

New insights into the mechanisms of plant isotope fractionation from combined analysis of intramolecular ¹³C and deuterium abundances in *Pinus nigra* tree-ring glucose

Thomas Wieloch^{1,2} (D), Meisha Holloway-Phillips³ (D), Jun Yu⁴ (D) and Totte Niittylä¹ (D)

¹Department of Forest Genetics and Plant Physiology, Swedish University of Agricultural Sciences, Umeå Plant Science Centre, 90183, Umeå, Sweden; ²Division of Geological and Planetary Sciences, California Institute of Technology, 91125 Pasadena, CA, USA; ³Research Unit of Forest Dynamics, Swiss Federal Institute for Forest, Snow and Landscape Research WSL, 8903, Birmendsorf, Switzerland; ⁴Department of Mathematics and Mathematical Statistics, Umeå University, 90187, Umeå, Sweden

Author for correspondence: Thomas Wieloch Email: thomas.wieloch@slu.se

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Summary

• Understanding isotope fractionation mechanisms is fundamental for analyses of plant ecophysiology and paleoclimate based on tree-ring isotope data.

• To gain new insights into isotope fractionation, we analysed intramolecular ¹³C discrimination in tree-ring glucose (Δ'_i , i = C-1 to C-6) and metabolic deuterium fractionation at H¹ and H² (ε_{met}) combinedly. This dual-isotope approach was used for isotope-signal deconvolution.

• We found evidence for metabolic processes affecting Δ_1' and Δ_3' , which respond to air vapour pressure deficit (VPD), and processes affecting Δ_1' , Δ_2' , and ε_{met} , which respond to precipitation but not VPD. These relationships exhibit change points dividing a period of homeostasis (1961–1980) from a period of metabolic adjustment (1983–1995). Homeostasis may result from sufficient groundwater availability. Additionally, we found Δ_5' and Δ_6' relationships with radiation and temperature, which are temporally stable and consistent with previously proposed isotope fractionation mechanisms.

• Based on the multitude of climate covariables, intramolecular carbon isotope analysis has a remarkable potential for climate reconstruction. While isotope fractionation beyond leaves is currently considered to be constant, we propose significant parts of the carbon and hydrogen isotope variation in tree-ring glucose originate in stems (precipitation-dependent signals). As basis for follow-up studies, we propose mechanisms introducing Δ_1' , Δ_2' , Δ_3' , and ε_{met} variability.

Introduction

Analysis of the systematic ¹³C/¹²C variation (commonly termed ¹³C signal'; abbreviations in Table 1) across tree-ring series is widely used to study past climate conditions, plant-environment interactions, and physiological traits such as leaf water-use efficiency (CO₂ uptake relative to H₂O loss) (Leavitt & Roden, 2022). Signals found at the whole-tissue or whole-molecule level (Fig. 1a, top and middle) are commonly interpreted based on a simplified mechanistic model of ¹³C discrimination, Δ (denoting ¹³C/¹²C variation caused by physiological processes) (Farguhar et al., 1982). This model considers isotope effects of CO2 diffusion from ambient air into intercellular air spaces (Craig, 1953) and CO₂ assimilation by rubisco (Roeske & O'Leary, 1984) and phosphoenolpyruvate carboxylase (PEPC; Fig. 2) (Farquhar, 1983; Farquhar & Richards, 1984). Manifestation of these effects as ¹³C discrimination depends on the ratio of intercellular-to-ambient CO₂ partial pressure $(p_i : p_a)$ (Farquhar et al., 1982), and a highly significant positive relationship between $p_i: p_a$ and leaf Δ was confirmed experimentally (Evans *et al.*, 1986). Environmental parameters influence $p_i : p_a$ and thus leaf Δ (Evans *et al.*, 1986) by affecting the stomatal aperture and CO₂ assimilation. For instance, in response to drought, isohydric plant species such as *Pinus nigra* (studied here) close their stomata (McDowell *et al.*, 2008). This can be expected to decrease $p_i : p_a$ and leaf Δ (Farquhar *et al.*, 1982; Evans *et al.*, 1986).

Isotope fractionation by metabolic processes downstream of CO_2 assimilation is complex (Hobbie & Werner, 2004), incompletely understood (Badeck *et al.*, 2005; Cernusak *et al.*, 2009), and has yet to be adequately integrated into ¹³C-discrimination models (Ubierna *et al.*, 2022). Specifically, the simplified ¹³C discrimination model described above requires multiple adaptations to enable correct interpretation of the ¹³C composition of tree-ring glucose (studied here). For instance, we recently argued that incorporation of carbon assimilated by PEPC into tree-ring glucose is negligible because leaves lack a high-flux pathway shutting this carbon into glucose metabolism (Fig. 2; Wieloch *et al.*, 2022c). Therefore, all carbon in tree-ring glucose the

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Table 1 Abbreviations and symbols.

Abbreviation	Definition
¹³ C signal	Systematic ¹³ C/ ¹² C variation
DAHPS	3-Deoxy-D-arabino-heptulosonate-7-phosphate synthase
F6P	Fructose 6-phosphate
G6P	Glucose 6-phosphate
G6PD	Glucose-6-phosphate dehydrogenase
GAP	Glyceraldehyde 3-phosphate
GAPDH	Glyceraldehyde-3-phosphate dehydrogenase
NMRS	Nuclear magnetic resonance spectroscopy
OPPP	Oxidative branch of the pentose phosphate pathway
PEP	Phosphoeno/pyruvate
PEPC	Phosphoeno/pyruvate carboxylase
PGA	3-Phosphoglycerate
PGI	Phosphoglucose isomerase
РК	Pyruvate kinase
RuBP	Ribulose 1,5-bisphosphate
TCAC	Tricarboxylic acid cycle
Symbol	Definition
D _a	Ambient CO ₂ partial pressure
Di	Intercellular CO ₂ partial pressure
PRE	Precipitation
RAD	Global radiation
SD	Sunshine duration
SPEI _i	Standardised precipitation-evapotranspiration index calculated for different timescales, <i>i</i> = 1, 3, 6, 8, 12, 16, 24, 36, 48 months
ТМР	Air temperature
VPD	Air vapour pressure deficit
Δ	¹³ C discrimination denoting ¹³ C/ ¹² C variation due to plant physiological processes
Δ_i'	Intramolecular ¹³ C discrimination where <i>i</i> denotes individual glucose carbon positions and the prime denotes data corrected for ¹³ C signal redistribution by heterotrophic triose phosphate cycling
Δ_{1-2}'	Arithmetic average of Δ_1' and Δ_2'
Δ_{1-3}'	Arithmetic average of Δ_1' , Δ_2' , and Δ_3'
Δ_{5-6}'	Arithmetic average of Δ_5' and Δ_6'
€ _{met}	Metabolic deuterium fractionation at glucose H ¹ and H ²

Moreover, we recently measured Δ intramolecularly at all six carbon positions, *i*, of glucose (Fig. 1a, bottom) extracted across an annually resolved tree-ring series of *P. nigra* (Wieloch *et al.*, 2018). The resultant Δ_i ' data set comprises 6 × 31 values (study period: 1961–1995; four years missing: 1977, 1978, 1981, 1982), which were corrected for ¹³C signal redistribution by heterotrophic triose phosphate cycling (indicated by prime, Supporting Information Notes S1). We found that, at least, four ¹³C signals contribute to the interannual ¹³C/¹²C variability in tree-ring glucose (Fig. 1b) and proposed the following theories on underlying mechanisms.

We initially proposed the diffusion-rubisco signal is preserved at C-1 to C-3 (Figs 1b, 2; Wieloch et al., 2018); although this view is modified here. Additionally, C-1 and C-2 are thought to carry ¹³C signals due to fractionation at phosphoglucose isomerase (PGI has carbon isotope effects at both C-1 and C-2) and glucose-6-phosphate dehydrogenase (G6PD has a carbon isotope effect at C-1) (Wieloch et al., 2018, 2022a). Two leaf-level mechanisms of signal introduction were proposed. First, with decreasing carbon assimilation, the PGI reaction in chloroplasts moves from being on the side of fructose 6-phosphate (F6P) towards equilibrium (Fig. 2; Dietz, 1985). This shift is expected to cause ¹³C enrichments at C-1 and C-2 of glucose 6-phosphate (G6P) and its derivatives starch and tree-ring glucose (Table 2; Wieloch et al., 2018). Moreover, shifts towards PGI equilibrium are associated with G6P increases (Dietz, 1985). Increasing G6P is thought to cause G6PD activation and thus increasing flux through the oxidative pentose phosphate pathway (OPPP) in chloroplasts (Cossar et al., 1984; Sharkey & Weise, 2016; Preiser et al., 2019) resulting in additional ¹³C enrichment at C-1 of

Fig. 1 Carbon isotope discrimination in tree rings. (a) Levels of resolution of stable carbon isotope analysis: whole plant materials, whole molecules, intramolecular carbon positions. (b) Hierarchical clustering of Δ_i' series for the period 1961–1995. Significance of series correlation: *, $P \le 0.05$; **, $P \le 0.01$; ***, $P \le 0.001$. Modified figure from Wieloch *et al.* (2018). Δ_i' denotes intramolecular ¹³C discrimination in glucose extracted across an annually resolved *Pinus nigra* tree-ring series.



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Fig. 2 Proposed metabolic origins of carbon and hydrogen isotope signals in tree-ring glucose. Dashed arrows indicate that intermediate reactions are not shown. 2PG. 2-phosphoglycolate; 2PGA, 2phosphoglycerate; 6PGL, 6phosphogluconolactone; ADP, adenosine diphosphate; ATP, adenosine triphosphate; CIT, citrate; COR cycle, cytosolic oxidation-reduction cycle; DAHP, 3-deoxy-D-arabino-heptulosonate 7-phosphate; DAHPS, DAHP synthase; E4P, erythrose 4phosphate; F6P, fructose 6-phosphate; FAD, flavin adenine dinucleotide; Fru, fructose; G6P, glucose 6-phosphate; G6PD, G6P dehydrogenase; GDC, glycine decarboxylase complex; Glc, glucose; MEP pathway, methylerythritol 4-phosphate pathway; NAD⁺, nicotinamide adenine dinucleotide; NADP⁺, nicotinamide adenine dinucleotide phosphate; np-GAPDH, nonphosphorylating glyceraldehyde-3phosphate dehydrogenase; OAA, oxaloacetate; OPPP, oxidative pentose phosphate pathway; PEP, phosphoeno/pyruvate; PEPC, PEP carboxylase; PGA, 3-phosphoglycerate; p-GAPDH, phosphorylating glyceraldehyde-3phosphate dehydrogenase; PGI, phosphoglucose isomerase; PGK, phosphoglycerate kinase; Pi, inorganic phosphate; PK, pyruvate kinase; PP, pentose phosphate; PRE, precipitation; Pyr, pyruvate; RAD, global radiation; Rubisco, ribulose-1,5bisphosphate carboxylase/oxygenase; RuBP, ribulose 1,5-bisphosphate; Suc, sucrose; SuSy, sucrose synthase; TCAC, tricarboxylic acid cycle; TK, transketolase; TMP, air temperature; TP, triose phosphates (glyceraldehyde 3-phosphate, dihydroxyacetone phosphate); UDPG, uridine diphosphate glucose; VPD, air vapour pressure deficit; Δ_i' , intramolecular ¹³C discrimination where *i* denotes individual glucose carbon positions and the prime denotes data corrected for ¹³C signal redistribution by heterotrophic triose phosphate cycling; ε_{met} , metabolic deuterium fractionation at glucose H¹ and H^2 .

G6P and its derivatives (Wieloch *et al.*, 2022a). Hydrogen isotope evidence consistent with these proposed metabolic shifts was reported recently (Wieloch *et al.*, 2022a). Second, the PGI reaction in chloroplasts is usually displaced from equilibrium on the side of F6P, whereas the PGI reaction in the cytosol is closer to or in equilibrium (Dietz, 1985; Gerhardt *et al.*, 1987; Leidreiter *et al.*, 1995; Schleucher *et al.*, 1999; Szecowka *et al.*, 2013). This is expected to result in ¹³C/¹²C differences between starch and

G6P. C-2

Table 2 Hydro Gilbert <i>et al</i> ., 2	ogen isotope effec 2012).	ts of phosphoglucose	isomerase (PGI, Rose	& O'Connell, 1961), an	d carbon isotope	e effects of glucose ison	nerase (GI,
PGI, $\alpha = k_{\rm H}/k_{\rm H}$	D		GI, $\alpha = k_{12}/k$	¹³ C			
F6P, H ^{1R}		G6P, H ²	F6P, C-1	F6P, C-2		G6P, C-1	G6P, C-2
	\rightarrow	2.2			\rightarrow	1.005	1.015
$\frac{0.9}{2}$	\leftrightarrow	1.1	<u>1.013</u> 1.018	<u>0.993</u> 1.008	\leftrightarrow	0.987	1.007
sucrose at	both hexose	C-1 and C-2 (T	able 2: Wieloch	This includes a	enolase, pyrus	vate kinase (PK). F	PEPC and 3
sucrose at	both herose	C-1 and C-2 (T	Table 2. Wieloch	This includes	enolose pyrus	rate kinase (PK) F	PEPC and 3
et al., 2022). By extension	n, changes in the re	lative contribution	deoxy-D-arabina	-heptulosonat	e-7-phosphate synth	ase (DAHPS)
of starch to	the biosynthes	is of tree-ring gluc	ose is expected to	the first enzyme	of the shikin	nate pathway. Breaki	ing the double
contribute to	o the ¹³ C signal	s at C-1 and C-2.		bond in PEP is	thought to p	roceed faster when	¹² C instead o
In additio	on to ¹⁵ C signa	als at C-1 and C-2	, tree-ring glucose	¹⁵ C forms this	bond (Wield	och <i>et al.</i> , 2022c).	Consequently
samples disc	cussed here carr	y deuterium signal	s caused by meta-	increasing relati	ve flux into n	netabolism downstre	eam of PEP i
bolic proces	ses at H ¹ and	H ² . These signals :	are strongly corre-	thought to ¹³ C	enrich remain	ning PEP at the dou	uble-bond car
lated and we	ere approximate	d as:		bons and their a	derivatives incl	luding glucose C-5 :	and C-6 (Wie

Table 2 Hydrogen isotope effects of phosphoglucose iso ose isomerase (GI, Gilbert et al., 2012).

<i>c</i> –	$(D_1 + D_2)/2$	1 Fan
$\epsilon_{\rm met}$ –	$(D_3 + D_4 + D_5 + D_{6S} + D_{6R})/5$	-i Equ

lated and were approximated as:

where D_i denotes relative deuterium abundances at individual H-C positions (Wieloch et al., 2022a,b). Variability of ε_{met} pertaining to glucose H¹ and H² was attributed to isotope effects of G6PD $(k_{\rm H}/k_{\rm D} = 2.97)$ (Hermes *et al.*, 1982) and PGI (Table 2; Rose & O'Connell, 1961; Wieloch et al., 2022a,b), respectively. Proposedly, G6PD and PGI-dependent metabolic processes in both leaves and tree rings may contribute to ε_{met} signal introduction (Wieloch et al., 2022b). Interestingly, Wacker (2022) recently reported that the commonly observed whole-molecule deuterium depletion of leaf starch that derives from deuterium depletion at starch glucose H² (Schleucher et al., 1999; Wieloch et al., 2022a) is not detectable in nocturnal sucrose. Proposedly, this depletion is either washed out at the level of cytosolic PGI or masked either by the vacuolar sucrose pool or deuterium enrichments at other sucrose hydrogen positions. Washout would imply that any ε_{met} signal present at leaf-level G6P H² is lost to the medium. In this case, the ε_{met} signal at tree-ring glucose H² may originate outside leaves.

At tree-ring glucose C-4 (Fig. 1b), the diffusion-rubisco 13 C signal is thought to be absent due to counteracting fractionation by leaf-cytosolic glyceraldehyde-3-phosphate dehydrogenases (GAPDH; Fig. 2) (Wieloch et al., 2021). Signal removal may involve both changes in 3-phosphoglycerate (PGA) flux into downstream metabolism including the tricarboxylic acid cycle (TCAC) relative to flux into tree-ring glucose and changes in flux through the cytosolic oxidation-reduction cycle (Wieloch, 2021; Wieloch et al., 2021).

The ¹³C signal at C-5 and C-6 (Fig. 1b) is thought to derive from the postulated (but not yet measured) isotope effects of leaf-level enzymes that modify the carbon double bond in phosphoenolpyruvate (PEP, Fig. 2) (Wieloch et al., 2022c).

PK), PEPC, and 3e synthase (DAHPS), Breaking the double when ¹²C instead of 22c). Consequently, ownstream of PEP is the double-bond carbons and their derivatives including glucose C-5 and C-6 (Wieloch et al., 2022c). For example, O₃ causes downregulation of rubisco, upregulation of PEPC, and DAHPS expression (Dizengremel, 2001; Janzik et al., 2005; Betz et al., 2009). This is expected to cause increasing relative flux into metabolism downstream of PEP (Wieloch et al., 2022c). Accordingly, we previously found a negative relationship between reconstructed tropospheric O₃ concentration and tree-ring glucose Δ_{5-6} ' (arithmetic average of Δ_5 ' and Δ_6 ', Table 1) (Wieloch *et al.*, 2022c).

By contrast, the diffusion-rubisco signal is not evident at C-5 and C-6 (Wieloch et al., 2022c). This was explained (inter alia) by interaction between photorespiration and the TCAC (Fig. 2; Wieloch et al., 2022c). Photorespiration increases with drought, which results in increasing supply of mitochondrial NADH via the glycine decarboxylase complex. Since this NADH can feed oxidative phosphorylation, NADH and FADH₂ supply by the TCAC, which requires injection of PEP into the TCAC via PK and PEPC, may be reduced. This should result in Δ_{5-6} ' increases counteracting drought-induced decreases in diffusion-rubisco discrimination (as mentioned in the first section).

The theories of isotope signal introduction outlined above require further testing. They derive from separate analyses of either the Δ_i ' or deuterium data set. However, some reactions exhibit both carbon and hydrogen isotope effects (e.g. G6PD at G6P C-1 and H¹; PGI at G6P C-1, C-2, and H² but not H¹) and should therefore introduce intercorrelated ¹³C and deuterium signals (suggested terminology: hydro-carbon isotope signals and hydro-carbon isotope fractionation). Combined analysis of intramolecular ¹³C and deuterium data can, in principle, help to separate those signals from signals introduced by reactions, which merely exhibit either carbon or hydrogen isotope effects. Therefore, we here studied the relationships between Δ_i' and ε_{met} and their dependence on environmental parameters. Based on our results, we critically examine and revise existing isotope theory and provide new insights into a central open question - whether

carbon and hydrogen isotope variability across tree rings derives from leaf-level processes only (as supported by current evidence) or whether processes in the stem contribute as well.

Materials and Methods

Isotope data

The Δ_i ' and ε_{met} data sets of *P. nigra* Arnold reanalysed here are described in Wieloch *et al.* (2018, 2022b) and in Notes S1. Δ may be affected by p_a (Schubert & Jahren, 2012). Annual p_a data were obtained for the Mauna Loa Observatory, HI (curators: Pieter Tans, NOAA/ESRL, Boulder, USA; Ralph Keeling, Scripps Institution of Oceanography, La Jolla, USA). However, we found no significant correlation between Δ and p_a (r=-0.32, P>0.05, n=31). Moreover, in line with the mechanistic model of diffusion–rubisco ¹³C discrimination (Farquhar *et al.*, 1989), most previous studies have reported positive slopes for the regression of these variables (Schubert & Jahren, 2012), but we found a negative slope (-0.02 ± 0.01 SE). Hence, in the present case, the effect of p_a on Δ is not verifiable. Therefore, we neither correct Δ nor Δ_i ' for p_a .

Estimation of ε_{met} according to Eqn 1 aims to remove fractionation by nonmetabolic processes such as leaf water deuterium enrichment (Wieloch *et al.*, 2022b). Series of ε_{met} calculated separately for H¹ and H² (by replacing the numerator in Eqn 1 with D_1 and D_2 , respectively) are essentially perfectly correlated considering that both series exhibit significant random variation due to the relatively large error of NMRS measurements (1983–1995, r = 0.92, $P = 10^{-5}$, n = 13).

Climate data

Data of relative humidity, precipitation (PRE), global radiation (RAD), sunshine duration (SD), and air temperature (TMP) are from the climate station Hohe Warte (Vienna, Austria, 48.23°N, 16.35°E, 198 m amsl) (Klein Tank et al., 2002). Air vapour pressure deficit (VPD) was calculated following published procedures (Abtew & Melesse, 2013). Data of the standardised precipitation-evapotranspiration index (SPEI_i) calculated for integrated periods of i = 1, 3, 6, 8, 12, 16, 24, 36, 48 months were obtained for 48.25°N, 16.25°E (Fan & van den Dool, 2004; Beguería et al., 2010). The SPEI is a multiscalar drought index approximating soil moisture variability when calculated for short timescales and groundwater variability when calculated for long timescales (Vicente-Serrano et al., 2010). The RAD series starts in 1964 while all other climate series start in 1961. Horizontal distances between the tree site and the climate station and grid point are < 15 km. Vertical offsets are small. Hence, climate data and site conditions are expected to be in good agreement.

Data analysis

Based on *TMP* during the study period (1961–1995), the growing season at the site was estimated to extend from March to November (Wieloch *et al.*, 2018). Conifers form tree rings over

the course of several months (Cuny *et al.*, 2014). Therefore, all statistical analyses exclusively consider periods comprising ≥ 4 growing season months. According to autocorrelation analyses on Δ_i ' and Δ series, the growth of the trees studied here has not been significantly affect by interannual carry-over of carbon (Wieloch *et al.*, 2018). Therefore, our statistical analyses do not consider the climate conditions of previous years.

After mean-centring and unit-variance scaling of Δ_i ' series, hierarchical cluster analysis was done with the functions dist() and hclust() of the STATS package in R (R Core Team, 2021), choosing Euclidean distances and Ward's fusion criterion as inputs (Ward, 1963). Pearson's correlation analysis, ordinary least squares regression analysis, and Shapiro–Wilk normality tests were respectively done with the functions cor(), lm(), and shapiro.test() of the STATS package in R (R Core Team, 2021). The fraction of systematic variance in isotope series was estimated according to published procedures (Nilsson *et al.*, 1996). Change point tests were done with the function detectChangePointBatch() of the CPM package in R (parametric generalised likelihood ratio test and nonparametric Lepage test) (Ross, 2015). *F*-tests and one-tailed *t*-tests were, respectively, done with the functions f.test() and t.test() in EXCEL (Microsoft Corp., Redmond, WA, USA).

Results

Hydro-carbon isotope signals at tree-ring glucose HC-1 and HC-2

Tree-ring glucose of our P. nigra samples exhibits strongly correlated hydrogen isotope signals at H^1 and H^2 (Wieloch et al., 2022b). These signals occur only after crossing a change point in 1980. Isotope-environment relationship analyses indicated that the trees had likely access to groundwater before 1980, which prevented changes in the processes introducing these isotope signals. We proposed the signals derive from the hydrogen isotope effects of G6PD ($k_{\rm H}/k_{\rm D} = 2.97$) (Hermes *et al.*, 1982) and PGI (Table 2; Rose & O'Connell, 1961; Wieloch et al., 2022a) in autotrophic and/or heterotrophic tissue (Fig. 2; see the Introduction section) (Wieloch et al., 2022a,b). If this proposal is correct then there should be related signals in Δ_1 ' and Δ_2 ' due to the carbon isotope effects of G6PD affecting C-1 $(k_{12}C/k_{13}C = 1.0165)$ (Hermes et al., 1982) and PGI affecting C-1 and C-2 (Table 2; Gilbert et al., 2012). Several findings support this hypothesis. First, among all $\Delta_i{'}$ series, $\Delta_1{'}$, $\Delta_{1-2}{'}$, $\Delta_{1-3}{'}$, and Δ are not normally distributed (Table S1, negative skew). Second, among these non-normal series, Δ_{1-2}' , Δ_{1-3}' , and Δ exhibit a change point in 1980 (Δ_{1-2}' : parametric test, P < 0.001, nonparametric test, P < 0.05; Δ_{1-3} ': parametric test, P < 0.01; Δ : parametric test: P < 0.05; n = 31). Third, 1983–1995 average values of Δ_1 ', Δ_2 ', Δ_{1-2} ', and Δ are significantly lower than the average values of 1961-1980, while the 1983–1995 variance is significantly larger (Table S2). By contrast, Δ_3 ' does not exhibit significant differences in average value or variance between the two periods. Fourth, Δ_1 ' and Δ_2 ' data pertaining to 1983–1995 are significantly correlated (r = 0.67, P = 0.01, n = 13). Fifth, ε_{met} approximates average hydrogen isotope fractionation at glucose H¹ and H² caused by metabolic processes

(Eqn 1). Using simple linear regression modelling, we found significant negative relationships between the 1983-1995 data of $\varepsilon_{\rm met}$ and Δ_1' as well as Δ_2' , but not Δ_3' or any other Δ_i' (Fig. 3, green circles; $\Delta_1 ' \sim \epsilon_{met}$: $R^2 = 0.35$, $adjR^2 = 0.29$, P = 0.03; $\Delta_2 ' \sim \epsilon_{met}$: $R^2 = 0.54$, $adjR^2 = 0.50$, P = 0.004; $\Delta_3' \sim \varepsilon_{\text{met}}$: $R^2 = 0.21$, $\operatorname{adj} R^2 = 0.13$, P > 0.1; n = 13; Table S3). Our ¹³C-NMRS data exhibit relatively large measurement errors. Based on estimates of this random error variance, c. 88% of the variance in the Δ_1 ' and Δ_2 ' data of 1983–1995 is systematic variance (Table S4). Hence, c. 33% and 57% of the systematic variance in Δ_1 ' and Δ_2 ' is explained by processes causing $\varepsilon_{\rm met}$ variation (0.29/0.88 and 0.5/0.88) while c. 67% and 43%, respectively, go back to other processes. Taken together, carbon and hydrogen isotope signals at glucose HC-1 and HC-2 are significantly associated during 1983-1995 but not during 1961-1980 (Notes S2). The processes introducing these signals cause concerted ¹³C and deuterium enrichments (Fig. 3a,b).

Isotope–environment relationships at tree-ring glucose C-1 to C-3

As evident from our previously published hierarchical cluster analysis and Pearson's correlation analyses for the whole period (1961–1995), Δ_1 ', Δ_2 ', and Δ_3 ' share common variability (Fig. 1b; Wieloch *et al.*, 2018). Since Δ_{1-2} ' and Δ_{1-3} ' exhibit change points in 1980 (as mentioned in the previous section and Tables S1, S2), we analysed the early (1961–1980) and late period (1983–1995) separately.

During the late period, Δ_1 ' and Δ_3 ' are more closely associated (Fig. 4a; r = 0.87, $P = 10^{-4}$, n = 13) than Δ_1 ' and Δ_2 ' (r=0.67, P=0.01, n=13). While this contrasts with results for the whole period (Fig. 1b), it is consistent with isotope-climate relationship patterns for the late period. Δ_1 ' and Δ_3 ' correlate similarly with numerous climate parameters and periods (Table 3; VPD, PRE, SPEI₁ to SPEI₁₆, TMP, SD). By contrast, Δ_2 ' correlates only with one VPD period and several PRE periods. A model including ε_{met} and growing season VPD as cofactors captures most of the systematic variance in Δ_1 ' of 88% (Tables 4 (M1), S4). Consistent with the findings described earlier (Fig. 3; Table 3), only ε_{met} but not growing season VPD contributes significantly to the Δ_2 ' model, whereas only growing season VPD but not ε_{met} contributes significantly to the Δ_3 ' model (Table 4 (M2, M3)). Removing insignificant terms, we find that ε_{met} explains 57% of the systematic variance in Δ_2 ', while growing season VPD explains the entire systematic variance in Δ_3 ' (Tables 4 (M4, M5), S4). The effect of VPD on Δ_1 ' is about twice as large as on Δ_3 ' (Table 4 (M1 vs M5)) while the effect of $\varepsilon_{\rm met}$ on Δ_1 ' is about half as large as on Δ_2 ' (M1 vs M4). Intriguingly, ${\it \Delta_1}\,'$ and ${\it \Delta_3}\,'$ are affected by processes that respond to growing season VPD. VPD-dependent processes can account for both the clustering and correlation between Δ_1 ' and Δ_3 ' data of 1983–1995 (Fig. 4a). By contrast, ε_{met} is significantly correlated only with PRE (especially March-July PRE) but no other climate parameter (Tables 4 (M11), S5). Furthermore, in our Δ_1 and Δ_2 ' models, ε_{met} can be substituted by March–July *PRE* (Table 4 (M1 vs M6, M4 vs M7)).

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Fig. 3 Relationship between the average hydrogen isotope fractionation caused by metabolic processes at glucose H¹ and H² (ε_{met}) and ¹³C discrimination at C-1, C-2, and C-3 (Δ_1' , Δ_2' , and Δ_3'). Glucose was extracted across an annually resolved tree-ring series of *Pinus nigra* from the Vienna Basin (black squares, 1961–1980; green circles, 1983–1995). Dashed line, relationship between the hydrogen and carbon isotope data of the period 1983–1995.



Fig. 4 Hierarchical clustering of Δ_i' series for the periods 1983–1995 (a) and 1961–1980 (b). Δ_i' denotes intramolecular ¹³C discrimination in treering glucose of *Pinus nigra* from the Vienna basin with *i* denoting individual glucose carbon positions. Significance of series correlation: **, $P \le 0.01$; ****, $P \le 10^{-4}$.

During the early period, Δ_1 ', Δ_2 ', and Δ_3 ' are not significantly correlated (Fig. 4b). Furthermore, isotope–environment models that work for the late period (Table 4 (M5–M7)) do not work for the early period (M8–M10). Compared with the late period, we found fewer and weaker isotope–climate correlations (Table 5).

Isotope-environment relationships at tree-ring glucose C-4 to C-6 $\,$

As evident from our previously published hierarchical cluster analysis and Pearson's correlation analyses for the whole period, Δ_4 ', Δ_5 ', and Δ_6 ' share common variability, and Δ_5 ' and Δ_6 ' are significantly correlated (Fig. 1b; r = 0.61, P < 0.001, n = 31) (Wieloch et al., 2018). This significant correlation holds for both the early and late period (Fig. 4). Furthermore, we did not find change points in the Δ_4 ', Δ_5 ', and Δ_6 ' series (Tables S1, S2). Therefore, we analysed isotope-environment relationships for the whole period. We found that Δ_5 ' and Δ_6 ' correlate with numerous climate parameters and periods but most significantly with RAD while significant Δ_4 '-climate correlations are rare (Table 6). Models including April-September RAD and March-October TMP as cofactors capture 96% of the systematic variance in Δ_{5-6} , Δ_{5} , and Δ_6 ' of 73%, 66%, and 45%, respectively (Table 7 (M1–M3); Δ_{5-6} ', $\operatorname{adj} R^2 = 0.70$, $P = 10^{-7}$; Δ_5 ', $\operatorname{adj} R^2 = 0.64$, $P = 10^{-6}$; Δ_6' , adj $R^2 = 0.43$, P < 0.001; n = 28; Table S4). Based on RAD regression slopes (which are better constrained than TMP regression slopes), the ¹³C discrimination at C-5 is c. 1.5 times larger than at C-6 (Table 7 (M2-M3)). The model works well for both the early and late period (Table 7 (M4-M5)). Furthermore, consistent with the weak association between Δ_4 ' and Δ_{5-6} ' (Fig. 1b), the model works reasonably well for Δ_4 ', considering the relatively low systematic variance in Δ_4 ' of 38% (Tables 7 (M6), S4).

Discussion

Intramolecular carbon isotope analysis of tree-ring glucose yields information about metabolic variability and water status of both leaves and stems

We found evidence for processes affecting Δ_1 ' and Δ_3 ', which respond to *VPD* (Table 4 (M1, M3, M5)). Intriguingly, we also

found evidence for processes simultaneously affecting $\varepsilon_{\rm met}$, Δ_1 ', and Δ_2 ', which respond to *PRE* but not *VPD* (Tables 4 (M1, M2, M4, M6, M7, M11), S5). This sensitivity to different hydrological properties may be explained by the fact that stem capacitance can buffer stem water status against changes in VPD (McCulloh et al., 2019), whereas leaf water status is tightly coupled to VPD (Grossiord et al., 2020). Changes in PRE will affect soil water potential and hence both stem and leaf water status. Variability in leaf water status may be impacted more by VPD than by soil water status, which would explain why VPD is the best predictor of the intercorrelated processes affecting Δ_1 ' and Δ_3 '. By contrast, VPDinsensitive processes affecting $\varepsilon_{\rm met}$, Δ_1 ', and Δ_2 ' may reside in stems. Hence, we propose intramolecular carbon and hydrogen isotope analysis of tree-ring glucose yields information about metabolic variability and water status not only of leaves but also of stems. PRE-dependent systemic changes in enzyme expression can be considered as an alternative explanation.

Isotope fractionation mechanisms in leaves affecting treering glucose C-1 to C-3 $\,$

The Δ_{1-2}' and Δ_{1-3}' series exhibit change points in 1980; that is, their frequency distributions do not align with the properties of a single theoretical probability distribution (see the Results section; Tables S1, S2). Consequently, we investigated the early (1961–1980) and late period (1983–1995) separately. During the late period, Δ_1' and Δ_3' are significantly intercorrelated (Fig. 4a) and correlate negatively with *VPD* and positively with short-term *SPEI*, whereas Δ_2' lacks most of these correlations (Table 3). Furthermore, during the late period, growing season *VPD* accounts for a significant fraction of the systematic variance in Δ_1' and the entire systematic variance in Δ_3' but does not contribute significantly to explaining Δ_2' (Tables 4 (M1, M2, M5), S4). Hence, increasing *VPD* during 1983–1995 causes ¹³C enrichments at treering glucose C-1 and C-3 but not C-2. At C-1, the effect is about twice as large as at C-3 (Table 4 (M1 and M5)).

As discussed above, the *VPD*-dependent processes affecting Δ_1 ' and Δ_3 ' are likely located in leaves. Qualitatively, *VPD*-induced ¹³C enrichments at C-1 and C-3 are consistent with the mechanisms of diffusion–rubisco fractionation (see the Introduction section). However, diffusion–rubisco fractionation affects all glucose carbon positions equally (Wieloch *et al.*, 2018). Hence,

	V	٩	I	PR 	щ		SPEI	_	νI	PEI ₃		SPi	El_4	· ·	SPEI ₆		SP	PEI ₈	l	SPEI ₁₂	ا یہ ا	PEI ₁₆	SPEI ₂₄	- SPE	-1 ₃₆	SPEI ₄₈	≧	4P		ß		RAD	
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ALLMAM	U		9	q	а	ъ	U	q	U		р	q		ъ	Р		а										ъ						
MAMJJAS	U		q	q	q		р	ъ	U		q	U		в	Р	а	р														ъ		
MAMJJASO	U		q	ъ	а		р	ъ	U	_	ъ	U		в	р	а	р		ъ	а											ъ		
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ALLM	q		q	ъ			р	ъ	q		q	q		q	р	q	q		ъ	а	а						ъ		ъ		ъ		
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JJAS	q		q				а		q		q	q		q	U	q	q		а	q	а												
OSALL	р		q						q		ъ	q		q	U.	q	U		ъ	q	ы												
NOSALL	9		ъ						q		ъ	q		ы	q	q	U		q	q	a	ъ											
JASO	ъ		ъ						q		ъ	q		ъ	U.	q	U		q	q	aa	в											
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Significance lev sunshine durati Climate data w discrimination a	vels: « ion; <u>5</u> ere a at glu	a, ≤' 5 <i>PEI;</i> iver∂	0.05 , sta 1, sta 1, sta	5 (lig unda 1 for 1, C	rdise all≥ -2, a	rey); ed pr 4-n	t b, ≤ ecipit nonth C-3, re	0.01 ation peric	(me -eva ods (dium apotra of the ly. Glu	grey anspi : gro	y); c, iratic wing e wa	≤0. Son inc seas	001 (Jex c son (i racte	(dark)f diff Marcl ≥d acr	grey eren h-Nc 'oss a	(). Un t peri ovem an an	nderlin iods (iber). inuall	ne d∉ ï = 1 Mor y res	enotes n , 3, 6, 8 iths wer olved tra	egativ , 12, 1 'e abbi ee-ring	re correlé 16, 24, 30 reviated g series c	ttion. Clim 5, 48 mon by their in M <i>Pinus ni</i>	iate par <i>e</i> ths); <i>TN</i> itial lette gra.	umeters: IP, air te ers. $\varDelta_1', .$	<i>PRE</i> , preo mperatur ∆₂′, and ∠	ipitatio e; <i>VPD</i> 3 [,] den	on; <i>R</i> , , air v ote ir	AD, { vapoi ntram	globa ur pre 10lecu	l radić sssure Jlar ¹³	ation; s defici C	t. D

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Table 3 Significance of Pearson's correlations among Δ_1' , Δ_2' , and Δ_3' and climate series for the period 1983–1995 (n = 13).

Table 4 Linear regression models of Δ_1' , Δ_2' , Δ_3' , and ε_{met} as function of ε_{met} , growing season air vapour pressure deficit (VPD) and March–July precipitation (*PRE*).

M1: ${\it \Delta_1'} \sim {\it arepsilon_{ m met}}$ +	<i>VPD</i> , 1983–1995		
$R^2 = 0.87$, adj R^2	$= 0.84, P < 10^{-4}, n = 13$	3	
	Estimate	±SE	P≤
Intercept	36.0	2.7	10 ⁻⁷
emet .	-0.0187	0.0057	0.01
VPD	-0.0295	0.0047	10 ⁻⁴
M2: ${\it \Delta_2'} \sim {\it arepsilon_{ m met}}$ +	+ <i>VPD</i> , 1983–1995		
$R^2 = 0.62$, adj R^2	= 0.54, <i>P</i> < 0.008, <i>n</i> = 1	3	
	Estimate	\pm SE	P≤
Intercept	17.6	4.5	0.003
$\varepsilon_{\rm met}$	-0.0315	0.0097	0.009
VPD	-0.0111	0.0080	0.2
M3: $arDelta_3'\sim arepsilon_{met}$ +	+ VPD, 1983–1995		
$R^2 = 0.64$, adj R^2	= 0.57, <i>P</i> < 0.006, <i>n</i> = 1	3	
	Estimate	\pm SE	P≤
Intercept	14.5	2.1	10^{-4}
ε _{met}	-0.00615	0.00449	0.2
VPD	-0.0129	0.0037	0.006
M4: $arDelta_2'\sim arepsilon_{metr}$	1983–1995		
$R^2 = 0.54$, adj R^2	= 0.50, <i>P</i> < 0.004, <i>n</i> = 1	3	
	Estimate	\pm SE	P≤
Intercept	11.8	1.7	10 ⁻⁴
£ _{met}	-0.0351	0.0097	0.004
M5: ${oldsymbol{\Delta}_{3}}'\sim {oldsymbol{VPD}}$,	1983–1995		
$\frac{\textbf{M5: } \boldsymbol{\Delta}_{3}' \sim \textbf{VPD},}{R^2 = 0.57, \text{ adj}R^2}$	1983–1995 = 0.53, <i>P</i> < 0.003, <i>n</i> = 1	3	
$\frac{\textbf{M5: } \boldsymbol{\Delta_3}' \sim \textbf{VPD,}}{R^2 = 0.57, \text{adj}R^2}$	1983–1995 = 0.53, <i>P</i> < 0.003, <i>n</i> = 1 Estimate	3 ± SE	P≤
$\frac{\textbf{M5: } \Delta_{\textbf{3}}' \sim \textbf{VPD},}{R^2 = 0.57, \text{ adj}R^2}$	1983–1995 = 0.53, <i>P</i> < 0.003, <i>n</i> = 1 Estimate 14.3	3 ± SE 2.2	<i>P</i> ≤ 10 ⁻⁴
$\frac{\textbf{M5: } \boldsymbol{\Delta_3}' \sim \textbf{VPD},}{R^2 = 0.57, \text{ adj}R^2}$ Intercept VPD	1983–1995 = 0.53, <i>P</i> < 0.003, <i>n</i> = 1 Estimate 14.3 -0.0143	3 ± SE 2.2 0.0037	<i>P</i> ≤ 10 ⁻⁴ 0.003
$M5: \Delta_{3}' \sim VPD,$ $R^{2} = 0.57, adjR^{2}$ Intercept VPD $M6: \Delta_{1}' \sim PRE$	1983–1995 = 0.53, <i>P</i> < 0.003, <i>n</i> = 1 Estimate 14.3 -0.0143 + <i>VPD</i> , 1983–1995	13 ± SE 2.2 0.0037	<i>P</i> ≤ 10 ⁻⁴ 0.003
M5: $\Delta_{3}' \sim VPD$, $R^2 = 0.57$, $adjR^2$ Intercept VPD M6: $\Delta_{1}' \sim PRE$ $R^2 = 0.82$, $adjR^2$	1983–1995 = 0.53, <i>P</i> < 0.003, <i>n</i> = 1 Estimate 14.3 -0.0143 + <i>VPD</i> , 1983–1995 = 0.79, <i>P</i> < 0.001, <i>n</i> = 1	2.2 0.0037	<i>P</i> ≤ 10 ⁻⁴ 0.003
M5: $\Delta_{3}' \sim VPD$, $R^{2} = 0.57$, $adjR^{2}$ Intercept VPD M6: $\Delta_{1}' \sim PRE$ $R^{2} = 0.82$, $adjR^{2}$	1983–1995 = 0.53, <i>P</i> < 0.003, <i>n</i> = 1 Estimate 14.3 -0.0143 + VPD, 1983–1995 = 0.79, <i>P</i> < 0.001, <i>n</i> = 1 Estimate	±SE 2.2 0.0037 3 ±SE	<i>P</i> ≤ 10 ⁻⁴ 0.003
M5: $\Delta_{3}' \sim VPD$, $R^2 = 0.57$, $adjR^2$ Intercept VPD M6: $\Delta_{1}' \sim PRE \cdot$ $R^2 = 0.82$, $adjR^2$ Intercept	1983–1995 = 0.53, <i>P</i> < 0.003, <i>n</i> = 1 Estimate 14.3 -0.0143 + VPD, 1983–1995 = 0.79, <i>P</i> < 0.001, <i>n</i> = 1 Estimate 27.8	3 ±SE 2.2 0.0037 3 ±SE 4.4	$P \le 10^{-4}$ 0.003 $P \le 10^{-4}$
M5: $\Delta_{3}' \sim VPD$, $R^2 = 0.57$, $adjR^2$ Intercept VPD M6: $\Delta_{1}' \sim PRE$ $R^2 = 0.82$, $adjR^2$ Intercept PRE	1983–1995 = 0.53, <i>P</i> < 0.003, <i>n</i> = 1 Estimate 14.3 -0.0143 + VPD, 1983–1995 = 0.79, <i>P</i> < 0.001, <i>n</i> = 1 Estimate 27.8 0.0146	3 ±SE 2.2 0.0037 3 ±SE 4.4 0.0061	$P \le 10^{-4}$ 0.003 $P \le 10^{-4}$ 0.04

 $R^2 = 0.43$, $adjR^2 = 0.37$, P < 0.02, n = 13

Table 4 (Continued)

	Estimate	\pm SE	P≤
Intercept PRE	-1.84 0.0274	2.79 0.0096	0.52 0.016

M8: $arDelta_3' \sim VPD$, 1961–1980

 $R^2 = 0.13$, $adjR^2 = 0.07$, P = 0.15, n = 18

	Estimate	\pm SE	P≤
Intercept	12.3	3.8	0.006
VPD	-0.0112	0.0074	0.15

M9: $\Delta_{1}' \sim PRE + VPD$, 1961–1980

 $\overline{R^2} = 0.07$, $\operatorname{adj} R^2 = 0$, P > 0.55, n = 18

	Estimate	±SE	P≤
Intercept	18.6	4.8	0.002
PRE	0.00273	0.00407	0.51
VPD	-0.00357	0.00781	0.65

M10: $\varDelta_{2}' \sim PRE$, 1961–1980

 $R^2 = 0.05$, $adjR^2 = 0$, P = 0.35, n = 18

	Estimate	±SE	P≤
Intercept	6.92	0.93	10 ⁻⁵
PRE	0.00276	0.00286	0.35

M11: $\varepsilon_{met} \sim PRE + VPD$, 1983–1995

 $R^2 = 0.71$, $adjR^2 = 0.66$, P < 0.002, n = 13

	Estimate	±SE	P≤
Intercept	437	118	0.004
PRE .	-0.777	0.164	0.001
VPD	-0.082	0.155	0.61

 ε_{met} , Δ_1' , Δ_2' , and Δ_3' denote hydrogen isotope fractionation caused by metabolic processes at glucose H¹ and H², and carbon isotope discrimination at glucose C-1, C-2, and C-3, respectively. Glucose was extracted across an annually resolved tree-ring series of *Pinus nigra* from the Vienna Basin.

the unequal VPD response of Δ_1 ', Δ_2 ', and Δ_3 ' points to postrubisco fractionations. In the following, we assume Δ_3 ' variation derives entirely from diffusion–rubisco fractionation and argue VPD-dependent isotope fractionation at PGI and G6PD in leaf chloroplasts and the cytosol may exert additional control over Δ_1 ' and Δ_2 ' variability. Generally, variability in PGI fractionation depends on three biochemical properties: (1) the equilibrium status of the PGI reaction, and relative flux of the PGI reactants (2) F6P and (3) G6P into competing metabolic pathways (Figs 2, 5):

(1) PGI reversibly converts F6P into G6P (Fig. 5a). Under nonstress conditions, the PGI reaction in chloroplasts is strongly

	ΔdΛ		PRE		SPE	11	5	SPE1 ₃		SP	EI_4	-1	SPEI ₆		SP	EI_8		SPE	12	S	PEI ₁₆	10	SPE	:124	SPE	:1 ₃₆	SF	EI ₄₈	TMF		SD		RAD
Period/ Δ_i'	1 2	m	1 2	m	-	5	- ~	1 2	m	-	2	 m	2	m	-	5	m	-	5	1 ~	2	<u>س</u>	-	2 3		5	~	2		2 3	1 2	3	1 2
MAMJ	ъ				ъ		σ	_		ъ									10	F		ъ									b		g
ILMAM							с Л	~											10	~		ъ											
ALLMAM																			10	~		ъ											
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SALLM			ы	q			5												10	~		q		ъ									
MJJASO				ъ																		q		ъ									
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averaged for all \geq 4-month periods of the growing season (March–November). Months were abbreviated by their initial letters. Δ_1' , Δ_2' , and Δ_3' denote intramolecular ¹³C discrimination at glucose C-1, C-2, and C-3, respectively. Glucose was extracted across an annually resolved tree-ring series of *Pinus nigra*. Significance levels: a, < 0.05 (light grey); b, < 0.01 (meaium grey). Underline derives negative concluation. Chinate parameters, prosport pressure deficit. Climate data were SPEI, standardised precipitation-evapotranspiration index of different periods (*i* = 1, 3, 6, 8, 12, 16, 24, 36, 48 months); *TMP*, air temperature; VPD, air vapour pressure deficit. Climate data were standardised precipitation-evapotranspiration index of different periods (*i* = 1, 3, 6, 8, 12, 16, 24, 36, 48 months); *TMP*, air temperature; VPD, air vapour pressure deficit. Climate data were standardised precipitation-evapotranspiration index of different periods (*i* = 1, 3, 6, 8, 12, 16, 24, 36, 48 months); *TMP*, air temperature; VPD, air vapour pressure deficit. Climate data were standardised precipitation-evapotranspiration index of different periods (*i* = 1, 3, 6, 8, 12, 16, 24, 36, 48 months); *TMP*, air temperature; VPD, air vapour pressure deficit. Climate data were

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	Δην		PRE		SF	<i>EI</i> 1		SPE	-1 ³		SPEI	-	S	٥Ele		SPE			SPEI.	12	S	PEI ₁₆		SPEI ₂	4	SPE	136	S	9EI ₄₈		TMP		SD		Ч	AD
Period/∆/	4 5	9	4	5 6	4	5	9	4	5	9	4	9	4	5	9	4	5	9	4	2	4	5	9	4 5	9	4	5	4	5	- 7 9	4 5	9	4	5	- 4	5
MAMJ																			а		9													р		2
ILMAM	а			а		а			в												а													q		q
ALLMAM	ы			а		а			а		סי	_																						9		9
MAMJJAS	ы								а		סי	_		а																				9		9
MAMJJASO									в		10	_		а			а																	ы		q
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ſſWÞ	9	ъ		p		U	ъ		Р		10	_									ы									ιu	a	ъ		0		0
AMJJA	9	ъ		р		q	а		р			~		а			а														а	ъ		0	~	0
AMJJAS	ъ	ы		а		р	ъ		Р			^		ъ			ъ															а		U	~	U
AMJJASO	5	а		а		q	ъ		а			^		q			ъ															ъ		q		0
AMJJASON	ъ	а				я	а		ъ		10			а			а															ษ		9		U
ALLM	q	в	ъ	а		q	а		U	а		~		q			а														а	9		0	~ ~	U
MJJAS	5	ы	ы	в		q	ъ		р	ы		a		q			ъ														ъ	9		q	~	q
MJJASO	g	ы	ы	а		q	ъ		q	ы		a		р	ъ		ъ		^v	сц												ъ		q		q
NOSALLM	ъ	а				а	а		ы	а		a		q	ъ		q		.0	сц											ษ	9		ъ		9
JJAS	а	а							q	а		a		q			а		10	с												9			~	ъ
DIASO		ы							q	а		а		q	ъ		ъ		10	е												ы				
NOSALI	а	ъ							а	а	10	l a		q	ъ		q		.0	с												ъ				ъ
IASO		а							а	а		a		q	ъ		q		10	с									а							
IASON		а								а	10	ı a		q	ъ		q	а	.0	5												а				
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sunshine duration; SPE/i, standardised precipitation-evapotranspiration index of different periods (i = 1, 3, 6, 8, 12, 16, 24, 36, 48 months); TMP, air temperature; VPD, air vapour pressure deficit. Climate data were averaged for all \ge 4-month periods of the growing season (March–November). Months were abbreviated by their initial letters. Δ_4' , Δ_5' , and Δ_6' denote intramolecular ¹³C discrimination at glucose C-4, C-5, and C-6, respectively. Glucose was extracted across an annually resolved tree-ring series of *Pinus nigra*.

Table 7 Linear regression models of Δ_4' , Δ_5' , and Δ_6' as function of April–September global radiation (*RAD*), and March–October air temperature (*TMP*).

M1: $\Delta_{5-6}' \sim RAD + TMP$, 1964–1995 $R^2 = 0.72$, $adjR^2 = 0.70$, $P = 10^{-7}$, $n = 28$			
Intercept	26.0	3.1	10 ⁻⁴
RAD	-0.00843	0.00105	10 ⁻⁷
ТМР	1.35	0.29	10 ⁻⁴
$M2: \Delta_5' \sim RAD$	+ <i>TMP</i> , 1964–1995		
$R^2 = 0.66$, adj R^2	$P^2 = 0.64, P = 10^{-6}, n = 28$	3	
	Estimate	±SE	P≤
Intercept	24.8	4.3	10 ⁻⁵
RAD	-0.0103	0.0015	10 ⁻⁶
ТМР	1.81	0.40	10 ⁻⁴
$M3: \Delta_6' \sim RAD$	+ <i>TMP</i> , 1964–1995		
$R^2 = 0.47$, adj R^2	² = 0.43, <i>P</i> < 0.001, <i>n</i> = 2	28	
	Estimate	±SE	P ≤
Intercept	27.3	4.2	10 ⁻⁶
RAD	-0.00658	0.00144	10^{-4}
ТМР	0.876	0.393	0.04
$M4: \Delta_{5-6'} \sim RA$	<i>D</i> + <i>IMP</i> , 1964–1980		
$\frac{\mathbf{M4:} \boldsymbol{\Delta_{5-6'}} \sim \mathbf{RA}}{R^2 = 0.69, \mathrm{adj}R^2}$	<i>D</i> + <i>TMP</i> , 1964–1980 ² = 0.63, <i>P</i> < 0.001, <i>n</i> = 1	15	
$\frac{M4: \Delta_{5-6} \sim RA}{R^2 = 0.69, \text{adj}R^2}$	D + 1MP, 1964–1980 P = 0.63, P < 0.001, n = 1 Estimate	15 ± SE	<i>P</i> ≤
$\frac{M4: \Delta_{5-6} \sim RA}{R^2 = 0.69, \text{ adj}R^2}$ Intercept	$\frac{D + 170P, 1964 - 1980}{P = 0.63, P < 0.001, n = 7}$ Estimate 31.2	15 ± SE 6.7	<i>P</i> ≤ 0.00
$ \frac{M4: \Delta_{5-6} \sim RA}{R^2 = 0.69, \operatorname{adj} R^2} $ Intercept RAD	$\frac{D + 100}{2} = 0.63, P < 0.001, n = 1$ Estimate $\frac{31.2}{-0.00906}$	± SE 6.7 0.00177	<i>P</i> ≤ 0.00 0.00
M4: $\Delta_{5-6} \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP	D + 100000000000000000000000000000000000	± SE 6.7 0.00177 0.47	<i>P</i> ≤ 0.00 0.00 0.04
M4: $\Delta_{5-6}' \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP M5: $\Delta_{5-6}' \sim RA$	D + TMP, 1964-1980 ² = 0.63, P < 0.001, n = 1 Estimate 31.2 -0.00906 1.11 D + TMP, 1983-1995	± SE 6.7 0.00177 0.47	<i>P</i> ≤ 0.00 0.00 0.04
M4: $\Delta_{5-6} \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP M5: $\Delta_{5-6}' \sim RA$ $R^2 = 0.82, adjR^2$	$D + TMP, 1964-1980$ $F = 0.63, P < 0.001, n = 1$ Estimate $31.2 \\ -0.00906 \\ 1.11$ $D + TMP, 1983-1995$ $F = 0.79, P < 0.001, n = 1$	15 ± SE 6.7 0.00177 0.47	<i>P</i> ≤ 0.00 0.00 0.04
M4: $\Delta_{5-6}' \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP M5: $\Delta_{5-6}' \sim RA$ $R^2 = 0.82, adjR^2$	D + 1/MP, 1964-1980 ² = 0.63, P < 0.001, n = 1 Estimate 31.2 -0.00906 1.11 D + 1/MP, 1983-1995 ² = 0.79, P < 0.001, n = 1 Estimate	15 ±SE 6.7 0.00177 0.47	<i>P</i> ≤ 0.00 0.00 0.04 <i>P</i> ≤ <i>P</i> ≤
M4: $\Delta_{5-6}' \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP M5: $\Delta_{5-6}' \sim RA$ $R^2 = 0.82, adjR^2$ Intercept	$D + TMP, 1964-1980$ $P^{2} = 0.63, P < 0.001, n = 1$ Estimate $31.2 \\ -0.00906 \\ 1.11$ $D + TMP, 1983-1995$ $P^{2} = 0.79, P < 0.001, n = 1$ Estimate 29.1	15 ±SE 6.7 0.00177 0.47 13 ±SE 4.5	<i>P</i> ≤ 0.00 0.00 0.04 <i>P</i> ≤ 10 ⁻⁴
M4: $\Delta_{5-6}' \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP M5: $\Delta_{5-6}' \sim RA$ $R^2 = 0.82, adjR^2$ Intercept RAD	$D + TMP, 1964-1980$ $P^{2} = 0.63, P < 0.001, n = 1$ Estimate $31.2 \\ -0.00906 \\ 1.11$ $D + TMP, 1983-1995$ $P^{2} = 0.79, P < 0.001, n = 1$ Estimate $29.1 \\ -0.00875$	15 ±SE 6.7 0.00177 0.47 13 ±SE 4.5 0.00132	$P \le$ 0.00 0.00 0.04 $P \le$ 10^{-4} 10^{-4}
M4: $\Delta_{5-6} \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP M5: $\Delta_{5-6} \sim RA$ $R^2 = 0.82, adjR^2$ Intercept RAD TMP	$D + 7MP, 1964-1980$ $P = 0.63, P < 0.001, n = 1$ Estimate $31.2 \\ -0.00906 \\ 1.11$ $D + 7MP, 1983-1995$ $P = 0.79, P < 0.001, n = 1$ Estimate $29.1 \\ -0.00875 \\ 1.22$	15 ±SE 6.7 0.00177 0.47 13 ±SE 4.5 0.00132 0.40	$P \le 0.00 \\ 0.00 \\ 0.04 \\ P \le 10^{-4} \\ 10^{-4} \\ 0.01 \\ 0.01 \\ 0.00 \\ $
M4: $\Delta_{5-6}' \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP M5: $\Delta_{5-6}' \sim RA$ $R^2 = 0.82, adjR^2$ Intercept RAD TMP M6: $\Delta_4' \sim RAD$	D + 17MP, 1964–1980 ² = 0.63, P < 0.001, n = 1 Estimate 31.2 -0.00906 1.11 D + 17MP, 1983–1995 ² = 0.79, P < 0.001, n = 1 Estimate 29.1 -0.00875 1.22 + 17MP, 1964–1995	15 ±SE 6.7 0.00177 0.47 13 ±SE 4.5 0.00132 0.40	$P \le$ 0.00 0.04 $P \le$ $P \le$ 10^{-4} 0.01
M4: $\Delta_{5-6}' \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP M5: $\Delta_{5-6}' \sim RA$ $R^2 = 0.82, adjR^2$ Intercept RAD TMP M6: $\Delta_4' \sim RAD$ $R^2 = 0.15, adjR^2$	$D + 17MP, 1964-1980$ $P^{2} = 0.63, P < 0.001, n = 1$ Estimate $31.2 \\ -0.00906 \\ 1.11$ $D + 17MP, 1983-1995$ $P^{2} = 0.79, P < 0.001, n = 1$ Estimate $29.1 \\ -0.00875 \\ 1.22$ $+ 17MP, 1964-1995$ $P^{2} = 0.09, P = 0.12, n = 28$	15 ±SE 6.7 0.00177 0.47 13 ±SE 4.5 0.00132 0.40	$P \le 0.00 \\ 0.00 \\ 0.04 \\ P \le 10^{-4} \\ 10^{-4} \\ 0.01 \\ $
M4: $\Delta_{5-6}' \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP M5: $\Delta_{5-6}' \sim RA$ $R^2 = 0.82, adjR^2$ Intercept RAD TMP M6: $\Delta_4' \sim RAD$ $R^2 = 0.15, adjR^2$	$D + 17MP, 1964-1980$ $P^{2} = 0.63, P < 0.001, n = 1$ Estimate $31.2 -0.00906 - 1.11$ $D + TMP, 1983-1995$ $P^{2} = 0.79, P < 0.001, n = 1$ Estimate $29.1 -0.00875 - 1.22$ $+ TMP, 1964-1995$ $P^{2} = 0.09, P = 0.12, n = 28$ Estimate	±SE 4.5 0.00177 0.47 13 ±SE 4.5 0.00132 0.40 	$P \le 0.00 \\ 0.00 \\ 0.04 \\ P \le 10^{-4} \\ 10^{-4} \\ 0.01 \\ P \le P \le 0$
M4: $\Delta_{5-6}' \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP M5: $\Delta_{5-6}' \sim RA$ $R^2 = 0.82, adjR^2$ Intercept RAD TMP M6: $\Delta_4' \sim RAD$ $R^2 = 0.15, adjR^2$ Intercept	$D + TMP, 1964-1980$ $P^{2} = 0.63, P < 0.001, n = 1$ Estimate $31.2 -0.00906 - 1.11$ $D + TMP, 1983-1995$ $P^{2} = 0.79, P < 0.001, n = 1$ Estimate $29.1 -0.00875 - 1.22$ $+ TMP, 1964-1995$ $P^{2} = 0.09, P = 0.12, n = 28$ Estimate 8.33	15 \pm SE 6.7 0.00177 0.47 13 \pm SE 4.5 0.00132 0.40 5 \pm SE \pm SE 4.73	$P \le 0.00 \\ 0.04 \\ 0.04 \\ P \le 10^{-4} \\ 10^{-4} \\ 0.01 \\ P \le 0.09 \\ 0.09 \\ 0.09 \\ 0.00 \\ 0.0$
M4: $\Delta_{5-6}' \sim RA$ $R^2 = 0.69, adjR^2$ Intercept RAD TMP M5: $\Delta_{5-6}' \sim RA$ $R^2 = 0.82, adjR^2$ Intercept RAD TMP M6: $\Delta_4' \sim RAD$ $R^2 = 0.15, adjR^2$ Intercept RAD	D + 1/MP, 1964-1980 $P = 0.63, P < 0.001, n = 1$ Estimate $31.2 -0.00906 -1.11$ $D + T/MP, 1983-1995$ $P = 0.79, P < 0.001, n = 1$ Estimate $29.1 -0.00875 -1.22$ $+ T/MP, 1964-1995$ $P = 0.12, n = 28$ Estimate $8.33 -0.00266$	$ \begin{array}{c} 15 \\ \pm SE \\ 6.7 \\ 0.00177 \\ 0.47 \\ 13 \\ \pm SE \\ 4.5 \\ 0.00132 \\ 0.40 \\ \hline \\ 5 \\ \pm SE \\ 4.73 \\ 0.00160 \\ \end{array} $	$P \le 0.00 \\ 0.00 \\ 0.04 \\ P \le 10^{-4} \\ 10^{-4} \\ 0.01 \\ P \le 0.09 \\ 0.11 \\ 0.01 \\ 0.09 \\ 0.11 \\ 0.01 \\ 0.01 \\ 0.00 \\ 0.0$

 Δ_4' , Δ_5' , and Δ_6' denote carbon isotope discrimination at glucose C-4, C-5, and C-6, respectively. Glucose was extracted across an annually resolved tree-ring series of *Pinus nigra* from the Vienna Basin.

displaced from equilibrium on the side of F6P (Dietz, 1985; Gerhardt et al., 1987; Kruckeberg et al., 1989; Schleucher et al., 1999; Wieloch, 2022; Wieloch et al., 2022a). With decreasing p_i , however, the reaction moves towards equilibrium (Dietz, 1985; Wieloch et al., 2022a). This shift is accompanied by ¹³C increases at C-1 and C-2 of G6P (Table 2), which will be transmitted to downstream derivatives such as starch and treering glucose (Wieloch et al., 2018). In isohydric species such as *P. nigra*, p_i decreases with drought due to stomatal closure (McDowell et al., 2008). Here, we found stronger VPD-induced ¹³C increases at tree-ring glucose C-1 than at C-3. This is consistent with the PGI-related isotope shift expected at C-1. However, the apparent absence of the diffusion-rubisco signal from C-2 contrasts with the expected isotope shift. That said, in Phaseolus vulgaris, the ratio of leaf sucrose-to-starch carbon partitioning was shown to increase steeply with decreasing p_i (Sharkey et al., 1985). Hence, the relative contribution of chloroplastic G6P and its isotope composition to downstream metabolism may decrease with increasing VPD, reducing the influence of the mechanism described on Δ_1 ' and Δ_2 ' variation.

(2) In natural systems, leaf night-time respiration is increased under drought (Fig. 5b; Schmiege *et al.*, 2023). Furthermore, in the dark, the cytosolic PGI reaction was found to be near equilibrium (Gerhardt *et al.*, 1987). Consequently, F6P would be ¹³C depleted at C-1 but ¹³C enriched at C-2 relative to the corresponding G6P positions (Table 2). Increasing relative F6P flux into mitochondrial respiration would then result in ¹³C increases at C-1 and ¹³C decreases at C-2 of G6P and downstream derivatives. Thus, this mechanism is consistent with both observations, stronger *VPD*-induced ¹³C increases at tree-ring glucose C-1 compared with C-3, and the apparent absence of the diffusion– rubisco signal from C-2.

(3) While carbon assimilation commonly decreases with drought (McDowell et al., 2008), the activity of leaf-cytosolic G6PD increases (Fig. 5c; Landi et al., 2016). This can be expected to result in increasing relative G6P flux into the OPPP. While some authors reported that the cytosolic PGI reaction in illuminated leaves is in equilibrium (Gerhardt et al., 1987), others found displacements from equilibrium (Leidreiter et al., 1995; Schleucher et al., 1999; Szecowka et al., 2013). Hence, PGIrelated isotope shifts in tree-ring glucose resulting from G6P flux into the leaf-cytosolic OPPP are hard to predict (Table 2). By contrast, the unidirectional conversion of G6P to 6phosphogluconolactone catalysed by G6PD proceeds faster with ¹²C-1 than ¹³C-1 G6P $(k_{12C}/k_{13C} = 1.0165)$ (Hermes et al., 1982). Hence, increasing relative flux through the leafcytosolic OPPP may contribute to the stronger VPD-induced ¹³C increases at tree-ring glucose C-1 compared with C-3.

Aside from these mechanisms, there are others that might introduce Δ_1 ' and Δ_2 ' variation. For instance, we recently reported evidence consistent with increasing relative flux through the chloroplastic OPPP in response to decreasing p_i under illumination (Fig. 5a; Wieloch *et al.*, 2022a, 2023). Furthermore, under illumination, chloroplastic F6P is used for both RuBP regeneration and starch biosynthesis (Fig. 5a). Increasing *VPD*



promotes photorespiration resulting in increasing RuBP regeneration relative to carbon export from the Calvin–Benson cycle into sinks such as starch. **Fig. 5** Processes invoked to explain isotope fractionation at tree-ring glucose HC-1 and HC-2: (a) in leaf chloroplasts under illumination, (b) in the leaf cytosol in the dark, (c) in the leaf cytosol under illumination, and (d) in the stem cytosol. F6P and G6P carbon atoms 1–6 occur in sequentially order from top to bottom. Atom positions affected by G6PD and PGI fractionation are given in blue and green, respectively. In some cases, carbon position 1 is given both as blue letter and green dot to indicate fractionation at both enzymes. Dashed arrows indicate that intermediate reactions are not shown. Wavy lines indicate fractional introduction of hydrogen from water by the PGI reaction. Note that G6PD in stem leucoplasts may additionally contribute to isotope fractionation at tree-ring glucose C-1 and H¹. F6P, fructose 6-phosphate; G6PD, G6P dehydrogenase; OPPP, oxidative pentose phosphate pathway; PGI, phosphoglucose isomerase; p_i , intercellular CO₂ partial pressure; *PRE*; precipitation; RuBP, ribulose 1,5-bisphosphate.

The mechanisms described above should also introduce hydrogen isotope signals because of the hydrogen isotope effects of G6PD affecting G6P H¹ (Hermes et al., 1982) and PGI affecting G6P H² (Table 2; Fig. 5). However, growing season VPD neither correlates with ε_{met} pertaining to tree-ring glucose H¹ nor H² (Tables \$5, \$6). Hence, either G6PD and PGI are not the sources of VPD-dependent carbon isotope fractionation in Δ_1 ' (and Δ_2), or the corresponding hydrogen isotope signals were washed out after introduction. Washout at H¹ may occur during equilibration of F6P with mannose 6-phosphate by phosphomannose isomerase (cf. Topper, 1957). Similarly, complete washout at H² may occur when the leaf-cytosolic PGI reaction is in equilibrium (Notes S3). Previously, this latter process was invoked (among others) to explain why a whole-molecule deuterium depletion observed in leaf starch was not transmitted to nocturnal sucrose (see the Introduction section; Wacker, 2022). As each conversion by PGI was found to be associated with a 0-50% probability for hydrogen exchange with the medium (Noltmann, 1972), partial washout of existing hydrogen isotope signals may also occur under nonequilibrium conditions (Notes S3).

In the mechanisms described above, we assumed diffusion– rubisco fractionation contributes to VPD-dependent Δ_i ' variation. However, diffusion–rubisco fractionation affects all glucose carbon positions equally (Wieloch *et al.*, 2018). Since merely two of six glucose carbon positions carry VPD-dependent isotope variation, the question arises of whether the diffusion–rubisco signal was already below the detection level on introduction. If this were the case, then VPD-dependent Δ_1 ' and Δ_3 ' variation would originate entirely from post-rubisco processes. Furthermore, postrubisco processes that were previously invoked to explain the absence of the diffusion–rubisco signal from C-4, C-5, and C-6 (see the Introduction section) would not occur.

Isotope fractionation mechanisms in stems affecting treering glucose HC-1 and HC-2

Previously, we found a change point in ε_{met} in 1980 (Wieloch *et al.*, 2022b). Here, we found the same change point in Δ_{1-2} ' (see the Results section). Consistent with this, Δ_1 ' and Δ_2 ' data of 1983–1995 exhibit a significantly lower average value and a

significantly larger variance than those of 1961–1980 (Tables S1, S2). Furthermore, Δ_1 ' and Δ_2 ' are significantly correlated during the late (Fig. 4a) but not the early period (Fig. 4b), and ε_{met} accounts for a significant fraction of the variance of both Δ_1 ' and Δ_2' during the late period (Table 4 (M1, M4); Fig. 3a,b). In Δ_2 ', the $\varepsilon_{\rm met}$ effect is about twice as large as in Δ_1 '. Processes affecting ε_{met} , Δ_1 ', and Δ_2 ' simultaneously respond to *PRE* but not VPD (Tables 4 (M1, M2, M4, M6, M7, M11), S_5). Δ_1 ' and Δ_2 ' respond to March–July *PRE* during the late but not the early period (Table 4 (M9-M10)). Previously, we reported evidence suggesting the groundwater table before 1980 was high enough to prevent metabolic changes causing ε_{met} variation (Wieloch et al., 2022b). By extension, this should also explain the properties of Δ_1 ' and Δ_2 ' listed above. That is, since the trees had access to groundwater during the early period, metabolic shifts that can cause intercorrelated variation in ε_{met} , Δ_1 ', and Δ_2 ' were not induced.

Processes causing intercorrelated variation in ε_{met} , Δ_1 ', and Δ_2 ' are probably located in the stem (see the first section of the Discussion section). The ε_{met} signal is present at glucose H¹ and, considerably more strongly, at H² (range: 64‰ and 240‰, respectively; 1983–1995). In the biochemical pathway leading to tree-ring cellulose, PGI is the last enzyme acting on precursors of glucose H² (Figs 2, 5d). With each conversion by PGI, there is a probability for hydrogen exchange with the medium of 0–50% (Noltmann, 1972). Thus, if we assume *P. nigra* stem PGI exchanges hydrogen with the medium as does spinach leaf PGI (Fedtke, 1969) and the reaction is in equilibrium, then any deuterium signal at G6P H² will be washed out. Among all H-C positions of tree-ring glucose, the deuterium abundance at H² is neither exceptionally high nor low during 1961–1980, whereas it is exceptionally high (and exceptionally variable) during



Fig. 6 Average intramolecular δD_i patterns of the periods 1961–1980 and 1983–1995 (black and blue, respectively). The data were acquired for tree-ringglucose of *Pinus nigra* laid down at a site in the Vienna basin. $\delta D_i = D_i / (\Sigma D_{ME}/6)$ -1 where D_i denotes the deuterium abundance at glucose hydrogen position *i* and $\Sigma D_{ME}/6$ denotes the average deuterium abundance of the six methyl-group hydrogens of the glucose derivative used for NMRS measurements. Error bars represent 95% confidence intervals. The figure shows discrete data. Dashed and dotted lines were added to guide the eye. Modified figure from Wieloch *et al.* (2022b).

1983-1995 (Fig. 6). This indicates that the PGI reaction was close to or in equilibrium during 1961-1980 but displaced from equilibrium on the side of G6P during 1983-1995 (Table 2). Additionally, shifts of the PGI reaction away from equilibrium towards the side of G6P should cause ¹³C enrichment at G6P C-1 and C-2 (Δ_1 ' and Δ_2 ' decreases), and Δ_2 ' should decrease three times more than Δ_1 ' (Table 2). Consistent with this, we found negative relationships between $\varepsilon_{
m met}$ and Δ_1 ', as well as Δ_2 ' (Table 4 (M1 and M4)). However, Δ_2 ' decreases only 1.88 times more than Δ_1 ', but this best estimate is associated with a relatively large error (SE interval: 1.04-3.45). That said, the offset from 3 is likely explained by increasing relative flux through the OPPP accompanying the putative PGI reaction shift (Figs 2, 5d). This is because G6P to 6-phosphogluconolactone conversion by G6PD exhibit ¹³C and D isotope effects $(k_{12C}/k_{13C} = 1.0165, k_H/k_D = 2.97)$ (Hermes *et al.*, 1982). Hence, increasing relative OPPP flux causes ¹³C enrichment at G6P C-1 (Δ_1 ' decreases) and deuterium enrichment at G6P H¹. This is consistent with both the apparently decreased PGI effects ratio (1.88 instead of 3) and, more importantly, ε_{met} increases at glucose H¹ of up to 64‰ during 1983–1995 (Fig. 6).

Sucrose translocated from leaves can be split into UDP-glucose and fructose via sucrose synthase or glucose and fructose via invertase (Fig. 2). UDP-glucose entering tree-ring cellulose biosynthesis directly via sucrose synthase is protected from isotope fractionation by PGI and G6PD. However, in stems of juvenile Quercus petraea and Picea abies, at least 79% and 43% of the precursors of tree-ring glucose went through PGI catalysis, respectively (Augusti et al., 2006). Theoretically, shifts of the PGI reaction away from equilibrium towards the side of G6P can cause ε_{met} increases at glucose H² of up to 611‰ (Notes S3, hydrogen exchange with the medium not considered). With 43% and 79% of all precursors of tree-ring glucose undergoing PGI catalysis, ε_{met} increases at glucose H² of up to 263‰ and 483‰ are possible, respectively. Thus, the PGI-related fractionation mechanism proposed here can potentially cause previously reported ε_{met} increases at glucose H² of up to 240‰ (Wieloch et al., 2022b). Shifts in sucrose cleavage by sucrose synthase vs invertase may exert additional control over the ε_{met} signal at glucose H^2 .

Based on results and interpretations presented above, decreasing stem water content is associated with both increasing OPPP flux and a shift of the PGI reaction away from equilibrium towards the side of G6P corresponding to low relative F6P concentration (Fig. 5d). We propose these concerted shifts may ensure redox homeostasis and balanced substrate supply to glycolysis as follows. In heterotrophic tissue, NADPH from the OPPP is believed to be central for maintaining redox homeostasis (Fig. 2; Stincone et al., 2015). Flux through the OPPP is regulated at G6PD. Heterotrophic G6PD activity reportedly increases with drought (Liu et al., 2013; Wang et al., 2016, 2020), oxidative load (Wang et al., 2016, 2020; Li et al., 2020), NADPH demand (Wendt et al., 2000; Esposito et al., 2001; Castiglia et al., 2015), and abscisic acid concentration (Cardi et al., 2011; Wang et al., 2016). Decreasing stem water content may cause increasing OPPP flux via increasing abscisic acid concentration (Brunetti et al., 2020), and possibly increasing

oxidative load increasing the demand for NADPH. In turn, increasing OPPP flux results in increasing supply of pentose phosphates, which may feed into glycolysis via the reductive part of the pentose phosphate pathway (Figs 2, 5d). This would reduce the demand for glycolytic substrates supplied via PGI. The shift of the PGI reaction away from equilibrium towards the side of G6P may reflect this decreased demand and result from PGI downregulation by intermediates of the pentose phosphate pathway such as erythrose 4-phosphate, ribulose 5-phosphate, and 6-phosphogluconate (Parr, 1956; Grazi et al., 1960; Salas et al., 1965). Furthermore, relative changes in G6P-to-F6P supply vs consumption may contribute to the shift of the PGI reaction. For instance, while starch storage consumes G6P, remobilisation supplies G6P (Noronha et al., 2018). Under drought, the storage-to-remobilisation balance may tilt towards remobilisation (Mitchell et al., 2013; Thalmann & Santelia, 2017; Tsamir-Rimon et al., 2021). Consequently, the PGI reaction may move towards the side of G6P. Similarly, we previously reported below-average tree-ring widths for years in which the PGI reaction is on the side of G6P (Wieloch et al., 2022b). Hence, in these years, G6P consumption by growth may have been reduced while F6P consumption by downstream metabolism may have been maintained.

1961–1980: A period of homeostasis with respect to processes affecting Δ_1' , Δ_2' , Δ_3' , and ε_{met}

During 1961–1980, Δ_1 ', Δ_2 ', and Δ_3 ' are not significantly correlated which contrasts with the period 1983-1995 (Fig. 4a,b). Similarly, relationships of Δ_1 ' and Δ_3 ' with VPD and Δ_1 ' and Δ_2 ' with *PRE* observed during the late period are largely absent during the early period (Tables 3-5) even though there is no difference in the magnitude of VPD and PRE variability between these periods (Fig. S1). Like Δ_{1-2}' and Δ_{1-3}' , ε_{met} exhibits a change point in 1980 and responds to PRE after but not before 1980 (Wieloch et al., 2022b). This shift in ε_{met} sensitivity was attributed to long-term drought, which intensified over the study period and proposedly lead to a groundwater depletion below a critical level in 1980 (Wieloch et al., 2022b). By extension, this groundwater depletion might also explain the insensitivity of Δ_1 ' and Δ_3 ' to VPD and Δ_1 ' and Δ_2 ' to PRE during 1961–1980 and their sensitivity from 1983 onwards. Thus, while the trees had access to groundwater, leaf- and stem-level processes affecting $\Delta_1', \Delta_2', \Delta_3'$ and $\varepsilon_{\rm met}$ could apparently maintain homeostasis despite changing atmospheric conditions.

Isotope fractionation mechanisms in leaves affecting tree-ring glucose C-5 and C-6

No change points were detected in Δ_5 ' and Δ_6 ' (see the Results section; Tables S1, S2). Furthermore, Δ_5 ' and Δ_6 ' remain significantly correlated across the entire study period (Figs 1b, 4a,b), and *RAD* is the most influential environmental cofactor (Table 6). Models including *RAD* and *TMP* as cofactors capture most of the systematic variance in Δ_{5-6} ', Δ_5 ', and Δ_6 ' (Tables 7 (M1–M3), S4). These relationships hold for both the early and

late study period (Table 7 (M4–M5)) with Δ_5 ' effects being c. 1.5-fold larger than Δ_6 ' effects (M2 vs M3, SE interval: 1.1–2.28).

Previously, we reported a negative relationship between treering glucose Δ_{5-6} ' and reconstructed tropospheric O₃ concentration (see the Introduction section; Wieloch *et al.*, 2022c). Light stimulates tropospheric O₃ formation (Lu *et al.*, 2019). This may explain the negative relationship between tree-ring glucose Δ_{5-6} ' and *RAD* reported here (Table 7 (M1–M3)). Furthermore, we previously explained the absence of the diffusion–rubisco signal from glucose C-5 and C-6 (*inter alia*) by interaction between photorespiration and the TCAC (see the Introduction section; Wieloch *et al.*, 2022c). As *TMP* increases, photorespiration increases more than photosynthesis (Long, 1991). This may result in decreasing flux of PEP into the TCAC (see the Introduction section) and explain the positive relationship between treering glucose Δ_{5-6} ' and *TMP* reported here (Table 7 (M1–M3)).

Isotope fractionation mechanisms in leaves affecting tree-ring glucose C-4

As for Δ_5 ' and Δ_6 ', no change point was detected in Δ_4 ' (see the Results section; Tables S1, S2). Considering the entire study period, Δ_4 ' is weakly associated with Δ_{5-6} ' (Fig. 1b). Consistent with this, the Δ_{5-6} '-climate model works reasonably well for Δ_4 ' considering the relatively low systematic variance in Δ_4 ' of 38% (Tables 7 (M1 and M6), S4). Introduction of the Δ_4 ' and Δ_{5-6} ' signals proposedly involves leaf-level consumption of PGA and PEP by downstream metabolism, respectively (Wieloch *et al.*, 2021, 2022c). Since PGA is a precursor of PEP (Fig. 2), our previously proposed theories of signal introduction are in line with the observation that Δ_4 ', Δ_5 ' and Δ_6 ' are associated and respond to the same environmental parameters.

Conclusions and future directions

Dual-isotope-environment analysis was used to deconvolute isotope signals and provide several new insights into plant isotope fractionation. First, the diffusion-rubisco signal was previously shown to be absent from tree-ring glucose C-4 to C-6 (Wieloch et al., 2021, 2022c) but believed to be present at C-1 to C-3 (Wieloch et al., 2018). Here, this signal was found to also be absent from C-2. Second, isotope fractionation beyond leaves is commonly considered to be constant for any given species (Roden et al., 2000; Gagen et al., 2022). However, our results suggest a significant part of the carbon and hydrogen isotope variation in tree-ring glucose originates in stems from processes affecting Δ_1 ', Δ_2 ', and ε_{met} simultaneously. Third, VPD affects Δ_1 ' and Δ_3 ' and *PRE* affects Δ_1 ', Δ_2 ', and ε_{met} (Table 4). These relationships proposedly reflect water content variability in leaves and stems, respectively. They apply to the late but not the early study period consistent with the finding of a change point in both the ε_{met} (Wieloch *et al.*, 2022b) and Δ_{1-3} ' series. This change point proposedly marks the crossing of a physiologically relevant groundwater threshold (Wieloch et al., 2022b). Additionally, we reported Δ_{5-6} ' relationships with RAD and

TMP, which apply to the entire study period (Table 4). These latter relationships are consistent with previously proposed isotope fractionation mechanisms (Wieloch *et al.*, 2022c). By contrast, we here revised and expanded our previous theory on the mechanisms introducing Δ_1 ', Δ_2 ', Δ_3 ', and ε_{met} variability. Given the multitude of isotope-environment relationships (including change point responses), intramolecular carbon isotope analysis has a remarkable potential for reconstructions of environmental conditions (*VPD*, *PRE*, *RAD*, *TMP*, soil moisture, groundwater thresholds, tropospheric O₃ concentration), tissue water content (leaf, stem), metabolic flux variability (various processes), and ecophysiological properties such as intrinsic water-use efficiency across space and time. Complementing hydrogen isotope analysis is expected to significantly enhance these capabilities.

Understanding isotope fractionation mechanisms is central for retrospective studies of plant physiology and climate based on tree-ring isotope data, and there is considerable room for improvement as shown above. Research in several largely unexploited areas is needed to make progress. First, there is a basic need for more in vitro data on intramolecular isotope effects of enzyme reactions including the reactions catalysed by triosephosphate isomerase, transketolase, PEPC, PK, and DAHPS. Second, intramolecular isotope analyses of leaf metabolites including starch and sucrose from both controlled and natural environments are needed to generate a baseline for mechanistic studies of isotope fractionation along the pathway from leaves to wood. Additionally, intramolecular isotope analysis of metabolites from wood slices acclimated to different ambient conditions (e.g. wet vs dry, varying sucrose supply) will be insightful. Third, combined analysis of intramolecular ¹³C and deuterium data can help to separate isotope signals. Fourth, genetic modification of key enzymes may help to test proposed isotope fractionation mechanisms in vivo. Fifth, intramolecular isotope fractionation affecting tree-ring glucose is complex. Software programs such as QIRN enable the convenient simulation of natural isotope abundances in complex metabolic networks (Mueller et al., 2022). If expanded, these programs may help to extract metabolic information from intramolecular tree-ring isotope data. This would require routines enabling control of relative flux at metabolic branchpoints by optimising regressions between (1) relative branchpoint flux and environmental parameters and (2) simulated and observed isotope data. In summary, intramolecular isotope analysis has an enormous potential to advance our knowledge about isotope fractionation mechanisms, plant ecophysiology, and paleoclimatology.

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Competing interests

None declared.

Author contributions

Conceptualisation, visualisation, project administration, and development of isotope theory were done by TW. Investigation and writing were done by TW with input from MH-P, JY, and TN.

ORCID

Meisha Holloway-Phillips D https://orcid.org/0000-0002-8353-3536

Totte Niittylä D https://orcid.org/0000-0001-8029-1503 Thomas Wieloch D https://orcid.org/0000-0001-9162-2291 Jun Yu D https://orcid.org/0000-0001-5673-620X

Data availability

Isotope data used here were published previously (Wieloch *et al.*, 2018, 2022b). Climate data used here are publicly accessible (see the Materials and Methods section). Data derivatives supporting the findings of this study are available within the paper (see the Results section; Tables 3–7; Figs 3, 4, 6) and its supporting information (see Notes S2, S3; Tables S1–S6; Fig. S1).

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Supporting Information

Additional Supporting Information may be found online in the Supporting Information section at the end of the article.

Fig. S1 Air vapour pressure deficit of the growing season and March–July precipitation over the period from 1961 to 1995 in the Vienna basin.

Notes S1 Materials and Methods (expanded).

Notes S2 Hydro-carbon isotope fractionation from 1961 to 1980.

Notes S3 Estimated deuterium fractionation due to shifts of the phosphoglucose isomerase reaction.

Table S1 Shapiro-Wilk normality test.

Table S2 F and T test.

Table S3 Pearson's correlations between Δ_i' and ε_{met} series of the period from 1983 to 1995.

Table S4 Components of variance in Δ_i ' series.

Table S5 Pearson's correlation coefficients and associated levels of significance of ε_{met} -climate relationships for the period from 1983 to 1995.

Table S6 Linear regression model of $\varepsilon_{met (H1)}$ as function of growing season air vapour pressure deficit and March–July precipitation.

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