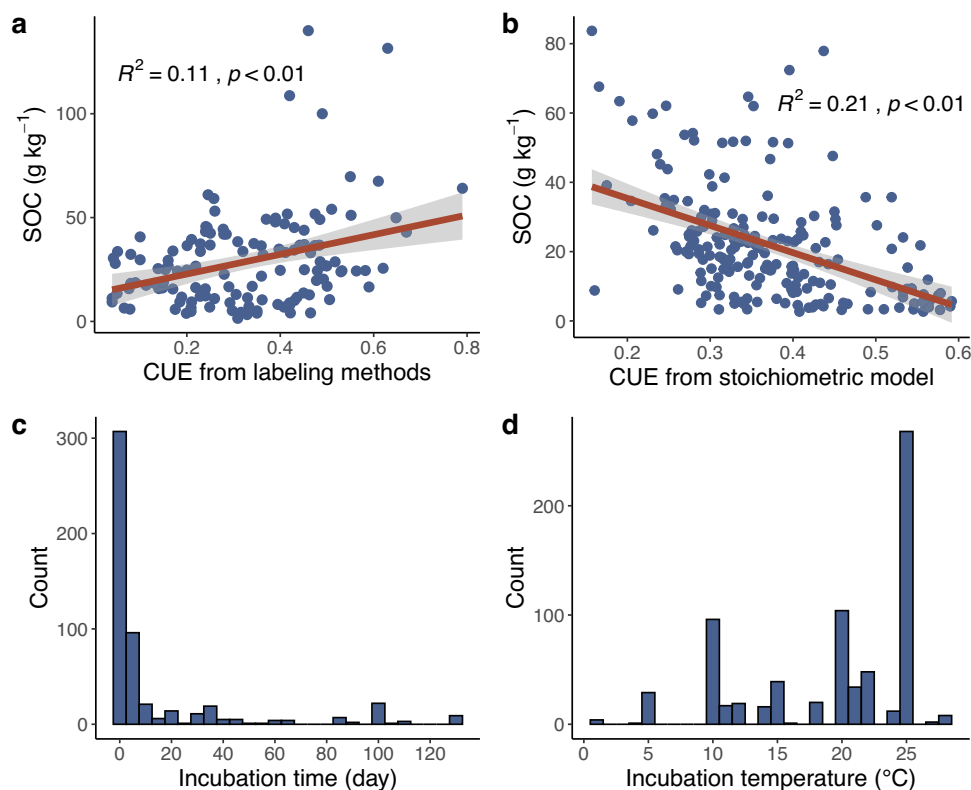


Fig. 3 | Impact of different research methods on the SOC-CUE relationship and variability in incubation conditions across studies. Panels **a** and **b** illustrate the relationships between soil organic carbon (SOC) concentration and CUE based on **a** isotopic labeling methods (^{14}C , ^{13}C -labeled substrates, and ^{18}O water) and

b stoichiometric modeling. Panel **c** presents the incubation durations, while panel **d** shows the temperatures employed in studies using labeling and incubation methods. Data sources: **a**²¹, **b**¹⁹, and **c**, **d**²⁰.

CUE_E considers C retention as net microbial growth over longer time scales (days to months), taking into account the drivers of CUE_P and CUE_C as well as microbial biomass turnover. On these time scales, a significant proportion of microbial biomass is converted to necromass following microbial death (MD)³⁰ such that $\text{CUE}_E = \frac{U - R - \text{EX} - \text{MD}}{U}$, encompassing all aspects of microbial C processing, including death and recycling processes.

Methods for measuring microbial CUE

Multiple approaches can be used to quantify CUE, such as isotopically labeling substrates^{31,32}, stoichiometric modeling^{22,33} and others³⁴. These methods rely on different assumptions and capture distinct microbial processes, which can explain the variability in CUE estimates across methods^{8,35,36} (Fig. 1a), including differences in the response of CUE to environmental changes³⁷, and the relationship between CUE and SOC (Fig. 3a, b).

The most common approach for measuring CUE is the tracking of isotopically labeled compounds (^{14}C , ^{13}C labeled substrate, or ^{18}O water) introduced to the system. Carbon isotopes in microbial substrates enable the differentiation between C allocated to microbial biomass and that released through respiration. Although this labeling technique is widely used, its results can be influenced by the choice and combination of substrates³¹, as well as the incubation period^{14,38}. A significant limitation of this approach is that measured CUE reflects only the efficiency of those microbes that use the introduced substrates, not the entire microbial community. Furthermore, the variation in incubation times and temperatures across different studies (Fig. 3c, d) presents a substantial obstacle to standardizing CUE measurements.

The method using ^{18}O -labeled water is based on the incorporation of the ^{18}O -atom into microbial DNA as a measure of growth as compared

to catabolic C losses as CO_2 ^{32,39}. This method has higher accuracy than the C labeling method as it is not substrate specific, does not perturb microbial metabolism like methods involving substrate addition, and exhibits comparatively less variability over time³⁵. Nonetheless, this method faces limitations such as higher cost and demanding technical procedures. Concerns also arise regarding the method's foundational assumptions, e.g., the presumption that water is the sole oxygen source for microbial DNA synthesis and the hypothesis that all microbial cells maintain a consistent DNA to biomass C ratio⁴⁰. Furthermore, its applicability in dry soils is challenging⁴¹.

Stoichiometric modeling is a common method for indirectly estimating CUE, which is based on the assumption that microbes growing on plant detritus allocate C to produce enzymes and other necessary components to acquire nutrients in the appropriate elemental ratios at the whole-community scale^{29,33}. This approach offers the advantage of requiring only a limited number of parameters, such as the activities of enzymes targeting C versus nitrogen (N) or phosphorus (P) acquisition and the C:N:P composition of the substrate and microbial biomass, which can be constrained by existing observations. However, it relies on highly simplified assumptions regarding elemental ratios and C allocation³⁶. This approach inherently suggests lower CUE in soils with high SOC due to its focus on the metabolic costs of nutrient acquisition under conditions where nutrients are scarce relative to C. This outcome (Fig. 3b) starkly contrasts with the positive correlation between CUE and SOC observed using isotopic labeling techniques (Fig. 3a), which are commonly considered to provide a more realistic insight into the relationship between CUE and SOC. The isotope labeling method estimates microbial growth and CUE by tracking the incorporation of labeled atoms into biomass or DNA, reflecting intracellular biochemical transformations. In contrast, the stoichiometry model method estimates CUE by analyzing the

activities of extracellular enzymes and the stoichiometric balance between organic matter and microbial biomass, focusing on extracellular metabolic processes⁴². Therefore, caution is advised when comparing results obtained from these two methods, even though they use the same term (CUE). We do not yet know the extent to which the stoichiometric and isotope methods are comparable. Until we understand which patterns can be accurately captured by the simpler stoichiometric method, we should rely on the more robust ¹⁸O method for measuring actual CUE and the ¹³C method for CUE associated with specific substrates.

In addition to the methods mentioned above, there are other less commonly used approaches, including the use of ¹⁸O in water vapor to minimize impact on soil moisture⁴¹, metabolic flux analysis¹⁷, and calorimetry⁴³. Each method offers unique advantages and faces specific limitations, grounded in their underlying assumptions and theoretical bases^{35–37}. These limitations not only affect the accuracy of these methods but also introduce significant comparability issues. Consequently, there is an urgent need to improve current methodologies and integrate innovative techniques to more accurately assess soil microbial CUE.

Data gap

Given the methodological challenges in measuring CUE in situ, field assessments of microbial CUE are rare. The vast majority of existing CUE observations have been obtained from lab incubations. Yet, these CUE observations remain scarce at the global scale, a situation which is exacerbated by the lack of harmonization of observations from different measurement approaches. For some ecosystems, observations are few or even nonexistent, including ecosystems that play a critical role in the global C cycle, such as tropical rainforests, wetlands, and peatlands^{44,45}.

Existing CUE measurements mostly come from studies of the litter and surface mineral soil¹⁶. Thus, our understanding of microbial CUE in subsurface soil remains limited, which is problematic as large amounts of C are stored in subsoils globally, and especially those of wetlands and peatlands. The few existing studies indicate that microbial CUE decreases with soil depth^{46,47} and that subsurface CUE may be less sensitive to warming³¹ but more sensitive to nutrient variations⁴⁸.

Moreover, data on temporal variations in CUE are lacking. A commonly overlooked factor that may contribute significantly to CUE variability in soil ecosystems, regardless of methodology, is seasonality in CUE. Seasonal changes are associated with significant variations in substrate availability, temperature and moisture, all of which may have a substantial impact on the growth and respiration of soil microorganisms, thereby altering microbial CUE³⁹. For example, CUE estimated using the ¹⁸O incorporation method ranged from 0.1 to 0.7 in soils from an agricultural field site and from 0.1 to 0.6 at a forest site within one year⁴⁹. It has also been reported that soil microbial CUE exhibits significant fluctuations within a short period (daily) after rewetting^{50,51}. This temporal dynamic in CUE values could contribute to the significant variability observed in CUE measurements.

Regulatory factors governing microbial CUE

The incorporation of soil microbial CUE dynamics into process-based models necessitates a comprehensive understanding of a range of regulatory factors influencing CUE (Fig. 4). CUE at a specific biological level is influenced by features of both the microbial community itself (biological controls) and its external environment (abiotic controls). These factors frequently interact, particularly at the community and ecosystem levels: abiotic controls can modify CUE_C or CUE_E by regulating biological controls, while biological controls may induce adaptation to abiotic factors, thereby influencing the impact of abiotic controls.

Biological controls

Microbial physiological state. Microbial CUE reflects the physiological state of microorganisms. Under natural conditions, only a small

proportion (values vary from 1% to >20% in different studies^{52,53}) of soil microbial cells are metabolically active, and soil respiration primarily originates from these metabolically active cells⁵³. Nonetheless, a high fraction of microbial cells in the soil are in a potentially active state (10 to 60% of the total microbial biomass), meaning that they are ready to start using available substrates within a few hours after easily available substrate is added. The shifts in physiological states of these microbial cells, resulting from changes in temperature, moisture, or substrate availability, significantly impact CUE⁵⁴. Consequently, CUE_P or CUE_C measurement methods relying on substrate addition may overestimate CUE⁴⁴, and shifts in physiological state can lead to seasonal variations in CUE⁴⁹.

Microbial community diversity and composition

Increased microbial diversity enriches the spectrum of metabolic functions within a community, potentially leading to greater microbial growth⁵⁵ and CUE_C by facilitating more efficient use of varied C sources^{10,56}. The composition of microbial communities, notably the ratio of fungal to bacterial biomass (F:B), plays a critical role in determining CUE_C⁵⁷. Communities dominated by fungi can show higher CUE_C, attributed to their higher biomass C to N) ratios (C:N) and their proficiency in decomposing complex organic materials⁵⁸, or lower CUE due to the high costs associated with resource acquisition by decomposer fungi⁵⁷. Therefore, this contrasting evidence from plant litter studies indicates that the relationship between F:B ratio and CUE is context-dependent^{57,59}. Alternatively, an approach categorizing microorganisms into copiotrophs (*r*-strategists with low CUE) versus oligotrophs (*K*-strategists with high CUE) has been promising for estimating CUE⁶⁰. For example, shifts from *r*-strategists to *K*-strategists explain increased CUE_C along a successional gradient in the south-eastern Tibetan Plateau⁶¹.

Changes in community composition may also enable microbial communities to alter their CUE in response to environmental changes or fluctuations^{62,63}. For instance, long-term warming experiments indicate a decline in the temperature sensitivity of CUE_C, suggesting that shifts in microbial composition can maintain CUE_C despite changes in temperature and substrate quality³¹. Similarly, modeling studies suggest that changing microbial community composition can reduce the sensitivity of CUE_C to substrate quality⁶⁴ and soil moisture fluctuations⁶⁵.

Biotic interactions

In the soil food web, biotic interactions such as mutualism, facilitation, competition, and predation can shape CUE_C⁵⁶. Interspecific microbial competition drives accelerated growth rates, accompanied by the release of secondary metabolites that can negatively affect CUE_C⁶⁶. Antagonistic interactions may trigger stress responses, further diminishing CUE_C⁶⁷. Conversely, facilitation enhances CUE_C by broadening species-realized niches, alleviating environmental stress, and reducing extracellular enzyme production costs⁶⁴. Biotic interactions at higher trophic levels, such as predation, can variably affect CUE_C by altering microbial density and influencing the outcomes of interspecific competition^{68,69}.

Abiotic controls

Temperature. Temperature significantly affects soil microbial CUE, with respiration often increasing more than growth in short-term incubations, resulting in a decrease in CUE_P^{9,34,70}. The impact on CUE_C and CUE_E is less clear⁶³, likely due to varied responses among microbial taxa^{71,72} and interactive effects with other environmental factors^{38,39,46,73}. Temperature shifts can lead to changes in community traits or select for taxa with distinct life strategies, known as trait modification and trait filtering, respectively^{74,75}. However, limited research on how CUE_P varies among different taxa in response to temperature impairs our ability to accurately predict changes in CUE_C^{76–78}.

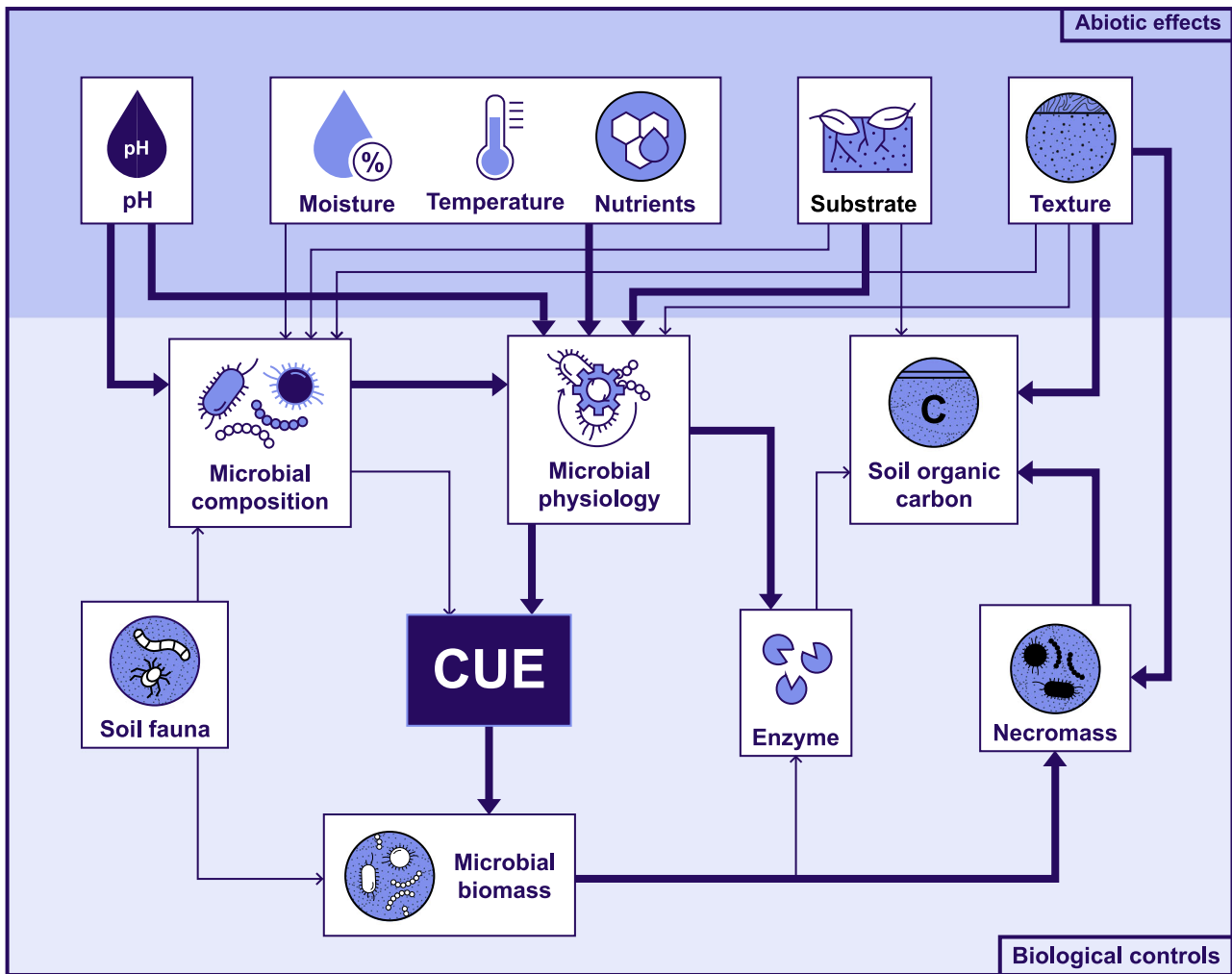


Fig. 4 | Framework of biological and abiotic determinants of CUE in a carbon cycle context. The darker-colored area in the figure indicates biological controls; the lighter-colored area indicates abiotic effects. The arrows depict implicit relationships and the width of the arrows corresponds to the levels of scientific

certainty: confident assertions are represented by thick lines, while less confident assertions are indicated by thinner lines. These confidence levels are based on the expertise of the authors.

The interplay between direct and indirect temperature effects on soil microbial CUE_C and CUE_E complicates our understanding of the impact of warming on CUE. Warming can intensify C-nutrient imbalances, potentially diminishing microbial CUE^{79} , but it can also improve the efficiency of substrate utilization, thereby enhancing $CUE^{32,72}$. Expected reductions in soil moisture due to increased evapotranspiration under warming conditions⁸⁰ add another layer of complexity, with the combined impacts of temperature and moisture on microbial CUE remaining inadequately explored^{10,81}. Some soil C models, including Millennium⁸² and MIMICS²⁵ have begun to account for the temperature dependency of CUE_C , indicating a growing recognition of the importance of including the dynamic response of microbial CUE to fluctuations in temperature.

Soil water availability

Increased soil moisture promotes microbial growth and CUE by improving substrate diffusivity and accessibility, and lowering investment in osmolyte synthesis, as long as conditions remainoxic^{8,10,83}. Prolonged water stress reduces soil substrate accessibility and increases the need to synthesize osmolytes to survive during dry periods, leading to lower CUE_C^{83} , even though the taxa that remain active in dry conditions can maintain relatively high growth rates⁸⁴. Furthermore, drought reduces plant C inputs to the soil⁸³, thus

potentially leaving microbes with fewer lower resources, resulting in lower CUE. The intricate interplay of drought-induced changes in microbial respiration and growth may leave CUE unchanged if the affected processes balance each other⁷⁸. High levels of soil moisture may also reduce microbial CUE. As soil pores fill with water, air spaces and oxygen diffusivity decline, potentially leading to anaerobic conditions if saturation occurs. Under O_2 limitation, soil microbes shift from aerobic to anaerobic respiration or fermentation, significantly reducing energy yield and leading to decreased microbial growth and CUE while having little impact on CO_2 production rate due to upregulated biochemical rates⁸³.

Microbial responses to rewetting of a dry soil also cause rapid changes in CUE, as shown in modeling studies⁵⁰ and confirmed by empirical evidence⁵¹. Upon rewetting, respiration increases while growth lags behind, especially when the soil has been dry for a long period⁵¹. As a result, just after rewetting, CUE is low and then increases as growth recovers during the first days after rewetting. However, after this initial pulse of microbial activity, CUE peaks and decreases again as substrates released during rewetting are consumed⁵¹.

Nutrient availability

The availability of nutrients such as N and P significantly affects microbial growth and respiration according to the concept of

stoichiometric homeostasis which assumes constrained biomass C:N:P ratios of microbial cells^{29,64}. Consequently, CUE decreases with increasing substrate C-to-nutrient ratios and increases with nutrient amendment when organic substrates are nutrient-poor^{22,29}. Several C cycle models, such as the one proposed by Manzoni et al.⁸⁵ and its later implementation²⁴, have integrated CUE dynamics as a function of stoichiometry. In contrast to the homeostasis concept, recent findings highlight the capability of microbes to store and use nutrients dynamically, contributing to a stable CUE across different environments by separating growth and respiration processes from immediate nutrient availability⁸⁶. This resilience to nutrient stress suggests that future C modeling should incorporate microbial nutrient storage dynamics for enhanced predictive accuracy.

Soil pH

Soil pH influences microbial CUE_C and CUE_E by affecting the bacterial community composition and acting as a potential stressor⁸⁷. It also impacts CUE by altering microbial community composition⁸⁸, nutrient solubility⁸³, and metal toxicity (e.g., aluminum⁸⁷). Habitats with neutral pH generally have higher bacterial diversity and biomass compared to acidic or alkaline soils⁷. The response of community composition to a shift in soil pH from acidic to neutral corresponded with a significant increase in CUE_C^{87,89}. However, recent research indicates a complex interplay between soil pH, microbial community composition, and CUE dynamics, evidenced by both negative correlations⁹⁰ and a U-shaped response curve, pinpointing a critical threshold at pH 6.4⁹¹, although the calculations to document this are complex and may necessitate refinement.

Soil texture and structure

Microbial growth is intricately linked to substrate accessibility, which is influenced by soil environmental conditions like texture and soil structure. Approximately 40–70% of soil bacteria are associated with microaggregates and clay particles⁹². The structural complexity of the soil environment also plays a crucial role in shaping the community structure and function of soil microorganisms at the ecosystem level⁹³. Heterogeneity of soil structure and composition creates diverse microhabitats that influence microbial interactions, diversity, distributions, and activity, as well as ecosystem processes like nutrient cycling and organic matter decomposition⁹⁴. Still, limited information exists on the relationship between soil texture or structure and microbial CUE. A recent meta-analysis found a significant positive link between microbial CUE_C or CUE_E for glucose and soil clay content⁹⁵, which was attributed to increased clay content enhancing substrate adsorption⁹⁶, thereby limiting substrate availability to microbes⁹⁷, and resulting in higher microbial CUE_C or CUE_E.

Substrate quality

Substrate quality, defined by the chemical characteristics of organic matter that influence its decomposability, such as the C:N ratio and molecular composition, significantly impacts soil microbial CUE⁹⁸. A “high-quality” substrate typically has a lower C:N ratio, indicating a balanced N content relative to C, and a lower content of recalcitrant compounds, which generally leads to faster decomposition and higher CUE by providing C and nutrients that microbes require for growth and metabolism⁸. Compounds requiring multiple enzymatic steps for degradation can lead to reduced efficiency in building biomass. Polymeric substrates like lignin and cellulose need depolymerization before cellular uptake, whereas smaller substrates readily diffuse across membranes⁶². Takriti et al. (2018) found a positive association between soil CUE_C and ratios of cellulase to phenol oxidase enzyme activity potential, which was considered to be indicative of soil organic matter (SOM) substrate quality⁴⁶. Different substrates necessitate distinct metabolic pathways, resulting in different respiration rates per unit C assimilated^{8,99}. Frey et al. (2013) observed lower microbial CUE_C

when soils were amended with oxalic acid or phenolic compounds compared to glucose, despite similar molecular sizes³¹.

Microbial CUE increases with the chemical energy per mole of C in the substrate, highlighting the importance of substrate chemistry for microbial CUE variability in soil⁸. This relationship is akin to the concept of energetic imbalance¹⁰⁰, which parallels the idea of stoichiometric imbalance. The energy content of soil microbial biomass and substrate can be quantified by the degree of reduction (γ), which refers to the average number of electrons available per C atom for biochemical reactions, indicating the energy density of the substrate or biomass⁸. The degree of reduction of soil microbial biomass (γ_B) is typically around 4.2, while that of substrate (γ_S) usually varies between 1 (e.g., for oxalate) and 8 (methane)⁸. Most of the substrates used by soil microorganisms have a γ_S of 3 (e.g., various organic acids), 4 (e.g., glucose and other carbohydrates), and rarely 5 or higher (e.g., leucine, polyhydroxyalkanoates or lipids)⁸. When γ_S is lower than γ_B , the substrate's energy content is insufficient to meet microbial demand, necessitating the oxidation of more substrate per unit of C assimilated, thereby reducing CUE¹⁰¹. These insights form the basis of the stoichiometric modeling for indirect CUE estimates.

SOC-CUE relationship

The relationship between CUE and SOC concentration at the ecosystem level can be positive, negative, or non-existent, depending on the interactions among multiple processes^{21,92,96,102–104}. Higher CUE can lead to increased SOC through biosynthesis and accumulation of microbial by-products—facilitating SOC formation via the entombing effect^{16,102,105}—or conversely, trigger SOC decline through the priming effect by ramping up microbial biomass and enzyme activity⁹. While some studies suggest a negative correlation between CUE and SOC^{103,104,106}, the majority of research supports a positive relationship^{21,74,107,108}, indicating that higher CUE is often linked to increased SOC levels. In a recent study, Tao et al.²¹ employed observational data and data assimilation algorithms and found that, on a global scale, CUE is positively correlated with SOC concentration, arguing for CUE as the major determinant for SOC formation. However, subsequent arguments have raised methodological concerns which might have obscured the importance of microbial community dynamics²⁷ and SOC stabilization processes¹⁰⁹.

Indeed, the link between microbial CUE and SOC is contingent upon the stabilization of microbial necromass within soil aggregates or its association with minerals^{96,102,105}. This stabilization process, pivotal for enhancing SOC, is significantly influenced by physico-chemical soil properties, which vary greatly and determine the potential for necromass protection^{110,111}. Positive SOC-CUE relationships could be anticipated in soils with high physicochemical C stabilization potential and microbial communities that convert simple chemical substrates into necromass¹¹¹. Conversely, when soil microbes face environmental stress, the relationship between CUE and SOC becomes less predictable. Particularly under conditions where nutrients are limited relative to carbon, the increased microbial respiration required to maintain stoichiometric balance leads to a decreased CUE^{29,33}. Further reductions in CUE may be driven by environmental challenges such as low oxygen or pH^{88,106}, as well as the physiological costs of microbial competition⁶⁶. However, these stressors on microbial activity may differently affect SOC, potentially leading to either a negative or negligible correlation between CUE and SOC¹⁰⁶. It's worth noting that in organic-rich soils, such as peat, C stabilization relies more on the accumulation of undecomposed plant material than on necromass formation¹¹², making the link between CUE and SOC less direct. Therefore, the CUE-SOC relationship in organic soils is expected to differ from mineral soils where C is mainly stabilized by mineral associations.

Additionally, it is important to recognize the distinct sensitivities of microbial CUE and SOC to environmental changes, as their

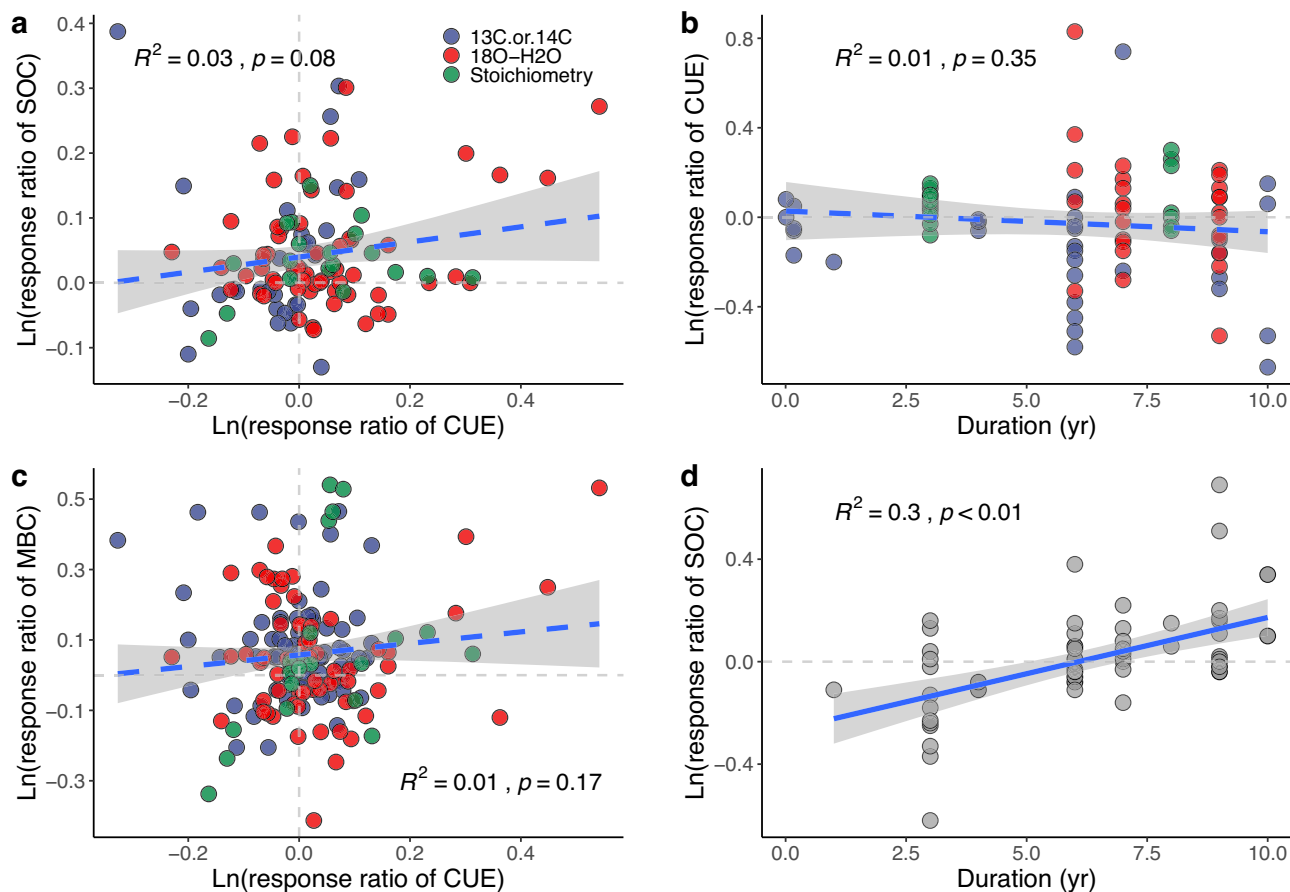


Fig. 5 | Contrasting responses of SOC and CUE to fertilization. Correlations between ln-transformed response ratios of microbial CUE and ln-transformed response ratios of (a) SOC and (c) microbial biomass C (MBC); and the correlation between experiment duration and ln-transformed response ratios of (b) CUE and (d) SOC. The response ratio is calculated as the ratio of the measured value in

treatment to the value in the control. Data are from meta-analyses^{27,37,114}. Both datasets include observations from all three methods of CUE measurement, i.e., C labeling, O labeling, and stoichiometry modeling as indicated by symbol colors in a–c.

responses are not synchronized. Microbial CUE can adjust rapidly, from days to months, in contrast to SOC, which may take years or even decades to respond to a measurable extent^{49,113}. Data from two meta-analyses highlight this disparity, showing that although fertilization positively affects both CUE_C and SOC^{37,114}, the response ratios of CUE_C were not significantly correlated with the response ratios of SOC, or even microbial biomass C content (Fig. 5a, c). Here, the “response ratio” is calculated as the ratio of the measured value in the treatment to the value in the control. Furthermore, the response ratios of microbial CUE_C were not significantly related to treatment duration (within ten years of treatment) (Fig. 5b), whereas the response ratios of SOC increased significantly with experiment duration (Fig. 5d). Therefore, SOC gradually approaches a new equilibrium over several decades, whereas CUE achieves equilibrium almost immediately. This discrepancy underscores the importance of considering the state (SOC and microbial biomass) dynamics of an ecosystem when evaluating the interplay between microbial CUE and SOC dynamics.

Using models and data across scales to clarify the microbial role in C cycling

Integrating genomic data with CUE and C models. With the rise of high throughput sequencing technology, the use of genomic datasets to help calibrate or validate C models has become both feasible and affordable. This capacity is especially valuable when predicting CUE¹¹⁵. As genomic data related to microbial traits becomes more readily available at both the population¹¹⁶ and community levels through metagenomics¹¹⁷, there is a growing need to effectively integrate this

data into C cycle models. This integration requires models that can handle complex microbial interactions, from individual populations to entire communities (Fig. 2).

One way to integrate genomic data is by converting the genetic sequences of microbes into information on metabolic pathways (e.g., cellulose degradation, lignin degradation, nitrogen reduction, and fermentation) using genome-scale metabolic models (GEMs)¹¹⁸. GEMs take into account the microbe’s environment, such as substrate availability, and predict the transformation of metabolites within a cell based on its genomic information. This process allows for the calculation of CUE at the population level by analyzing substrate use and CO₂ production¹¹⁸. For community-level CUE, GEMs can be combined into microbial community models that simulate interactions between different microbial taxa: The ‘computation of microbial ecosystems in time and space metabolic modeling platform’ (COMETS) extends GEMs to include dynamics of microbial growth and interactions, providing a tool for predicting CUE_C under various environmental conditions¹⁵.

An alternative modeling approach at the community level is based on traits (e.g., quantity of cellulase produced, maximum rate of reaction (V_{max}) of cellulose decay by cellulase, V_{max} of cellulose-monomer uptake, and turnover rate), such as the DEMENT model, which uses data on microbial traits to simulate substrate use and CO₂ production¹¹⁹. This model can predict both CUE_P and CUE_C under different environmental conditions and over time. However, translating genomic data into traits remains challenging¹²⁰. Genomic datasets typically indicate the presence or absence of certain genes or

pathways, but additional information, such as that from GEMs or experimental data, is necessary to accurately map these genes to functional traits in the models.

Validating genomic and trait-based models is crucial and can be achieved using community-level genomic datasets, which offer insights into microbial strategies that affect CUE, such as nutrient recycling and stress tolerance^{117,121}. Combining these models with traditional CUE measurements and omics data allows for the creation of detailed maps of community-level CUE, offering new insights into C cycling dynamics and providing input information for C cycle models.

A major challenge in this field is the high computational demand of integrating omic data into complex models. One solution is the development of computational emulators that can simulate the dynamics of microbial models more efficiently, bridging the gap between detailed, small-scale models and broader applications in C cycle studies¹²². This approach promises to improve our understanding of microbial contributions to C cycling, leveraging the power of genomic data to inform and validate complex ESMs.

Harmonization of CUE measurements and aligning measured and modeled CUE

Harmonizing soil microbial CUE measurements across different methods, i.e., aligning results from different methodologies, poses a challenge due to the differences across measurement techniques. While adopting a universal protocol for CUE measurement—a single, standardized measurement method—would be ideal, it may not be feasible given the complexities of CUE. Therefore, a more practical approach involves providing a clear and comprehensive description of the methodologies used in different studies. This detailed reporting should include information on the physiological processes considered, such as maintenance, enzyme production, biomass generation, and mortality rates. This level of detail helps in understanding and comparing results across studies, as well as in selecting appropriate data for model calibration¹⁷.

In contemporary soil C models that explicitly incorporate microbial processes^{25,82}, the CUE is close to empirically measured CUE_C. To achieve a uniform approach to CUE measurement, microbial models that resolve key processes influencing CUE, such as uptake, respiration, exudation, and microbial death could be used¹⁷. Such models can generate CUE metrics that align with different measurement methodologies by incorporating a complete or partial set of these processes into their calculations. Furthermore, these models can be adapted to conduct numerical experiments with specific substrates or to incorporate isotopic tracers (e.g., ¹³C, ¹⁴C, ¹⁸O) to simulate outcomes from labeling experiments. This adaptability allows for the exploration of hypotheses regarding discrepancies in measurements under diverse conditions by modifying model boundary conditions. Additionally, microbial models serve as foundational tools for integrating microbial metabolism into broader global C models, potentially enhanced by machine learning emulators for improved scalability and applicability.

Constraining CUE using model-data fusion

Data assimilation encompasses a collection of techniques, including Bayesian inference, that refine biogeochemical models by integrating observational data. This process not only updates model parameters to reflect the most likely values based on available data but also quantifies their uncertainties, thus bridging the gap between empirical observations and theoretical models¹⁰⁷. This approach is particularly valuable for parameters like microbial CUE, which are challenging to measure directly in the field due to technical limitations. An innovative application of data assimilation is demonstrated by Tao et al.²¹, who developed the PROcess-guided deep learning and DATA-driven (PRODA) approach^{123,124}. This method integrates global-scale SOC data with a microbially explicit model to produce a global map of microbial CUE. PRODA employs traditional Bayesian data assimilation

to estimate parameters at specific sites and then uses deep learning to extrapolate these site-specific parameter estimates to a global scale. The result is a set of parameters that optimally align with observed data, offering a detailed view of microbial CUE and SOC storage patterns worldwide, along with other soil C cycle dynamics such as decomposition rates, environmental impacts on soil respiration, and vertical C transport²¹.

Despite the potential of approaches like PRODA to harness large datasets for enhancing our understanding of the soil C cycle, their computational intensity—stemming from the extensive data sampling required by Bayesian inference—may limit their application in models with complex structures. The next wave of data assimilation techniques will likely integrate process-based models with deep learning algorithms more seamlessly¹²¹. Such advancements could offer quicker parameter optimization and facilitate comparisons across different models, paving the way for more accurate and comprehensive assessments of microbial CUE and C cycle dynamics on a global scale.

Long-term SOC records and ecosystem manipulation experiments

Ecosystem manipulation experiments and observations of natural gradients offer invaluable insights into how microbial communities and CUE adapt to global change factors. Especially insightful are field experiments (or studies leveraging natural gradients) that alter environmental factors such as soil temperature, precipitation patterns, or nutrient levels^{76,125} over long durations. These experiments provide critical data on the enduring effects of global change drivers on CUE, while simultaneously highlighting the limitations of current models and enhancing our comprehension of ecological processes. Integrating the results from these experiments with model simulations, supported by proven site modeling protocols and extra observational data, is crucial for steadily enhancing the accuracy and complexity of models¹²⁶.

Incorporating radiocarbon (¹⁴C) data and long-term SOC records into models is also vital for refining CUE forecasts across longer (decadal to centennial) time scales. This temporal information is essential for capturing the dynamics of CUE over time, thereby improving the precision of models in depicting spatial and temporal fluctuations¹²⁷.

Diagnosing CUE from existing models or simulation archives

In global C modeling, approaches to quantify the environmental impact on organic matter decomposition and stabilization differ significantly. An effective method for estimating microbial CUE at the ecosystem level as emerging from model simulations involves the calculation of the ratio between soil heterotrophic respiration (R) and gross decomposition (D) within these models. Gross decomposition refers to the sum of all C fluxes transferred between the modeled soil C pools that are mediated by microbial processes, excluding physically mediated transfers (e.g., sorption, aggregation, or leaching). This includes all C removed from organic matter pools, whether it is lost as CO₂ or transferred to another pool (SI-Text 1). This ratio effectively quantifies microbial-mediated C losses from SOC pools, integrating both growth (anabolic processes) and respiration (catabolic processes). Under steady-state conditions, it is assumed that heterotrophic respiration aligns with microbial C uptake, resulting in the formula: CUE = 1 - R/D. The steady-state assumption implies that microbial communities and SOC stock are stable in time (i.e., in equilibrium with boundary conditions). This is an approximation of real systems where SOC varies due to anthropogenic and natural changes (e.g., Holocene climatic variations). This diagnosed CUE, emerging as a property inherent to the model, is not susceptible to the equifinality issues that can affect the underlying intrinsic model parameters (like CUE_C), and it does not necessitate the incorporation of explicitly microbial models, offering a simplified yet insightful metric. These

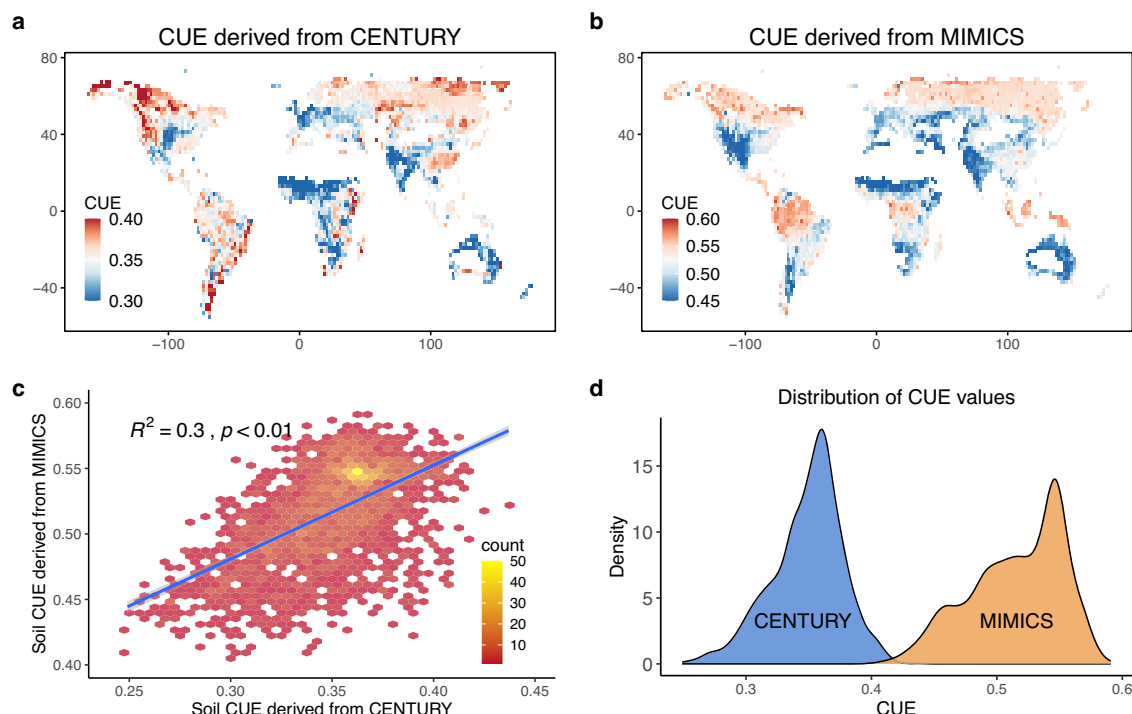


Fig. 6 | Diagnosed CUE from two existing soil C models. CUE diagnosed from a nutrient-enabled version of the the Organizing Carbon and Hydrology In Dynamic Ecosystems land surface model (ORCHIDEE-CNP) deploying a soil module based on (a) the CENTURY model¹²⁸, or (b) the MIMICS model with constant intrinsic CUE_C¹³². c Correlation between diagnosed CUE values from the CENTURY-based model and the MIMICS-based model. d Distribution frequency of CUE for the two scenarios.

model-based CUE estimates, derived from long-term flux averages (e.g., 20 years), represent stable C stocks. In contrast, measurement-based estimates, taken over shorter periods, are more susceptible to significant CUE variations due to asynchronous fluctuations in components such as respiration and degradation, potentially introducing estimation inaccuracies. This timescale discrepancy likely accounts for the greater variability observed in measurement-based CUE compared to model-based CUE. We propose this “model-diagnosed CUE” as a novel metric, designed to estimate microbial CUE from model outputs without direct measurements of microbial uptake.

Analyzing diagnosed CUE and its relationship with SOC across various models, such as those evaluated in the Trends in the land carbon cycle (TRENDY) model intercomparison project², facilitates the identification of differences attributable to unique model structures and assumptions. For example, warming-induced CO₂ emissions should be higher in models with low diagnosed CUE compared to high CUE as the warming-induced stimulation of microbial activity will result in relatively more C being respired than cycled within the soil systems. This approach further allows the benchmarking and subsequent refinement of diagnosed CUE estimates using observed CUE_E data.

For instance, we derived CUE estimates from simulations conducted with two different versions of the Organizing Carbon and Hydrology In Dynamic Ecosystems (ORCHIDEE) land surface model¹²⁸, which differ in the SOC model deployed. The CENTURY SOC model (Fig. S1), which is widely used but does not resolve microbial processes, uses first-order decay, while the MIMICS model (Fig. S2) resolves microbial physiology, providing a more mechanistic understanding of microbial processes. The resulting global CUE maps (the average of simulation results over 20 consecutive years) revealed significant spatial variability (Fig. 6a, b). While the two maps showed a good correlation (Fig. 6c), the CUE values diagnosed from the MIMICS model were higher than those from the CENTURY model (Fig. 6d). These findings underscore the importance of incorporating

observational data into model calibration efforts to enhance the accuracy and reliability of SOC predictions by realistically resolving CUE.

In conclusion, the inherent structure of a model significantly shapes its outcomes, making the integration of empirical data with data-constrained models a fundamental step toward realistic predictions^{129,130}. Precisely delineating the spatial and temporal dynamics of CUE in models that specifically address microbial activities is crucial for the reliability of their predictions of SOC status and dynamics. Moreover, future soil C models must navigate the intricate balance between the complex regulatory mechanisms of CUE, other processes governing SOC formation and stabilization, and the practicality of model use to promote more precise projections of CUE responses under diverse environmental scenarios. This Perspective underscores the importance of combining different data sources with sophisticated modeling techniques to refine global CUE predictions. By incorporating genomic data, standardizing measurement protocols, applying data assimilation practices and critically evaluating CUE within existing frameworks, our comprehension of the global dynamics of microbial CUE can be markedly improved. This Perspective provides a roadmap for establishing an effective modeling approach to accurately represent global soil microbial CUE and its interactions with other biological and abiotic processes that regulate SOC dynamics.

References

1. Eyring, V. et al. Overview of the Coupled Model Intercomparison Project Phase 6 (CMIP6) experimental design and organization. *Geosci. Model Dev.* **9**, 1937–1958 (2016).
2. Friedlingstein, P. et al. Global carbon budget 2023. *Earth Syst. Sci. Data* **15**, 5301–5369 (2023).
3. Shi, Z. et al. Global-scale convergence obscures inconsistencies in soil carbon change predicted by earth system models. *AGU Adv.* **5**, e2023AV001068 (2024).

4. Varney, R. M., Chadburn, S. E., Burke, E. J. & Cox, P. M. Evaluation of soil carbon simulation in CMIP6 Earth system models. *Biogeosciences* **19**, 4671–4704 (2022).
5. Crowther, T. W. et al. The global soil community and its influence on biogeochemistry. *Science* **365**, eaav0550 (2019).
6. Ranheim Sveen, T., Hannula, S. E. & Bahram, M. Microbial regulation of feedbacks to ecosystem change. *Trends Microbiol.* S0966842X23001919 (2023) <https://doi.org/10.1016/j.tim.2023.06.006> (2023).
7. Fierer, N. Embracing the unknown: disentangling the complexities of the soil microbiome. *Nat. Rev. Microbiol.* **15**, 579–590 (2017).
8. Manzoni, S., Taylor, P., Richter, A., Porporato, A. & Ågren, G. I. Environmental and stoichiometric controls on microbial carbon-use efficiency in soils. *New Phytol.* **196**, 79–91 (2012).
9. Allison, S. D., Wallenstein, M. D. & Bradford, M. A. Soil-carbon response to warming dependent on microbial physiology. *Nat. Geosci.* **3**, 336–340 (2010).
10. Domeignoz-Horta, L. A. et al. Microbial diversity drives carbon use efficiency in a model soil. *Nat. Commun.* **11**, 3684 (2020).
11. Karhu, K. et al. Temperature sensitivity of soil respiration rates enhanced by microbial community response. *Nature* **513**, 81–84 (2014).
12. Luo, Z. et al. Convergent modelling of past soil organic carbon stocks but divergent projections. *Biogeosciences* **12**, 4373–4383 (2015).
13. Wieder, W. R., Cleveland, C. C., Smith, W. K. & Todd-Brown, K. Future productivity and carbon storage limited by terrestrial nutrient availability. *Nat. Geosci.* **8**, 441–444 (2015).
14. Geyer, K. M., Kyker-Snowman, E., Grandy, A. S. & Frey, S. D. Microbial carbon use efficiency: accounting for population, community, and ecosystem-scale controls over the fate of metabolized organic matter. *Biogeochemistry* **127**, 173–188 (2016).
15. Treseder, K. K. et al. Integrating microbial ecology into ecosystem models: challenges and priorities. *Biogeochemistry* **109**, 7–18 (2012).
16. Manzoni, S. et al. Reviews and syntheses: carbon use efficiency from organisms to ecosystems—definitions, theories, and empirical evidence. *Biogeosciences* **15**, 5929–5949 (2018).
17. Dijkstra, P. et al. On maintenance and metabolisms in soil microbial communities. *Plant Soil* **476**, 385–396 (2022).
18. Hagerty, S. B., Allison, S. D. & Schimel, J. P. Evaluating soil microbial carbon use efficiency explicitly as a function of cellular processes: implications for measurements and models. *Biogeochemistry* **140**, 269–283 (2018).
19. He, P., Zhang, Y., Shen, Q., Ling, N. & Nan, Z. Microbial carbon use efficiency in different ecosystems: a meta-analysis based on a biogeochemical equilibrium model. *Glob. Change Biol.* **00**, 1–17 (2023).
20. Qiao, Y. et al. Global variation of soil microbial carbon-use efficiency in relation to growth temperature and substrate supply. *Sci. Rep.* **9**, 5621 (2019).
21. Tao, F. et al. Microbial carbon use efficiency promotes global soil carbon storage. *Nature* <https://doi.org/10.1038/s41586-023-06042-3> (2023).
22. Manzoni, S. Flexible carbon-use efficiency across litter types and during decomposition partly compensates nutrient imbalances—results from analytical stoichiometric models. *Front. Microbiol.* **8**, 661 (2017).
23. Wieder, W. R., Bonan, G. B. & Allison, S. D. Global soil carbon projections are improved by modelling microbial processes. *Nat. Clim. Change* **3**, 909–912 (2013).
24. Zhang, H. et al. Modeling the effects of litter stoichiometry and soil mineral N availability on soil organic matter formation using CENTURY-CUE (v1.0). *Geosci. Model Dev.* **11**, 4779–4796 (2018).
25. Wieder, W. R., Grandy, A. S., Kallenbach, C. M. & Bonan, G. B. Integrating microbial physiology and physio-chemical principles in soils with the Microbial-Mineral Carbon Stabilization (MIMICS) model. *Biogeosciences* **11**, 3899–3917 (2014).
26. Sulman, B. N., Phillips, R. P., Oishi, A. C., Shevliakova, E. & Pacala, S. W. Microbe-driven turnover offsets mineral-mediated storage of soil carbon under elevated CO₂. *Nat. Clim. Change* **4**, 1099–1102 (2014).
27. He, X. et al. Model uncertainty obscures major driver of soil carbon. *Nature* **627**, E1–E3 (2024).
28. Shi, Z., Crowell, S., Luo, Y. & Moore, B. Model structures amplify uncertainty in predicted soil carbon responses to climate change. *Nat. Commun.* **9**, 2171 (2018).
29. Sinsabaugh, R. L. et al. Stoichiometry of microbial carbon use efficiency in soils. *Ecol. Monogr.* **86**, 172–189 (2016).
30. Camenzind, T., Mason-Jones, K., Mansour, I., Rillig, M. C. & Lehmann, J. Formation of necromass-derived soil organic carbon determined by microbial death pathways. *Nat. Geosci.* **16**, 115–122 (2023).
31. Frey, S. D., Lee, J., Melillo, J. M. & Six, J. The temperature response of soil microbial efficiency and its feedback to climate. *Nat. Clim. Change* **3**, 395–398 (2013).
32. Spohn, M., Klaus, K., Wanek, W. & Richter, A. Microbial carbon use efficiency and biomass turnover times depending on soil depth—implications for carbon cycling. *Soil Biol. Biochem.* **96**, 74–81 (2016).
33. Sinsabaugh, R. L., Manzoni, S., Moorhead, D. L. & Richter, A. Carbon use efficiency of microbial communities: stoichiometry, methodology and modelling. *Ecol. Lett.* **16**, 930–939 (2013).
34. Zhang, Q., Qin, W., Feng, J. & Zhu, B. Responses of soil microbial carbon use efficiency to warming: review and prospects. *Soil Ecol. Lett.* **4**, 307–318 (2022).
35. Geyer, K. M., Dijkstra, P., Sinsabaugh, R. & Frey, S. D. Clarifying the interpretation of carbon use efficiency in soil through methods comparison. *Soil Biol. Biochem.* **128**, 79–88 (2019).
36. Schimel, J., Weintraub, M. N. & Moorhead, D. Estimating microbial carbon use efficiency in soil: Isotope-based and enzyme-based methods measure fundamentally different aspects of microbial resource use. *Soil Biol. Biochem.* **169**, 108677 (2022).
37. Hu, J., Huang, C., Zhou, S. & Kuzyakov, Y. Nitrogen addition to soil affects microbial carbon use efficiency: meta-analysis of similarities and differences in ¹³C and ¹⁸O approaches. *Glob. Change Biol.* **28**, 4977–4988 (2022).
38. Hagerty, S. B. et al. Accelerated microbial turnover but constant growth efficiency with warming in soil. *Nat. Clim. Change* **4**, 903–906 (2014).
39. Simon, E. et al. Microbial growth and carbon use efficiency show seasonal responses in a multifactorial climate change experiment. *Commun. Biol.* **3**, 584 (2020).
40. Qu, L., Wang, C. & Bai, E. Evaluation of the 18O-H₂O incubation method for measurement of soil microbial carbon use efficiency. *Soil Biol. Biochem.* **145**, 107802 (2020).
41. Canarini, A. et al. Quantifying microbial growth and carbon use efficiency in dry soil environments via ¹⁸O water vapor equilibration. *Glob. Change Biol.* **26**, 5333–5341 (2020).
42. Sun, L. et al. Interpreting the differences in microbial carbon and nitrogen use efficiencies estimated by 18O labeling and ecoenzyme stoichiometry. *Geoderma* **444**, 116856 (2024).
43. Yang, S. et al. Enhancing insights: exploring the information content of calorespirometric ratio in dynamic soil microbial growth processes through calorimetry. *Front. Microbiol.* **15**, 1321059 (2024).
44. Fewster, R. E. et al. Imminent loss of climate space for permafrost peatlands in Europe and Western Siberia. *Nat. Clim. Change* **12**, 373–379 (2022).
45. Hugelius, G. et al. Large stocks of peatland carbon and nitrogen are vulnerable to permafrost thaw. *Proc. Natl. Acad. Sci. USA* **117**, 20438–20446 (2020).

46. Takriti, M. et al. Soil organic matter quality exerts a stronger control than stoichiometry on microbial substrate use efficiency along a latitudinal transect. *Soil Biol. Biochem.* **121**, 212–220 (2018).
47. Zhang, Q. et al. Whole-soil-profile warming does not change microbial carbon use efficiency in surface and deep soils. *Proc. Natl. Acad. Sci. USA* **120**, e2302190120 (2023).
48. Jiang, Y. et al. Deep soil microbial carbon use efficiency responds stronger to nitrogen deposition than top soil in tropical forests, southern China. *Plant Soil* <https://doi.org/10.1007/s11104-024-06509-w> (2024).
49. Schnecker, J. et al. Seasonal dynamics of soil microbial growth, respiration, biomass, and carbon use efficiency in temperate soils. *Geoderma* **440**, 116693 (2023).
50. Brangari, A. C., Manzoni, S. & Rousk, J. A soil microbial model to analyze decoupled microbial growth and respiration during soil drying and rewetting. *Soil Biol. Biochem.* **148**, 107871 (2020).
51. Li, X., Leizeaga, A., Rousk, J., Hugelius, G. & Manzoni, S. Drying intensity and acidity slow down microbial growth recovery after rewetting dry soils. *Soil Biol. Biochem.* **184**, 109115 (2023).
52. Couradeau, E. et al. Probing the active fraction of soil microbiomes using BONCAT-FACS. *Nat. Commun.* **10**, 2770 (2019).
53. Blagodatskaya, E. & Kuzyakov, Y. Active microorganisms in soil: critical review of estimation criteria and approaches. *Soil Biol. Biochem.* **67**, 192–211 (2013).
54. Hasby, F. A., Barbi, F., Manzoni, S. & Lindahl, B. D. Transcriptomic markers of fungal growth, respiration and carbon-use efficiency. *FEMS Microbiol. Lett.* **368**, fnab100 (2021).
55. Khurana, S. et al. Interactive effects of microbial functional diversity and carbon availability on decomposition—a theoretical exploration. *Ecol. Model.* **486**, 110507 (2023).
56. Anthony, M. A., Crowther, T. W., Maynard, D. S., Van Den Hoogen, J. & Averill, C. Distinct assembly processes and microbial communities constrain soil organic carbon formation. *One Earth* **2**, 349–360 (2020).
57. Soares, M. & Rousk, J. Microbial growth and carbon use efficiency in soil: Links to fungal-bacterial dominance, SOC-quality and stoichiometry. *Soil Biol. Biochem.* **131**, 195–205 (2019).
58. Malik, A. A. et al. Soil fungal:bacterial ratios are linked to altered carbon cycling. *Front. Microbiol.* **7**, 1247 (2016).
59. Keiblinger, K. M. et al. The effect of resource quantity and resource stoichiometry on microbial carbon-use-efficiency: resource quantity/quality drives microbial C-use-efficiency. *FEMS Microbiol. Ecol.* <https://doi.org/10.1111/j.1574-6941.2010.00912.x> (2010).
60. Six, J., Frey, S. D., Thiet, R. K., & Batten, K. M. Bacterial and fungal contributions to carbon sequestration in agroecosystems. *Soil Sci. Soc. Am. J.* **70**, 555–569 (2006).
61. Ma, S., Zhu, W., Wang, W., Li, X. & Sheng, Z. Microbial assemblies with distinct trophic strategies drive changes in soil microbial carbon use efficiency along vegetation primary succession in a glacier retreat area of the southeastern Tibetan Plateau. *Sci. Total Environ.* **867**, 161587 (2023).
62. Allison, S. D. Modeling adaptation of carbon use efficiency in microbial communities. *Front. Microbiol.* **5**, 571 (2014).
63. Qu, L. et al. Stronger compensatory thermal adaptation of soil microbial respiration with higher substrate availability. *ISME J. wrae025*. <https://doi.org/10.1093/ismejo/wrae025> (2024).
64. Kaiser, C., Franklin, O., Dieckmann, U. & Richter, A. Microbial community dynamics alleviate stoichiometric constraints during litter decay. *Ecol. Lett.* **17**, 680–690 (2014).
65. Brangari, A. C., Manzoni, S. & Rousk, J. The mechanisms underpinning microbial resilience to drying and rewetting—a model analysis. *Soil Biol. Biochem.* **162**, 108400 (2021).
66. Maynard, D. S., Crowther, T. W. & Bradford, M. A. Fungal interactions reduce carbon use efficiency. *Ecol. Lett.* **20**, 1034–1042 (2017).
67. Iven, H., Walker, T. W. N. & Anthony, M. Biotic interactions in soil are underestimated drivers of microbial carbon use efficiency. *Curr. Microbiol.* **80**, 13 (2023).
68. Frey, S. D. Protozoan grazing affects estimates of carbon utilization efficiency of the soil microbial community. *Soil Biol. Biochem.* **33**, 1759–1768 (2001).
69. Ma, L. et al. Long-term conservation tillage enhances microbial carbon use efficiency by altering multitrophic interactions in soil. *Sci. Total Environ.* **915**, 170018 (2024).
70. Tian, W. et al. Thermal adaptation occurs in the respiration and growth of widely distributed bacteria. *Glob. Change Biol.* **28**, 2820–2829 (2022).
71. Pold, G. et al. Carbon use efficiency and its temperature sensitivity covary in soil bacteria. *mBio* **11**, e02293–19 (2020).
72. Tian, J. et al. Microbially mediated mechanisms underlie soil carbon accrual by conservation agriculture under decade-long warming. *Nat. Commun.* **15**, 377 (2024).
73. Walker, T. W. N. et al. Microbial temperature sensitivity and biomass change explain soil carbon loss with warming. *Nat. Clim. Change* **8**, 885–889 (2018).
74. Kallenbach, C. M., Wallenstein, M. D., Schipanski, M. E. & Grandy, A. S. Managing agroecosystems for soil microbial carbon use efficiency: ecological unknowns, potential outcomes, and a path forward. *Front. Microbiol.* **10**, 1146 (2019).
75. Ye, J., Bradford, M. A., Maestre, F. T., Li, F. & García-Palacios, P. Compensatory thermal adaptation of soil microbial respiration rates in global croplands. *Glob. Biogeochem. Cycles* **34**, e2019GB006507 (2020).
76. Metze, D. et al. Soil warming increases the number of growing bacterial taxa but not their growth rates. *Sci. Adv.* **10**, eadk6295 (2024).
77. Saifuddin, M., Bhatnagar, J. M., Segrè, D. & Finzi, A. C. Microbial carbon use efficiency predicted from genome-scale metabolic models. *Nat. Commun.* **10**, 3568 (2019).
78. Smith, T. P., Clegg, T., Bell, T. & Pawar, S. Systematic variation in the temperature dependence of bacterial carbon use efficiency. *Ecol. Lett.* **24**, 2123–2133 (2021).
79. Sun, Y. et al. A global meta-analysis on the responses of C and N concentrations to warming in terrestrial ecosystems. *CATENA* **208**, 105762 (2022).
80. Xu, W. et al. A meta-analysis of the response of soil moisture to experimental warming. *Environ. Res. Lett.* **8**, 044027 (2013).
81. Fuchs-lueger, L. et al. Microbial carbon and nitrogen cycling responses to drought and temperature in differently managed mountain grasslands. *Soil Biol. Biochem.* **135**, 144–153 (2019).
82. Abramoff, R. Z. et al. Improved global-scale predictions of soil carbon stocks with Millennial Version 2. *Soil Biol. Biochem.* **164**, 108466 (2022).
83. Zheng, Q. et al. Growth explains microbial carbon use efficiency across soils differing in land use and geology. *Soil Biol. Biochem.* **128**, 45–55 (2019).
84. Metze, D. et al. Microbial growth under drought is confined to distinct taxa and modified by potential future climate conditions. *Nat. Commun.* **14**, 5895 (2023).
85. Manzoni, S. Optimal metabolic regulation along resource stoichiometry gradients. *Ecol. Lett.* **20**, 1182–1191 (2017).
86. Mason-Jones, K., Breidenbach, A., Dyckmans, J., Banfield, C. C. & Dippold, M. A. Intracellular carbon storage by microorganisms is an overlooked pathway of biomass growth. *Nat. Commun.* **14**, 2240 (2023).
87. Jones, D. L., Cooledge, E. C., Hoyle, F. C., Griffiths, R. I. & Murphy, D. V. pH and exchangeable aluminum are major regulators of microbial energy flow and carbon use efficiency in soil microbial communities. *Soil Biol. Biochem.* **138**, 107584 (2019).

88. Malik, A. A. et al. Land use driven change in soil pH affects microbial carbon cycling processes. *Nat. Commun.* **9**, 3591 (2018).
89. Silva-Sánchez, A., Soares, M. & Rousk, J. Testing the dependence of microbial growth and carbon use efficiency on nitrogen availability, pH, and organic matter quality. *Soil Biol. Biochem.* **134**, 25–35 (2019).
90. Zhang, X. et al. Erosion effects on soil microbial carbon use efficiency in the mollisol cropland in northeast China. *Soil Ecol. Lett.* **5**, 230176 (2023).
91. Schroeder, J. et al. Liming effects on microbial carbon use efficiency and its potential consequences for soil organic carbon stocks. *Soil Biol. Biochem.* **191**, 109342 (2024).
92. Schmidt, M. W. I. et al. Persistence of soil organic matter as an ecosystem property. *Nature* **478**, 49–56 (2011).
93. Young, I. M. & Crawford, J. W. Interactions and self-organization in the soil-microbe complex. *Science* **304**, 1634–1637 (2004).
94. Kuz'yakov, Y. & Blagodatskaya, E. Microbial hotspots and hot moments in soil: concept & review. *Soil Biol. Biochem.* **83**, 184–199 (2015).
95. Islam, Md. R., Singh, B. & Dijkstra, F. A. Microbial carbon use efficiency of glucose varies with soil clay content: a meta-analysis. *Appl. Soil Ecol.* **181**, 104636 (2023).
96. Cai, Y. et al. Assessing the accumulation efficiency of various microbial carbon components in soils of different minerals. *Geoderma* **407**, 115562 (2022).
97. Jeewani, P. H. et al. The stoichiometric C-Fe ratio regulates glucose mineralization and stabilization via microbial processes. *Geoderma* **383**, 114769 (2021).
98. Bölscher, T., Wadsö, L., Börjesson, G. & Herrmann, A. M. Differences in substrate use efficiency: impacts of microbial community composition, land use management, and substrate complexity. *Biol. Fertil. Soils* **52**, 547–559 (2016).
99. Jones, D. L. et al. Role of substrate supply on microbial carbon use efficiency and its role in interpreting soil microbial community-level physiological profiles (CLPP). *Soil Biol. Biochem.* **123**, 1–6 (2018).
100. Chakrawal, A., Calabrese, S., Herrmann, A. M. & Manzoni, S. Interacting bioenergetic and stoichiometric controls on microbial growth. *Front. Microbiol.* **13**, 859063 (2022).
101. Kleerebezem, R. & Van Loosdrecht, M. C. M. A generalized method for thermodynamic state analysis of environmental systems. *Crit. Rev. Environ. Sci. Technol.* **40**, 1–54 (2010).
102. Cotrufo, M. F., Wallenstein, M. D., Boot, C. M., Deneff, K. & Paul, E. The Microbial Efficiency-Matrix Stabilization (MEMS) framework integrates plant litter decomposition with soil organic matter stabilization: do labile plant inputs form stable soil organic matter? *Glob. Change Biol.* **19**, 988–995 (2013).
103. Craig, M. E. et al. Fast-decaying plant litter enhances soil carbon in temperate forests but not through microbial physiological traits. *Nat. Commun.* **13**, 1229 (2022).
104. Sokol, N. W. et al. The path from root input to mineral-associated soil carbon is shaped by habitat-specific microbial traits and soil moisture. *Soil Biol. Biochem.* 109367 <https://doi.org/10.1016/j.soilbio.2024.109367> (2024).
105. Liang, C., Schimel, J. P. & Jastrow, J. D. The importance of anaerobism in microbial control over soil carbon storage. *Nat. Microbiol.* **2**, 17105 (2017).
106. Li, Z. et al. Microbial metabolic capacity regulates the accrual of mineral-associated organic carbon in subtropical paddy soils. *Soil Biol. Biochem.* 109457 <https://doi.org/10.1016/j.soilbio.2024.109457> (2024).
107. Luo, Y. & Schuur, E. A. G. Model parameterization to represent processes at unresolved scales and changing properties of evolving systems. *Glob. Change Biol.* **26**, 1109–1117 (2020).
108. Kallenbach, C. M., Frey, S. D. & Grandy, A. S. Direct evidence for microbial-derived soil organic matter formation and its ecophysiological controls. *Nat. Commun.* **7**, 13630 (2016).
109. Xiao, K.-Q. et al. Beyond microbial carbon use efficiency. *Natl. Sci. Rev.* nwa059 <https://doi.org/10.1093/nsr/nwae059> (2024).
110. Georgiou, K. et al. Global stocks and capacity of mineral-associated soil organic carbon. *Nat. Commun.* **13**, 3797 (2022).
111. Zhu, E. et al. Enhanced mineral preservation rather than microbial residue production dictates the accrual of mineral-associated organic carbon along a weathering gradient. *Geophys. Res. Lett.* **51**, e2024GL108466 (2024).
112. Garcia-Palacios, P. et al. Dominance of particulate organic carbon in top mineral soils in cold regions. *Nat. Geosci.* <https://doi.org/10.1038/s41561-023-01354-5> (2024).
113. Lí, J. et al. Subarctic winter warming promotes soil microbial resilience to freeze–thaw cycles and enhances the microbial carbon use efficiency. *Glob. Change Biol.* **30**, e17040 (2024).
114. Wu, J., Cheng, X. & Liu, G. Increased soil organic carbon response to fertilization is associated with increasing microbial carbon use efficiency: data synthesis. *Soil Biol. Biochem.* **171**, 108731 (2022).
115. Dukovski, I. et al. A metabolic modeling platform for the computation of microbial ecosystems in time and space (COMETS). *Nat. Protoc.* **16**, 5030–5082 (2021).
116. Karaoz, U. & Brodie, E. L. microTrait: a toolset for a trait-based representation of microbial genomes. *Front. Bioinforma.* **2**, 918853 (2022).
117. Piton, G. et al. Life history strategies of soil bacterial communities across global terrestrial biomes. *Nat. Microbiol.* **8**, 2093–2102 (2023).
118. Gu, C., Kim, G. B., Kim, W. J., Kim, H. U. & Lee, S. Y. Current status and applications of genome-scale metabolic models. *Genome Biol.* **20**, 121 (2019).
119. Abs, E., Albright, M. B. N. & Allison, S. D. Invasions eliminate the legacy effects of substrate history on microbial nitrogen cycling. *Ecosphere* **15**, e4754 (2024).
120. Bernstein, D. B., Sulheim, S., Almaas, E. & Segrè, D. Addressing uncertainty in genome-scale metabolic model reconstruction and analysis. *Genome Biol.* **22**, 64 (2021).
121. Malik, A. A. et al. Defining trait-based microbial strategies with consequences for soil carbon cycling under climate change. *ISME J.* **14**, 1–9 (2020).
122. Demirel, E. et al. Improving the performance of reactive transport simulations using artificial neural networks. *Transp. Porous Media* **149**, 271–297 (2023).
123. Tao, F. et al. Deep learning optimizes data-driven representation of soil organic carbon in earth system model over the conterminous United States. *Front. Big Data* **3**, 17 (2020).
124. Reichstein, M. et al. Deep learning and process understanding for data-driven Earth system science. *Nature* **566**, 195–204 (2019).
125. Song, J. et al. A meta-analysis of 1,119 manipulative experiments on terrestrial carbon-cycling responses to global change. *Nat. Ecol. Evol.* **3**, 1309–1320 (2019).
126. Norby, R. J. et al. Model–data synthesis for the next generation of forest free-air CO₂FACE experiments. *New Phytol.* **209**, 17–28 (2016).
127. Tifafi, M. et al. The use of radiocarbon ¹⁴C to constrain carbon dynamics in the soil module of the land surface model ORCHIDEE (SVN r5165). *Geosci. Model Dev.* **11**, 4711–4726 (2018).
128. Goll, D. S. et al. A representation of the phosphorus cycle for ORCHIDEE (revision 4520). *Geosci. Model Dev.* **10**, 3745–3770 (2017).
129. Luo, Y. et al. Toward more realistic projections of soil carbon dynamics by Earth system models. *Glob. Biogeochem. Cycles* **30**, 40–56 (2016).
130. Tao, F. et al. Convergence in simulating global soil organic carbon by structurally different models after data assimilation. *Glob. Change Biol.* **30**, e17297 (2024).

131. Wieder, W. R., Grandy, A. S., Kallenbach, C. M., Taylor, P. G. & Bonan, G. B. Representing life in the Earth system with soil microbial functional traits in the MIMICS model. *Geosci. Model Dev.* **8**, 1789–1808 (2015).
132. Zhang, H. et al. Microbial dynamics and soil physicochemical properties explain large-scale variations in soil organic carbon. *Glob. Change Biol.* **26**, 2668–2685 (2020).

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X.H. and D.S.G. initiated the writing and led the design and writing of the article. S.A., S.M., P.C., G.H., K.G., Y.L., N.N., and Y.W. participated in the initial design of the content. E.A. and S.A. drafted section 4.1. F.T. and Y.H. drafted section 4.3. R.A., E.B., S.P.K.B., A.C., L.E., P.F., L.B.H., W.L., G.M., C.Q., and S.S. provided input on the manuscript text, figures and discussion of scientific content.

Competing interests

The authors declare no competing interests.

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